### Lakes Watershed Simulation and Forecasting System WSFS www.environment.fi/waterforecast

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#### SYKE

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#### Lake surface temperature

Lake surface temperature to improve lake evaporation simulation
Simple equation by Lindström et.al. (\*):

$$T(t) = (1-k) \cdot T(t-1) + k \cdot T_{air}(t)$$

# *k* is a parameter, calibrated using observations of lake surface temperature

\*) Lindström, G., Gardelin, M. and Persson, M. 1994. *Conceptual modeling of evapotranspiration for simulations of climate change effects*. SMHI RH No. 10, 25 p

#### **Improved model**

- Simple model poor in midsummer: too low temperatures
- Supposed to be mainly due to radiation energy from the sun, and was added as follows:

$$T(t) = (1 - k) \cdot T(t - 1) + k \cdot (T_{air}(t) + r(t) \cdot l)$$

r(t) = intensity of the short-wave radiation
I = influence of the short-wave radiation on lake surface warming

#### Implementation

- Parameters calibrated using lake surface observations over a long (30-40 y) period for 30 lakes.
- Additionally there is a large number of observation points in rivers for some of which the same model is calibrated and used operationally
- The lake and river temperature results are presented as graphs and as maps.

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Järven pintalämpötila ( $^{\circ}C$ ) 1. 8.2012 Forecast day 1. 8. Lake surface temperature ( $^{\circ}C$ ) 1. 8.2012

### Simulated lake surface temperature 1.8.2012

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### Simulated lake surface temperature 7.9.2012



Järven pintalämpötila (°C) 7. 9.2012 Forecast day 7. 9. Lake surface temperature (°C) 7. 9.2012



**Observations and** observation sites 7.9.2012

### Lake Saimaa, surface temperature





Lake evaporation

- Input: lake surface temperature , short wave radiation index and precipitation
- Short-wave radiation by latitude and date
- Precipitation included to indicate cloudiness and high relative humidity

Approach is more statistical than physical.

- Method was used to simulate potential evaporation measured by Class A –pans.
- Class A evaporation calculated by air temperature, short-wave radiation and precipitation; calibrated against long -time observation series
- Same dependencies were applied to lake evaporation, replacing air temperature by lake surface temperature.

### **Physical lake evaporation model**

• Dalton's equation (Aerodynamic formula)

$$Q_e = \rho_a C_e (q_s - q_a) U_a$$

- q<sub>s</sub> and q<sub>a</sub> are specific humidities at lake surface and air 2m level
- U<sub>a</sub> is wind speed
- C<sub>e</sub> is bulk transfer coefficient
- P<sub>a</sub> is air density
- Affecting factors
  - Air temperature
  - Lake surface temperature
  - Wind speed
  - Relative humidity
  - Near surface (atmospheric) stratification
    - Difference between air and lake surface temperature
    - Wind speed
    - Surface roughness

### **Bulk transfer coefficient (C**<sub>e</sub>)

Near surface atmospheric stratification according to Launiainen (1995, fig 4.)

- Richardson number (R<sub>z</sub>)

$$R_{Z} = \frac{g}{\theta_{s}} z_{a} \frac{\theta_{va} - \theta_{vs}}{U_{a}^{2}},$$

- $\Theta_s$ ,  $\Theta_{va}$  and  $\Theta_{vs}$  are potentia temperature at surface and virtual potential temperatures at surface and in air
- Surface roughness
   at lake surface z<sub>a</sub> 4\*10<sup>-4</sup>
   (Neutral case 1.56\*10<sup>-3</sup>)



Fig. 4. Drag coefficient and bulk heat transfer coefficient (referred to 10 m) for diabatic cases. The continuous and dashed lines depict the  $C_D/C_{DN}$  and  $C_H/C_{HN}$  dependences based on iteration, while the crosses are those calculated according to (A5)–(A7) using the  $\zeta$  from Equations (9) and (11). The roughness lengths used,  $z_0 = 10^{-4}$  m and  $z_0/z_T = 1$ , correspond to the neutral case  $(C_{DN} = C_{HN} = 1.2 \times 10^{-3})$ .

#### Saimaa lake evaporation

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Järvihaihdunta Lake evaporation mm



#### Old method

- Large evaporation in Spring ~ 1-3 mm/d in May
- Small evaporation in Autumn < 1.5 mm/d in September and October

#### New method

- Summer and Autumn: **Evaporation increases** ~ 2-3 mm/d
- Spring: Decreases ~ 2 mm/d

NOTE! Preliminary results, limited amount of wind speed and humidity observations



### **Comparison of lake evaporation equations**

#### **Evaporation equations**

New method =  $\rho_a * C_e * (q_0 - q_2) * U_{10}$ Equation 2 = (0.15+0.108\*U<sub>2</sub>)(e<sub>0</sub>-e<sub>2</sub>) Equation 3 = (0.236+0.068\*U<sub>2</sub>)(e<sub>0</sub>-e<sub>2</sub>) Equation 4 = (0.127+0.16\*U<sub>2</sub><sup>0.62</sup>) (e<sub>0</sub>-e<sub>2</sub>) \*Equations 2-4 from Järvinen and Huttula, 1982



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## Ice thickness model

- Calculates ice thickness at observation points
  - All lake and river points in Finland
  - Simulates black ice, snow ice, slush and snow on ice
- Ice growth is driven by
  - Air temperature
  - Water temperature (beginning of freezing)
  - Thickness of snow and slush
  - Thickness of snow ice and black ice
- Melting of ice
  - Air temperature, snow

Snow
Snow ice
Black ice
Water

#### Ice thickness model

- Ice growth calculated using heat transfer
  - Stefan's equations
  - Snow ice and black ice separately
  - $T_a$  =Air temperature, L =latent heat of water,
  - $h_{i,}$ ,  $h_s$  and  $h_{si}$  = thickness of black ice, snow ice and slush
  - $\rho_{si}$  and  $\rho_i$  =densitiv of snow ice and ice
  - $k_{i,}, k_s$  and  $k_{si}$  =heat transfer capacity
  - Indices *i*, *s* and *si* refer to (black) ice, snow and snow ice.
- Melting calculated using degreeday factor





#### Ice thickness model: results at Tornionjoki river Pello 28.11.2011



Observed (black diamond) and simulated total ice (solid blue line), snow ice (solid cyan line) and snow (dashed blue line above zero level)

#### Ice thickness model: results at Tornionjoki river Pello 25.4.2012



Observed (black diamond) and simulated total ice (solid blue line), snow ice (solid cyan line) and snow (dashed blue line above zero level)

#### Ice observation points 2006-2010 63 points total, ice thickness calculated and calibrated



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Järven jäänpaksuus (cm) 15. 3.2010 Forecast day 15. 3. Lake ice thickness (cm) 15. 3.2010

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