



Treatment and Impact of Specification of Errors in Data Assimilation

Sue Ballard, NetFAM/COST workshop April 2006

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Met Office, Exeter

1. Data assimilation

- Combining information, sources of information and Bayes' Theorem
- Variational data assimilation
- Background errors in 3D-Var and 4D-Var

2. Observation processing – preparing obs for data assimilation

- Bias Correction and Quality Control
- Observation error
- Thinning

3. Convective scale data assimilation for nowcasting/hydrology

- Impact of latent heat and moisture nudging
and treatment of “errors”

4. Conclusions

The ultimate problem in meteorology



Initial value problem

Deterministic (?)

Observations -> analysis -> forecast

Now accept that we need to allow for uncertainty and
can gain information
ensembles, probability, distributions

Allow for error in observations
error in forecast model

Observations – direct and indirect

eg screen temperature, mast wind, aircraft temperature

v satellite radiance, radar doppler radial wind, reflectivity

No uniform network, incomplete information

However information is advected and evolves nonlinearly

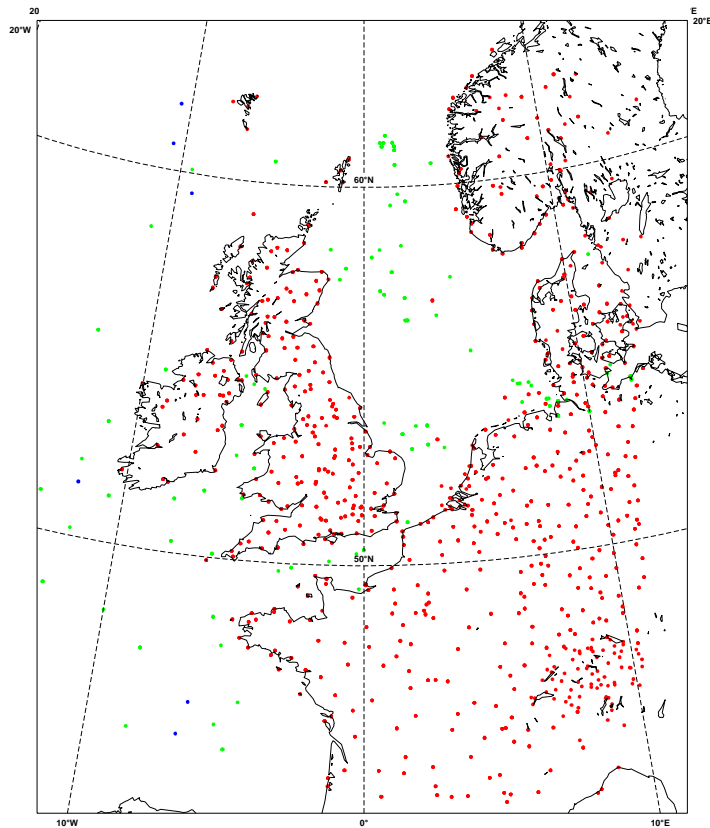
Previous forecast contains information from previous observations

Therefore combine latest forecast with latest observations

Sources of information

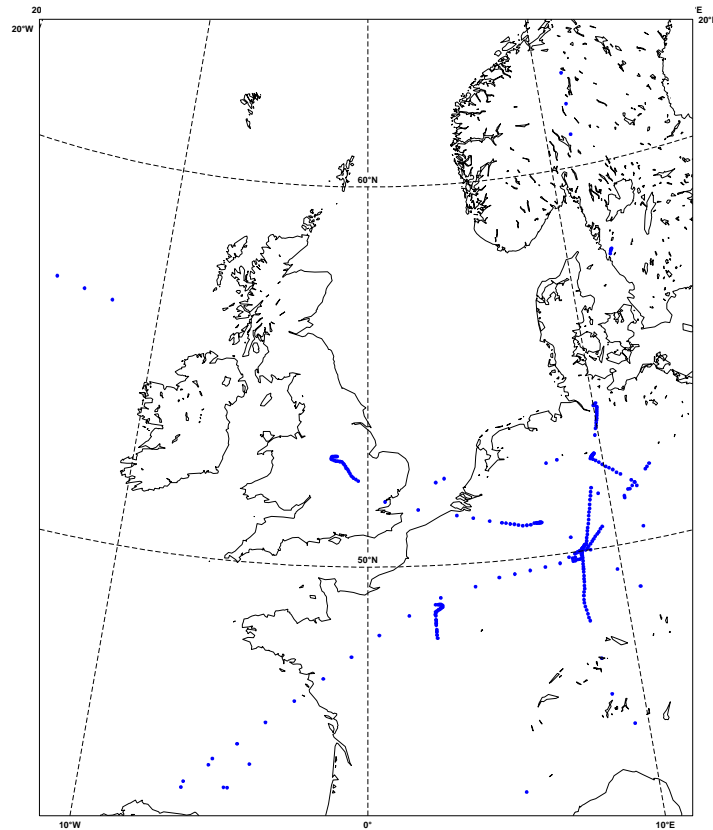
Data Coverage: Surface (21/3/2006, 0 UTC, qm00)
Total number of observations assimilated: 2041

LNDSYN (1840) SHPSYN (194) BUOY (7) MOPS PPN (0)

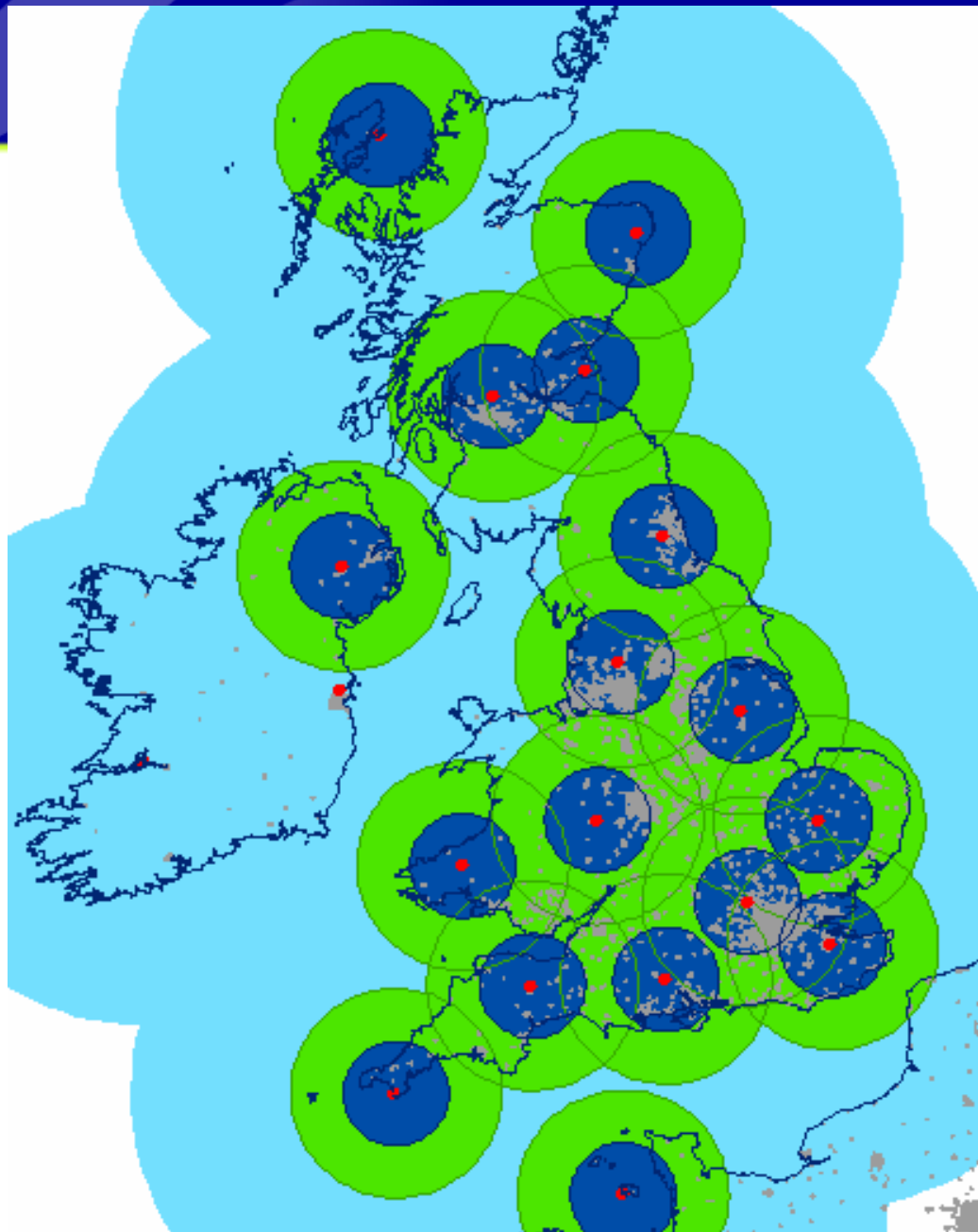


Data Coverage: Aircraft (21/3/2006, 0 UTC, qm00)
Total number of observations assimilated: 270

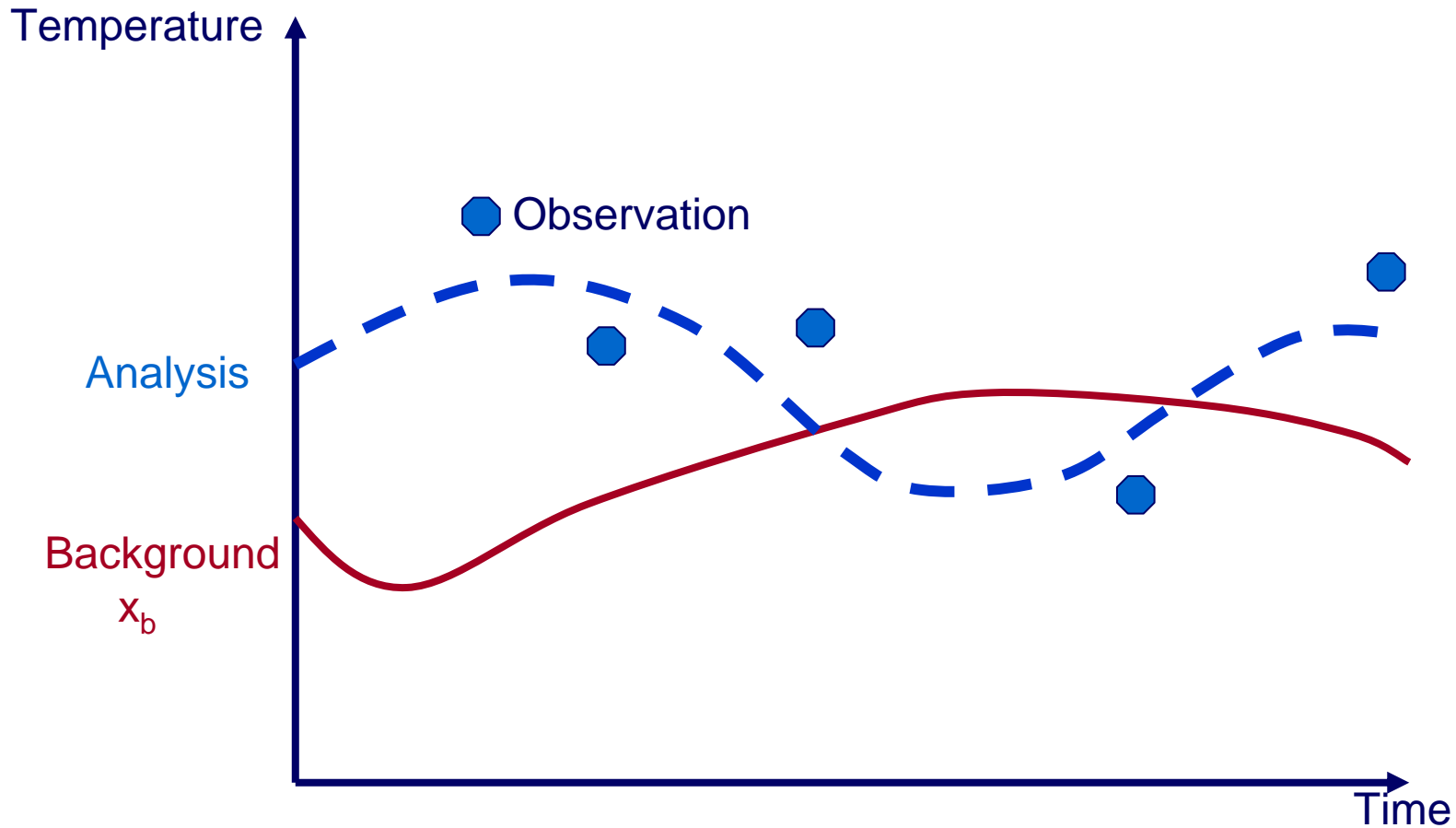
AMDARS (270) AIREPS (0) BOGUS (0)



Radar Network Coverage – 2006



Four-dimensional variational assimilation (4D-Var)



Gauss, 1809: *Theoria Motus Corporum Coelestium*

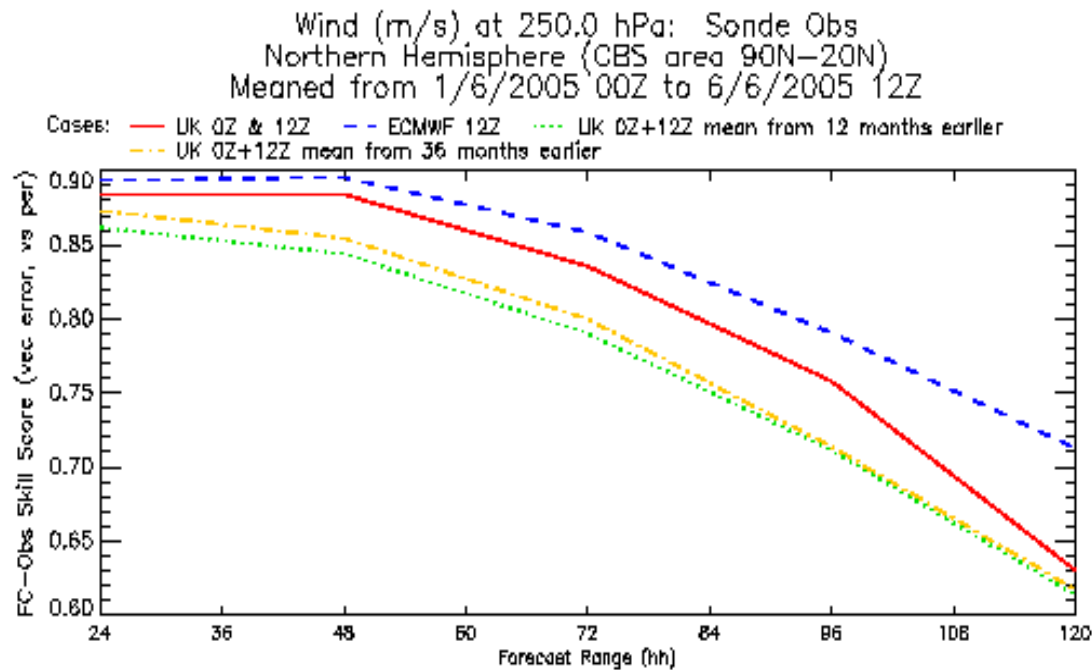
-1823: *Theoria combinationis observationum erroribus minimis obnoxiae*

- all models and observations are approximate
- the resulting analysis will also be approximate
- the observations must be combined in some optimal fashion
- it is better to have enough observations to over-determine the problem
- the model is used to provide a preliminary estimate
- the final estimate should fit the observations within their (presumed) observational error

Note: $O(10^6)$ obs

$O(10^7)$ model variables * grid-points

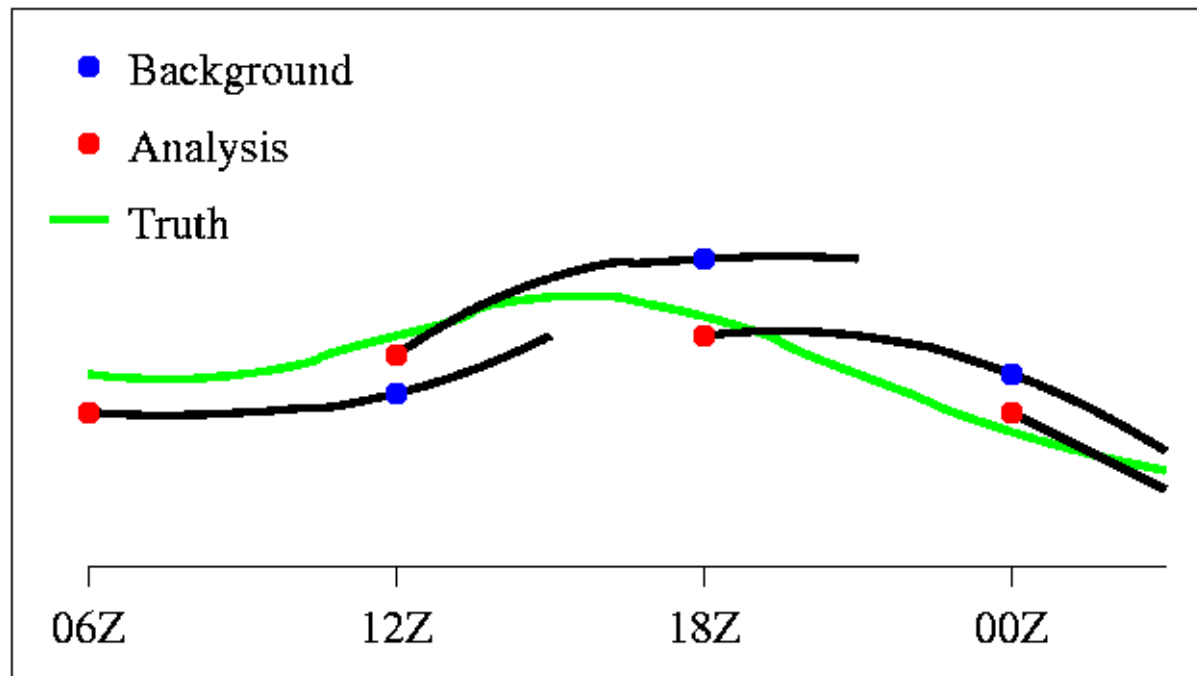
The background has information...



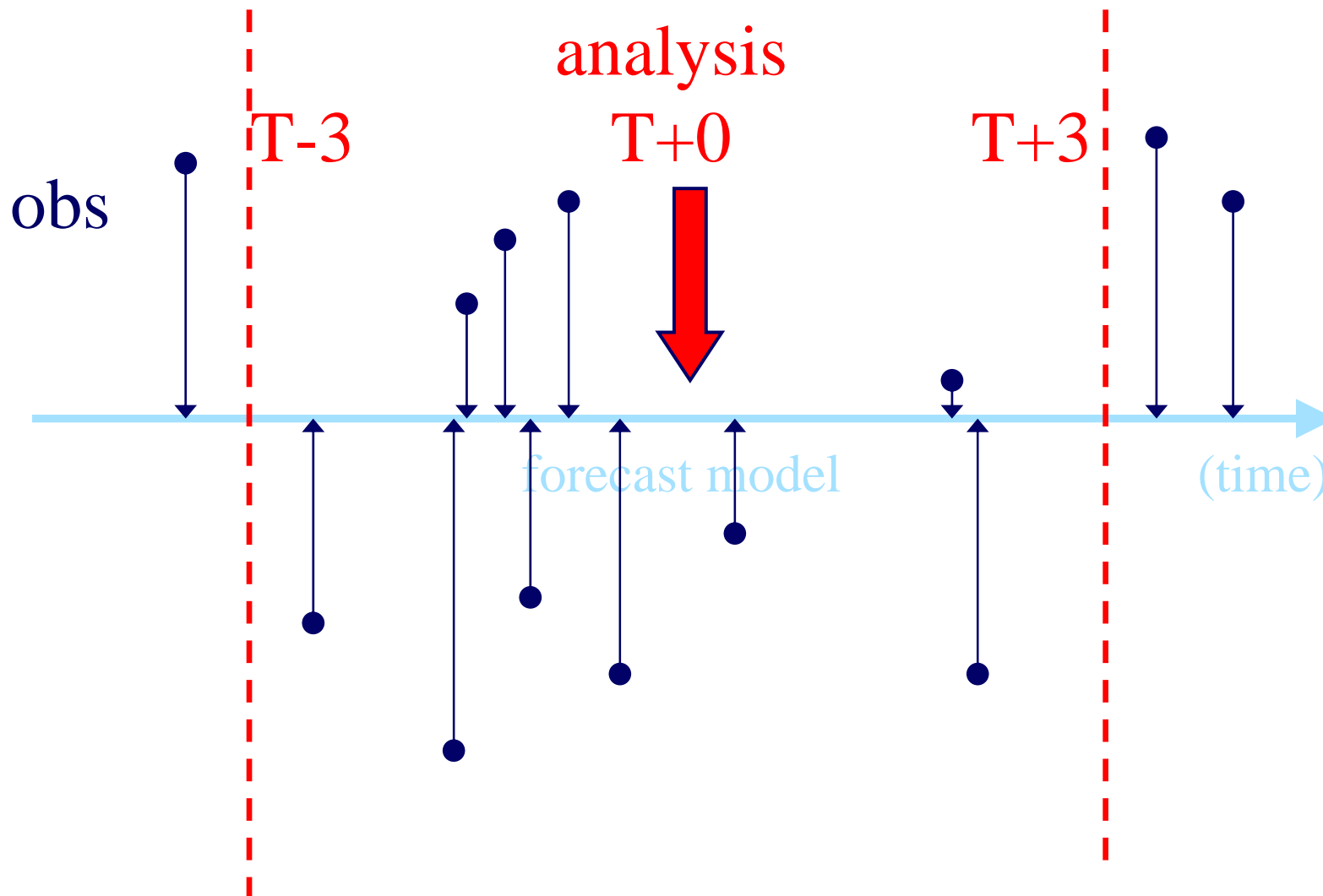
What is data assimilation?

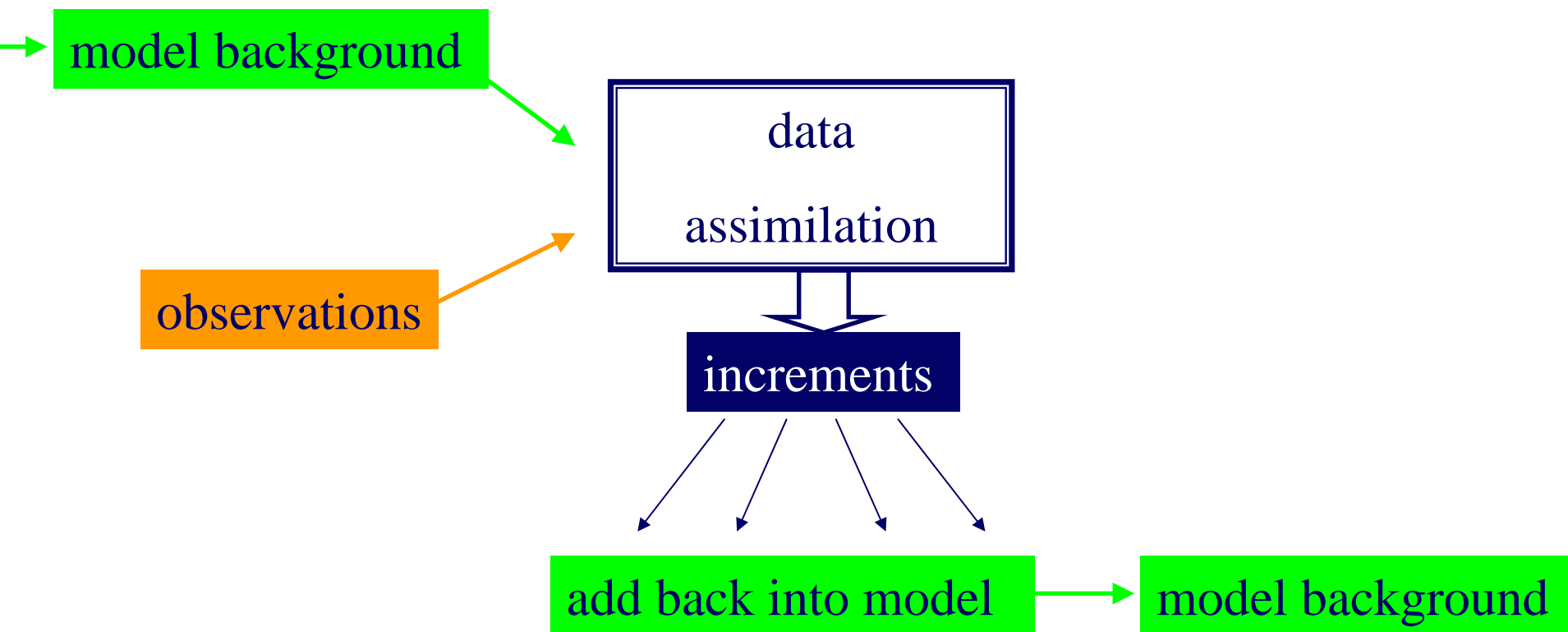


- The best and most powerful analysis systems are obtained by incorporating numerical models into analysis algorithms.
- The model encapsulates our understanding of the physical laws, and can be used to propagate observational information forwards in time.
- **Data assimilation:** the production of regular analyses for, and in conjunction with, a forecast model.
- Typically, an *intermittent* data assimilation cycle is used:



Intermittent Data Assimilation





Maximise $P(x|y, x^b)$ where $y = H(x) + e^o$ (observations) $x^b = x + e^b$ (background)

In case of Gaussian errors in the first guess and observations minimise

$$J(\mathbf{x}) = \frac{1}{2} (\mathbf{x} - \mathbf{x}^b)^T \mathbf{B}^{-1} (\mathbf{x} - \mathbf{x}^b) + \frac{1}{2} (H(\mathbf{x}) - \mathbf{y}^o)^T \mathbf{R}^{-1} (H(\mathbf{x}) - \mathbf{y}^o)$$

Where $\mathbf{B} = \langle e^b (e^b)^T \rangle$ background error covariances

$\mathbf{R} = \langle e^o (e^o)^T \rangle$ observation error covariances

- Assumptions: Gaussian errors, no bias, no correlations between background and observation errors
- Quality control needed to make errors more Gaussian
- The evaluation of $H(x)$ may involve time integration
- Data assimilation requires cycling in time (and ideally updating of \mathbf{B})
- The background can be thought of as a special set of observations

- To make VAR more practical, we can approximate the penalty J in terms of increments \mathbf{w} to a simplified model, and linearise about a guess \mathbf{x}^g state

$$J(\mathbf{x}) = \frac{1}{2} (\mathbf{x} - \mathbf{x}^b)^T \mathbf{B}^{-1} (\mathbf{x} - \mathbf{x}^b) + \frac{1}{2} H(\mathbf{x}) \mathbf{y}^o{}^T \mathbf{R}^{-1} H(\mathbf{x}) \mathbf{y}^o$$



$$J(\mathbf{w}) = \frac{1}{2} \mathbf{w}^T \mathbf{B}_{(\mathbf{w})}^{-1} \mathbf{w} + \frac{1}{2} \mathbf{G} \mathbf{w}^T \mathbf{R}^{-1} \mathbf{G} \mathbf{w} + \mathbf{d}^T \mathbf{R}^{-1} \mathbf{d}$$

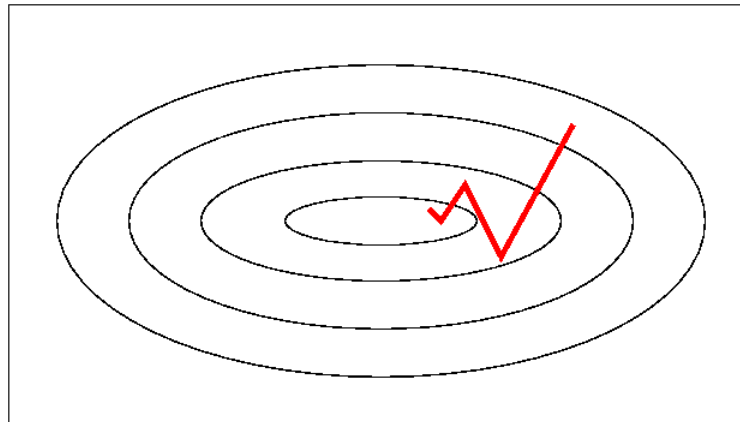
$$\mathbf{w}^g = \mathbf{S} (\mathbf{x}^g - \mathbf{x}^b)$$

$$\mathbf{d} = \mathbf{y}^o - H(\mathbf{x}^g)$$

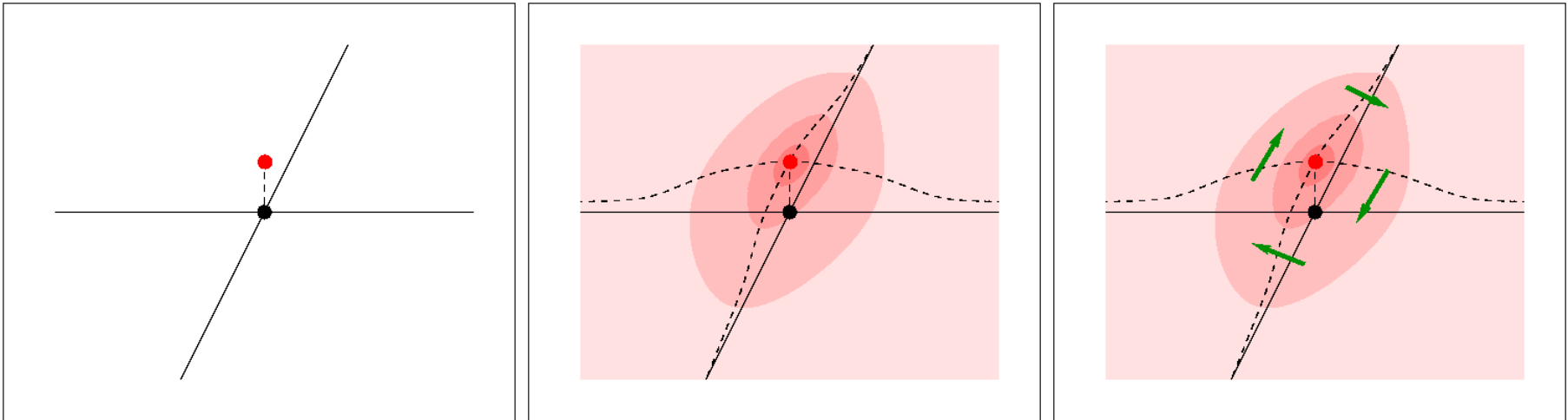
$$\mathbf{G} = \mathbf{H} \mathbf{S}^T$$

$$\mathbf{x} = \mathbf{x}^g + \mathbf{S} \mathbf{w}$$

- VAR is an inverse problem: we don't 'interpolate' observations onto the model grid, we **vary** the model state until we find that which is most compatible with the data, as defined by J .
- We can assimilate variables not directly related to the model variables, as long as we can write a reasonably accurate observation operator.
- For example, we can use satellite radiance measurements directly.
- To minimise J , we use an iterative **descent algorithm**. On each iteration, the algorithm needs J and its gradient wrt \mathbf{x} .



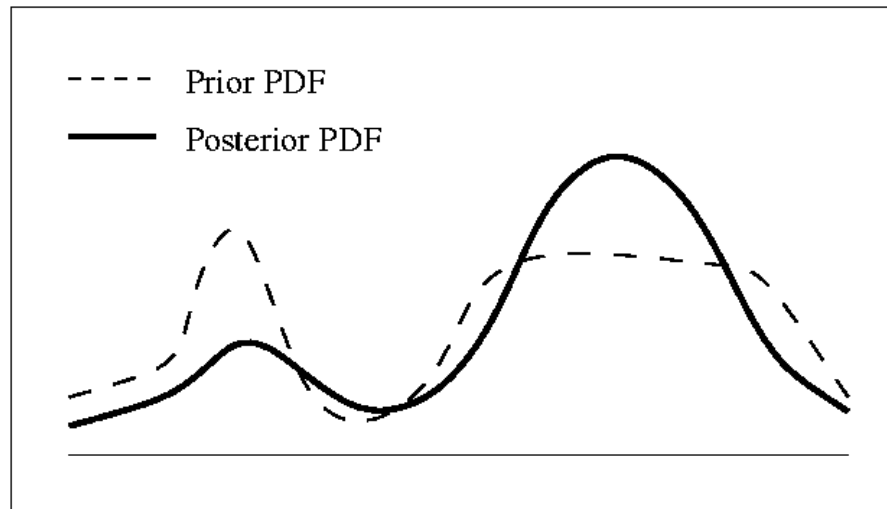
- The background error covariance matrix \mathbf{B} describes the error variance for each model variable, and the **correlations** between errors in different model variables; ie. how information from observations should be spread:



- Incorporating better approximations of the `true' background error covariance matrix is perhaps THE most important theoretical challenge in data assimilation.

- **B** is $N \times N$, where N is the number of (simplified) model variables.
- In order to make **B** manageable, we need to make some assumptions.
- Transform to variables that are approximately uncorrelated:
 - Transform to **streamfunction, velocity potential, unbalanced pressure** and **rh**.
 - Project onto vertical modes.
 - Project onto horizontal spectral modes.
- **B** is then diagonal, and easily dealt with.

- If the statistics of the background and observation errors were known, we could in theory use Bayes' Theorem to deduce the PDF $p(\mathbf{x} | \mathbf{x}_t)$ and choose an appropriate analysis:



- Statistics are not known, and in general could not be represented anyway.
- Need to make some assumptions.

Generation of Background Errors - Not known so have to be estimated eg

NMC method

Statistics are gathered by comparing pairs of forecasts valid at the same time (eg. T+48 - T+24)

The statistics are **climatological**, and approximately **homogeneous** and **isotropic**

can extend to vary latitudinally or use defined horizontal correlations

Idealised horizontal and vertical correlations

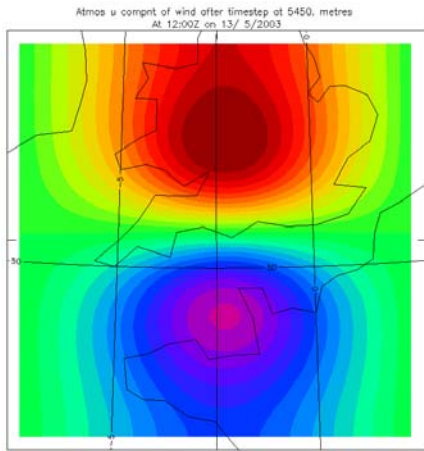
Ensembles of forecasts valid at forecast range of background

Analyses and Forecasts are very sensitive to specification of background errors variances, correlations and lengthscales

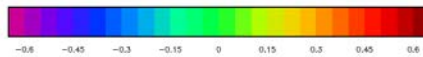
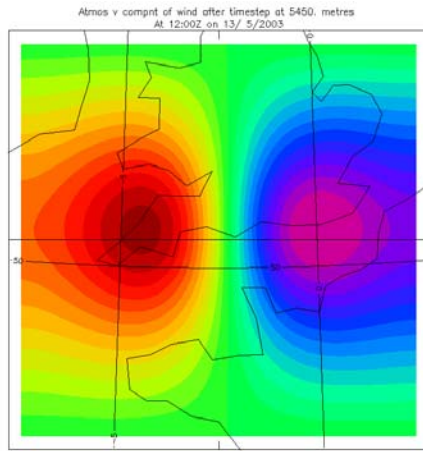
Spreading information – background errors



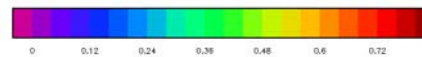
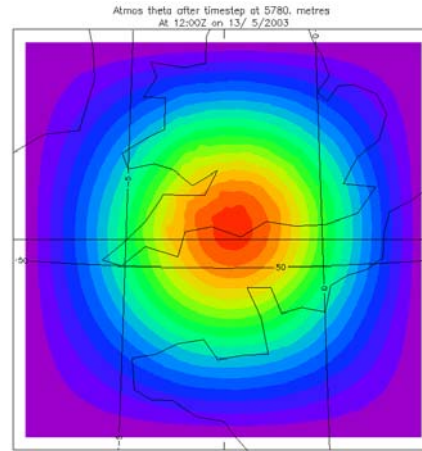
u



v



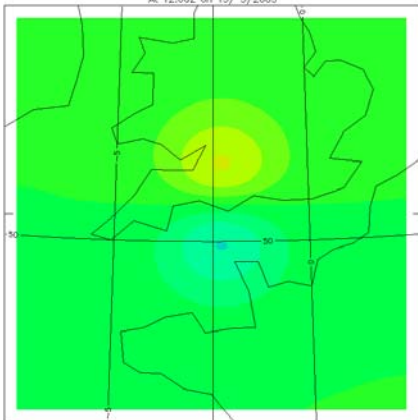
theta



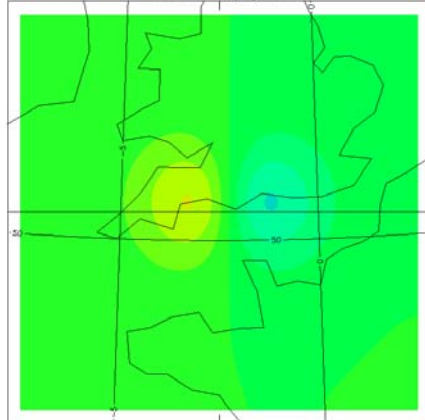
Derived for
12km model
Used in 4km

5.78 km
Approx
500hPa

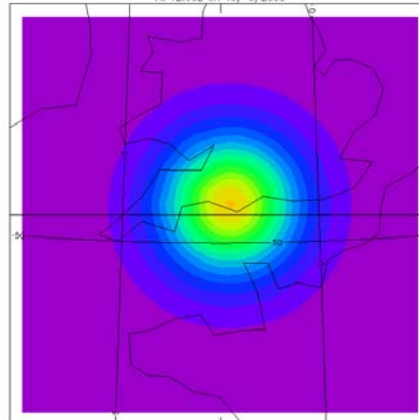
Atmos u compnt of wind after timestep at 5450. metres
At 12:00Z on 13/ 5/2003



Atmos v compnt of wind after timestep at 5450. metres
At 12:00Z on 13/ 5/2003



Atmos theta after timestep at 5780. metres
At 12:00Z on 13/ 5/2003



1/2 lengthscale

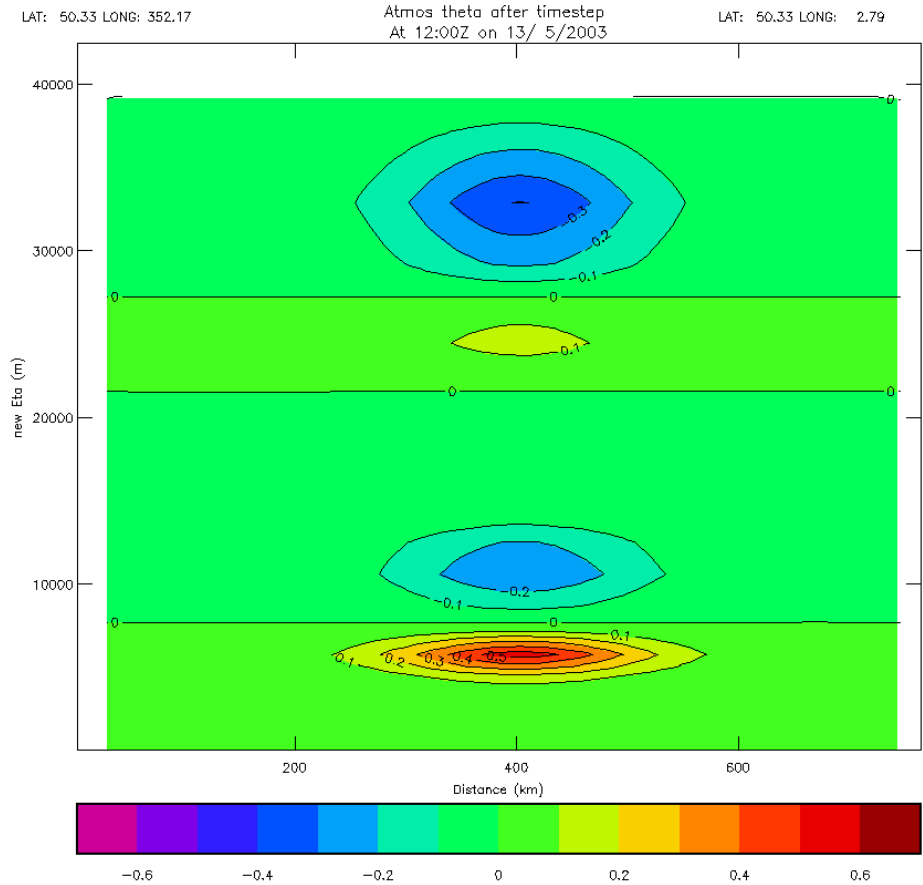
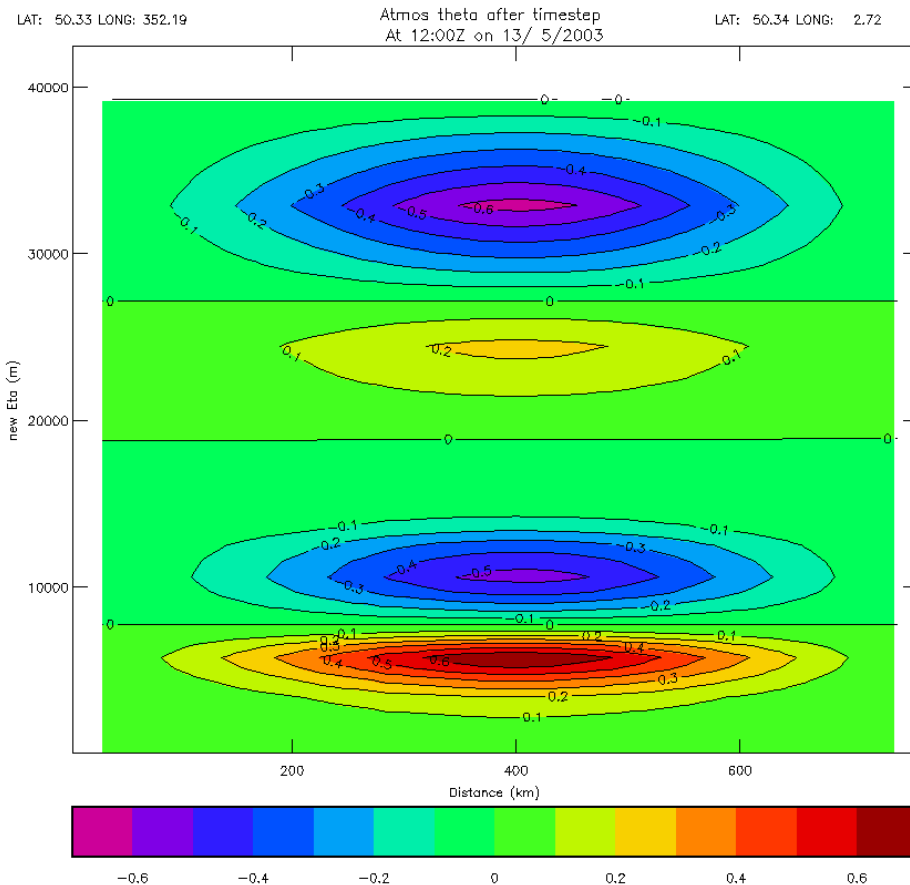
Spreading information – background errors theta increment at 5.78km



Derived for 12km model

theta

$\frac{1}{2}$ lengthscale



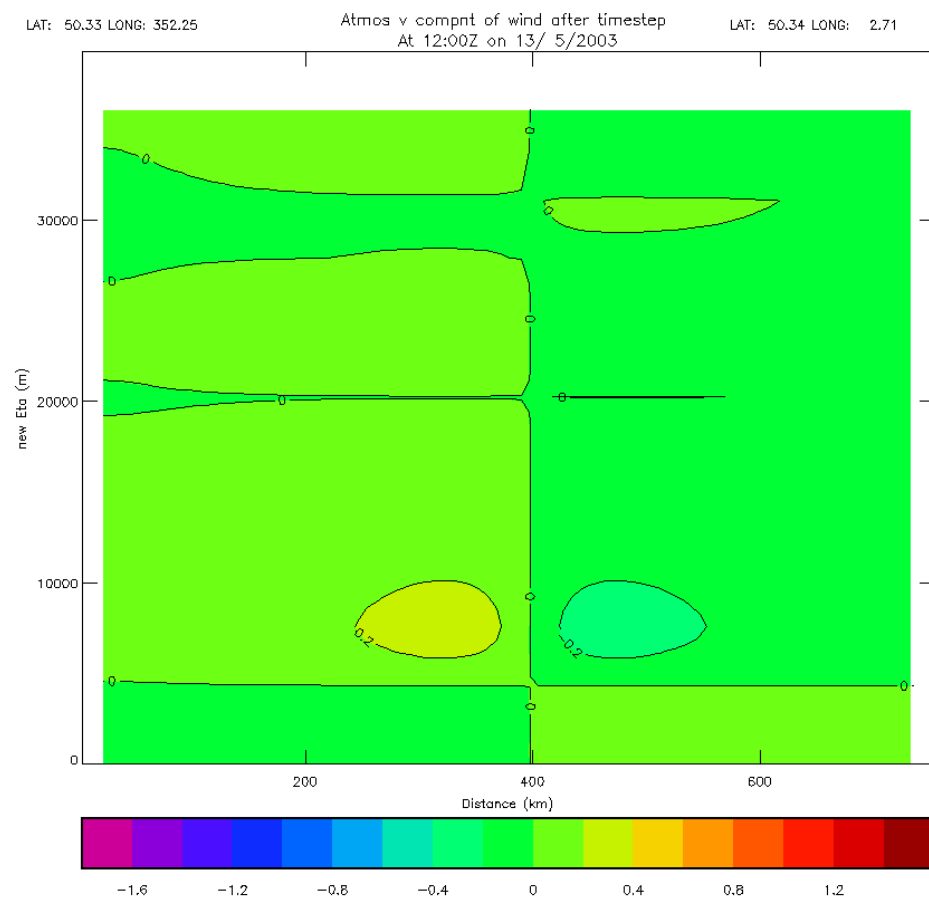
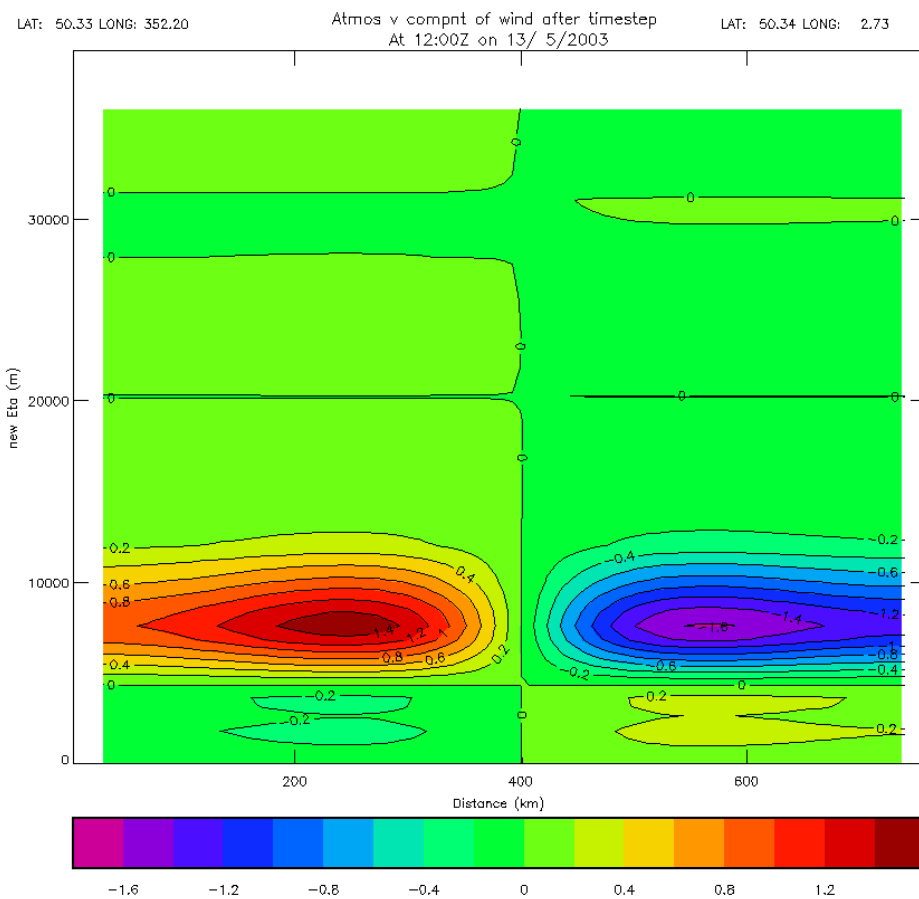
Spreading information – background errors v increment at 5.78km



Derived for 12km model

v

1/2 lengthscale

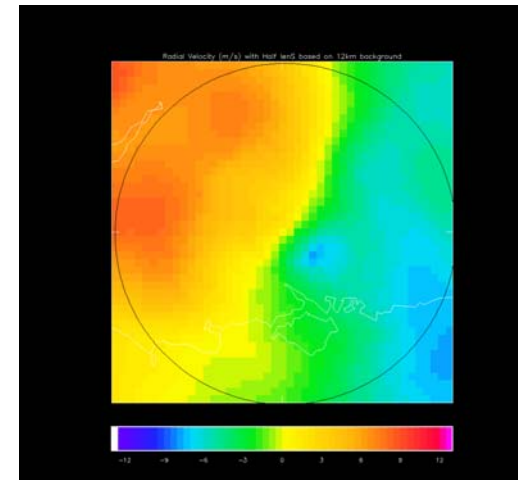
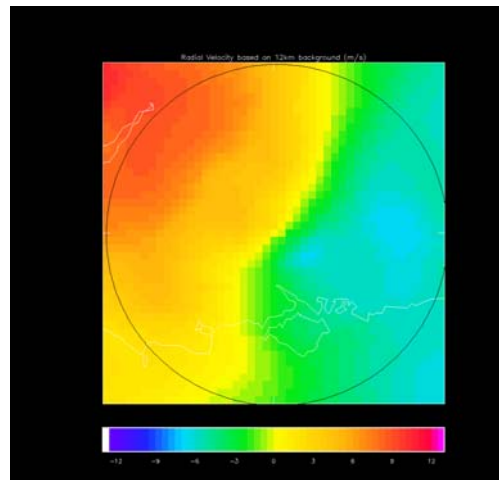
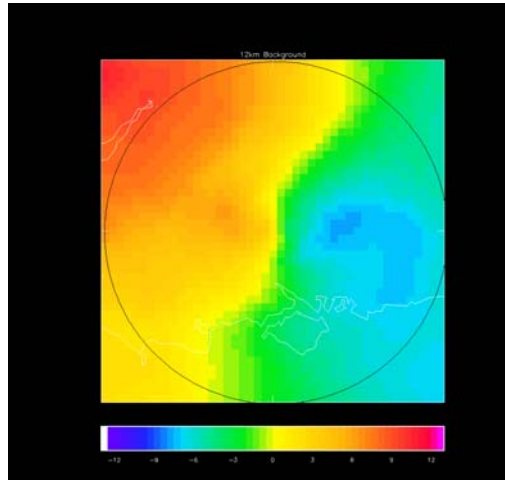


Impact of S-band radial radar wind data - radial wind on 1deg scan elevation

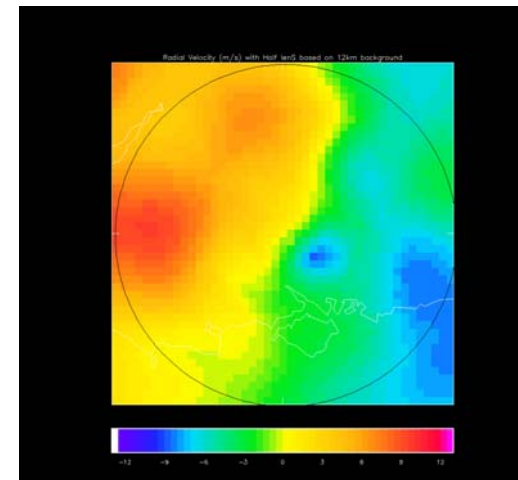
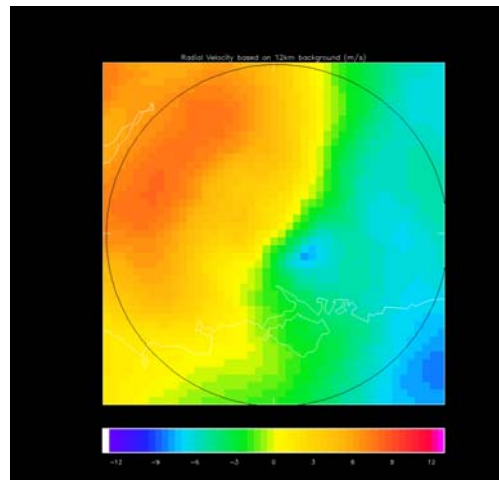
Analysis

1/2 Length scale

12km
Back-
ground



Super-
obbed
Radar
Doppler
wind

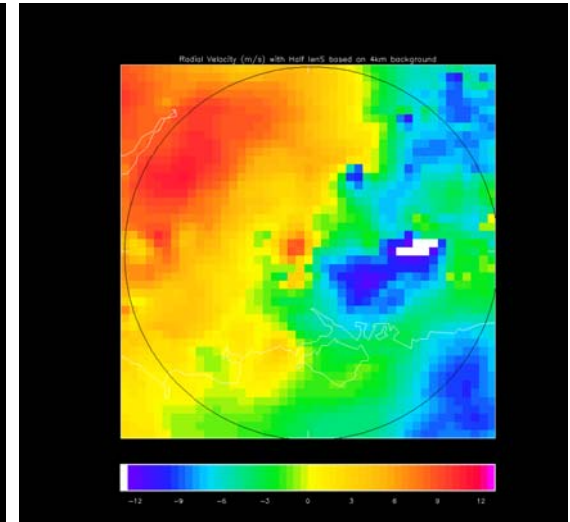
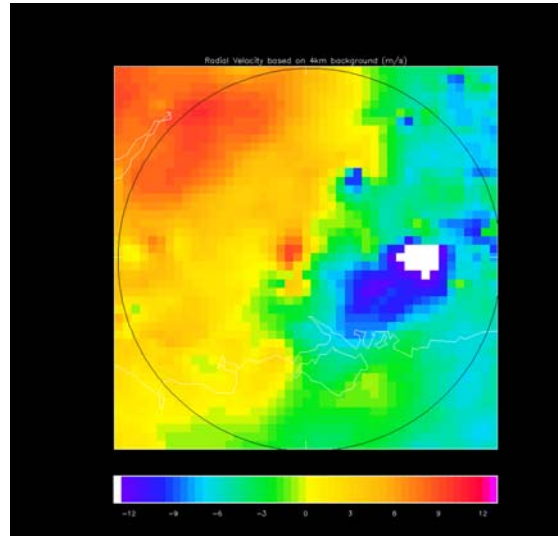
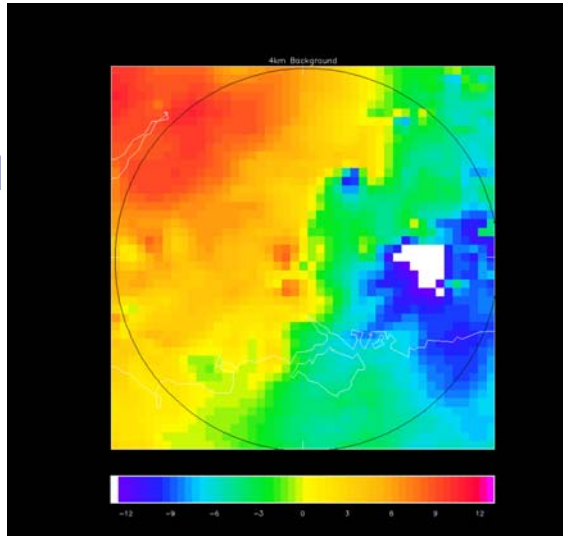


Impact of S-band radial radar wind data - radial wind on scan elevation

Analysis

1/2 Length scale

4km
Back-
ground

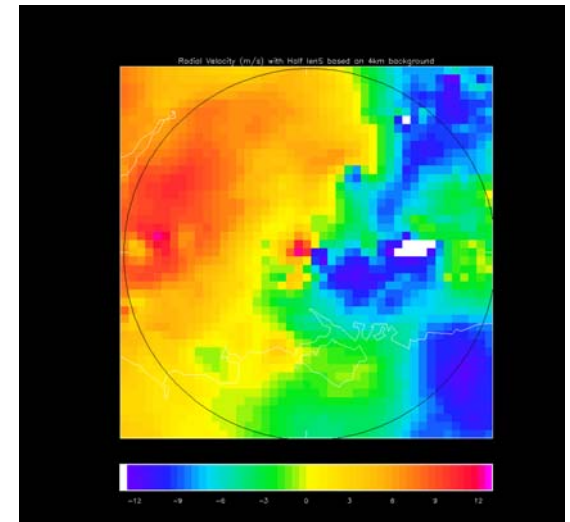
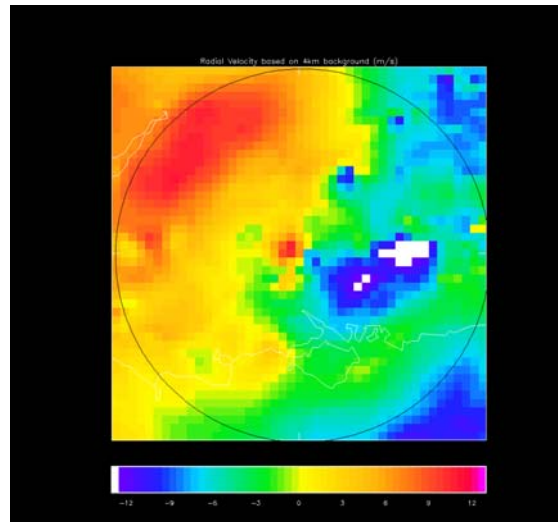


Super-
obbed
Radar
Doppler
wind



Raw Radial Velocity Obs

Chilbolton PPI scan 12z 010703, 1 deg elev, max range 90 km

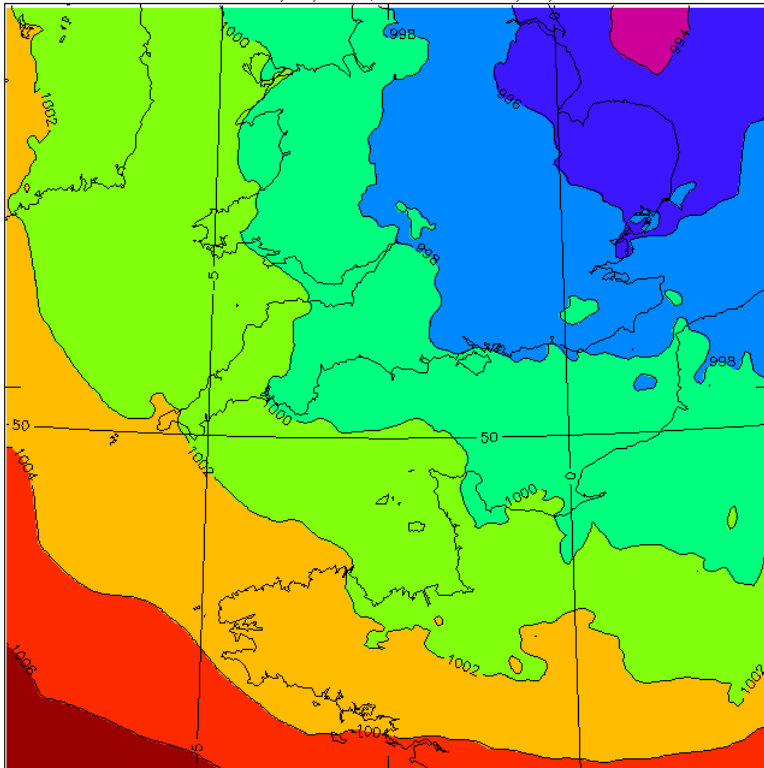


Reduced background wt

Impact of S-band radial radar wind data - T+2 surface pressure

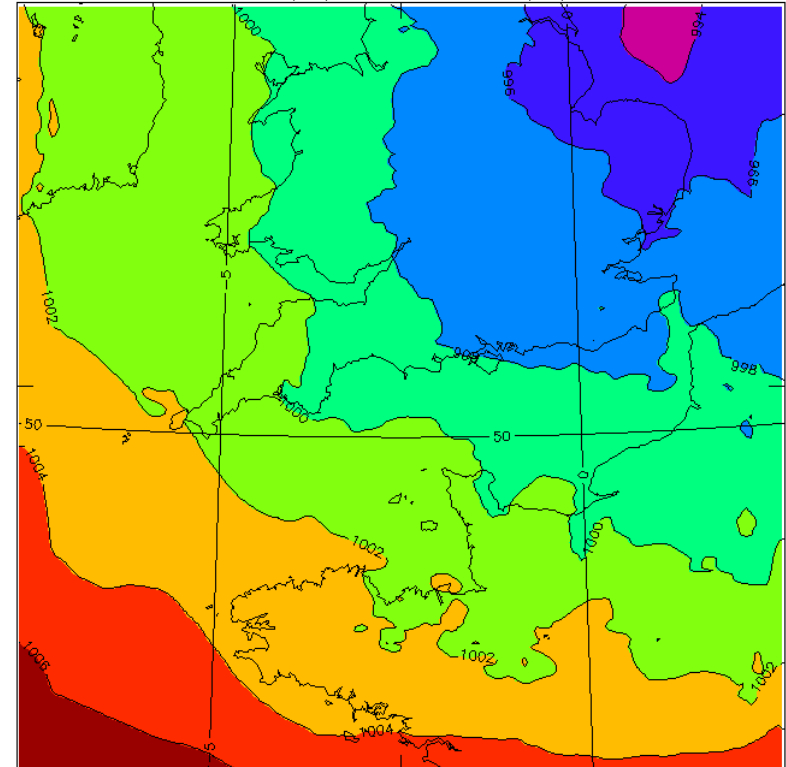
Background forecast

XAKVE Mean sea Level Atmos pressure at mean sea level
At 14Z on 1/ 7/2003, from 12Z on 1/ 7/2003

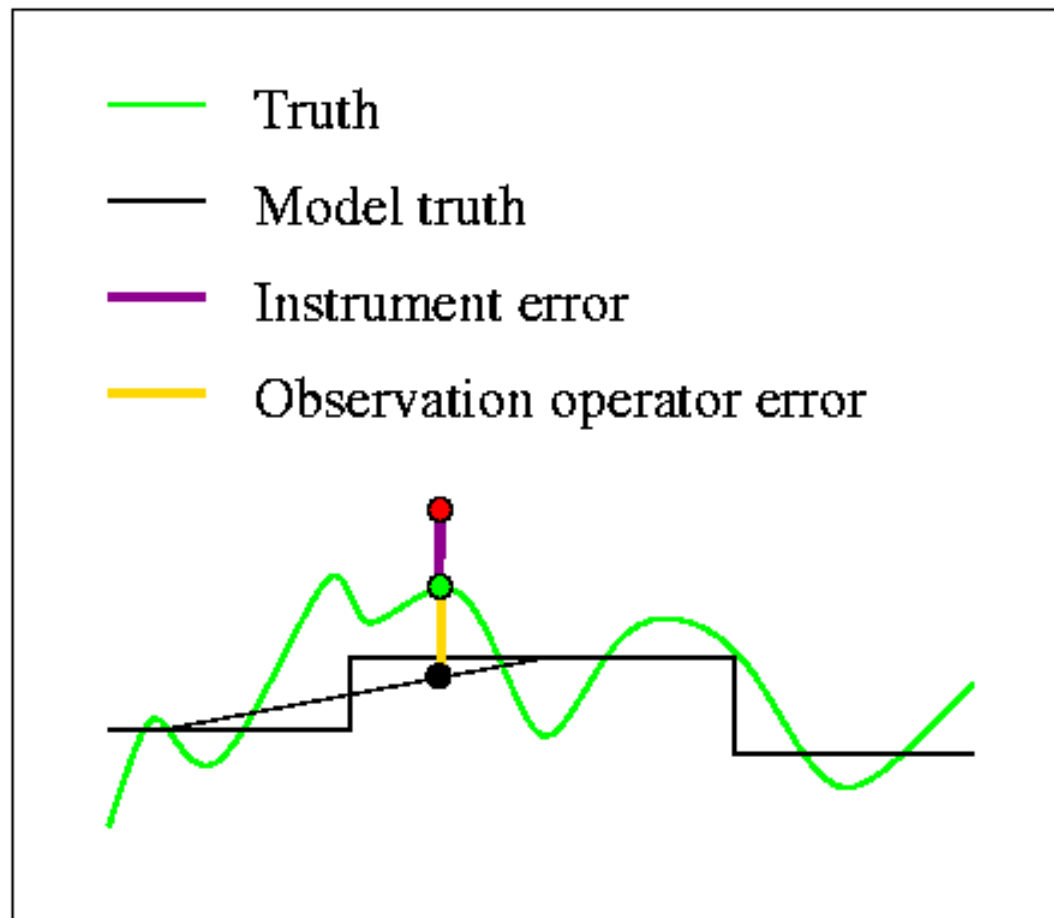


Including radar radial doppler winds

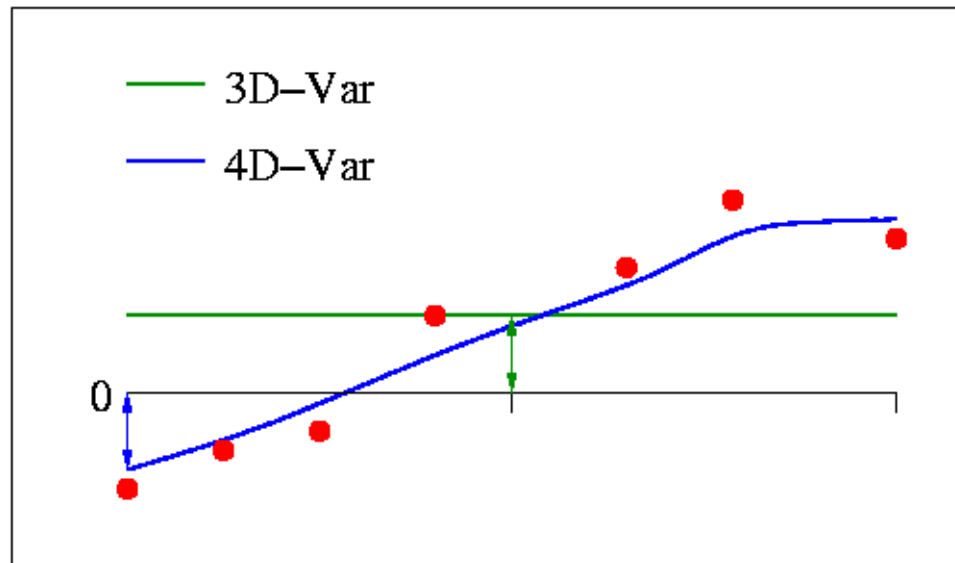
XAKVE Mean sea Level Atmos pressure at mean sea level
At 14Z on 1/ 7/2003, from 12Z on 1/ 7/2003



- **R** describes the observation error, which combines **instrument error** and errors in the observation operator **H** (**representativeness error**):

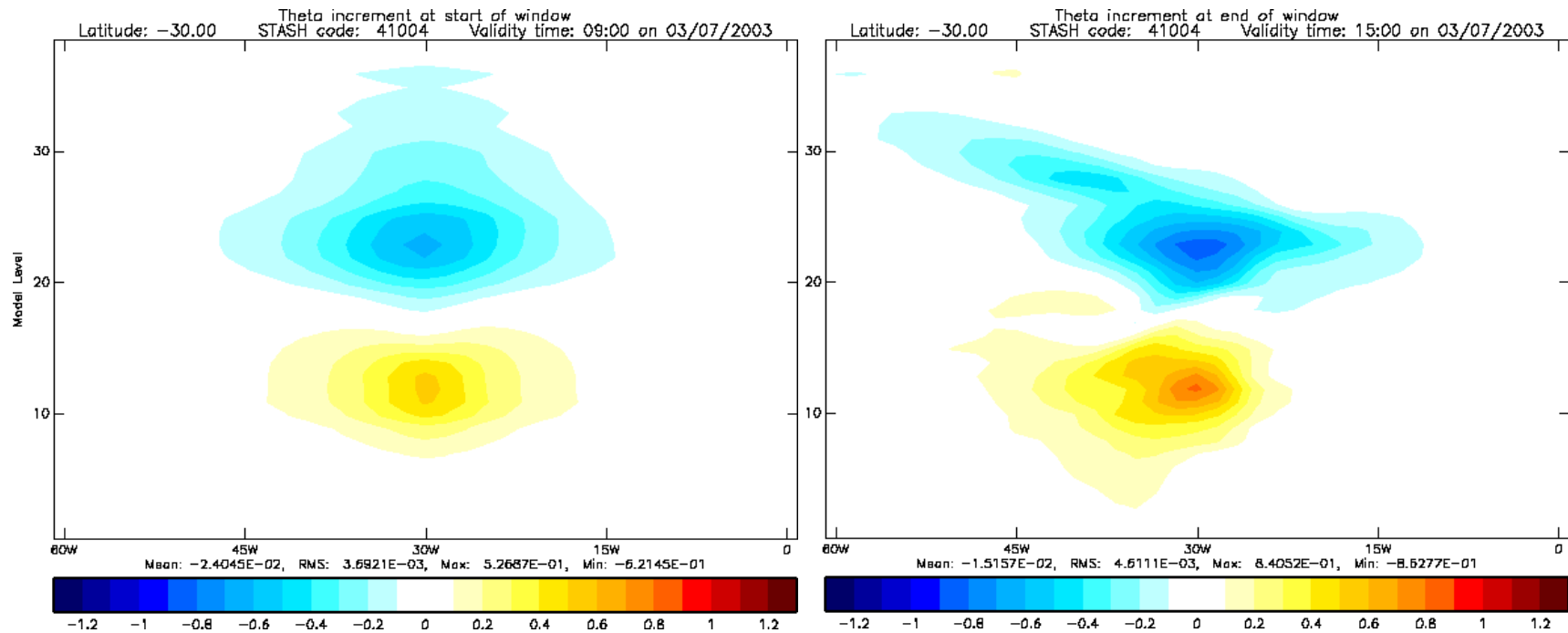


- In 4D-Var, the observation operator \mathbf{H} includes model forecasts to the observation times:



- Initial “background” forecast with nonlinear model
- 3D-Var either one value at centre of time window or forecast at time of observation and analyse increments assuming valid at analysis time FGAT
- Forecasts are updated with a simplified linear model; one forecast plus one **adjoint** integration per iteration. (Usually run about 50 iterations.)

- In 4D-Var, the background error covariance matrix \mathbf{B} is implicitly evolved by the linear model:



Pseudo pressure ob at beginning of window

Pseudo pressure ob at end of window

- This imposes some degree of dynamical consistency on the increments, and is probably a key advantage of 4D-Var.

- The occurrence of observations with gross errors can be handled
 - either in a prior Bayesian quality control, rejecting those not likely to come from the assumed Gaussian distribution characterising "good" observations
 - or by altering the variational penalty function to have a similar effect.

- In either method it is important to have a reliable estimate of the background error variance for that case, otherwise we can end up rejecting just those observations which show up a significant error in the background state.

- Theory assumes unbiased data

Conventional observations - Checks for:

- Physically plausible
- Position (e.g. ships over land)
- Track (movement since last report)
- Buddy checking (against neighbours)
- Model background O-B comparison
- Rejection lists from regular monitoring (O-B, O-A)

Both background and observation can be biased

- Correct for O-B difference bias
- Recalculate each month
- Vary strongly with scan position and latitude
- Scan – difference from nadir
- Airmass – linear regression
- If not possible reject data

$$y_{cor} = y_{raw} - C_{scan} - C_1 y_1 - \dots - C_n y_n - C_{air}$$

- Bad raw radiance T
- O-B threshold test
- Bad scan position
- Masked out
- Error in RTTOV
- Failed to converge
- High altitude
- Bad retrieved brightness temperature

Is radar data biased? What does bias depend on?

Need to bias correct the data and quality control it either externally or using variational quality control eg as at ECMWF

How do we specify error variances for radar data?

Precipitation rates

Reflectivity

Doppler radial winds

Refractivity

Do we have an absolute measure?

However in fact the errors need to be relative to errors for background and other observation types (tuning)

Errors correlated as in satellite scans

Observations: Radar Radial Wind Errors



Sources of error

Ground clutter, anomalous propagation, sea clutter, velocity folding, noise

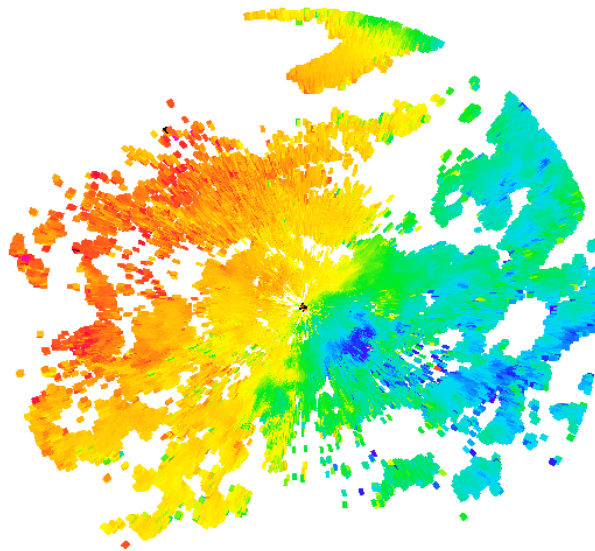
Velocity gradient in pulse volume

?

Superobbing of radar doppler winds

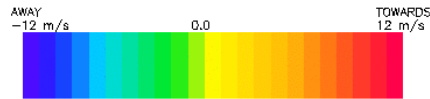


Raw data

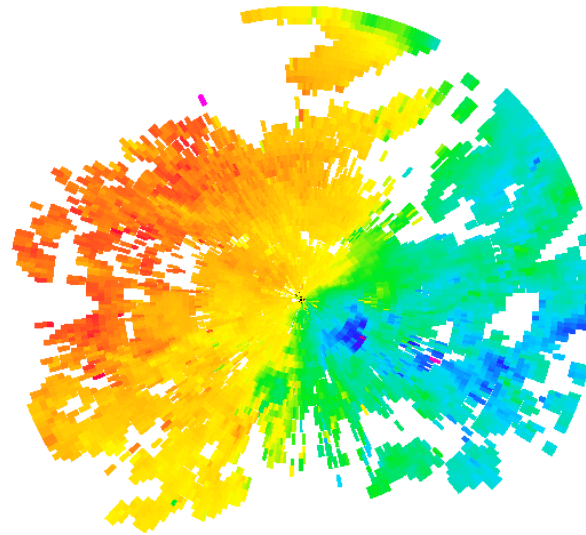


Raw Radial Velocity Obs

Chilbolton PPI scan 12z 010703, 1 deg elev, max range 90 km



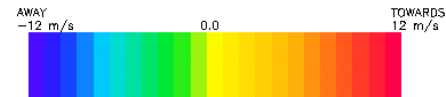
Superobbed data
Difference from background
method



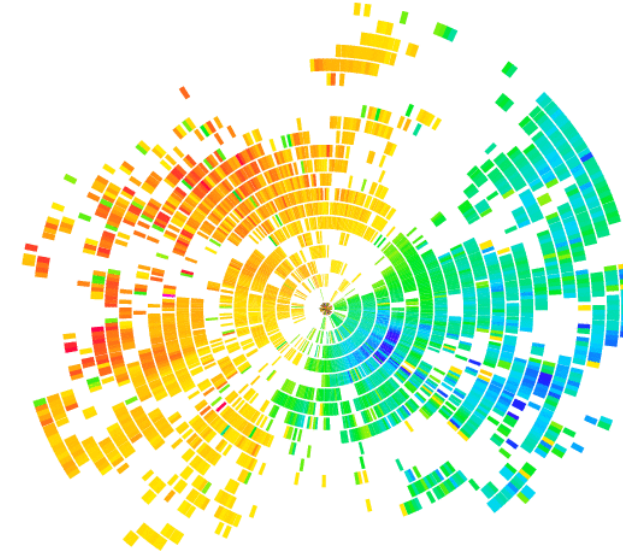
Radial Velocity Superobs

Cell Size 1 degree x 2 km

Chilbolton PPI scan 12z 010703, 1 deg elev, max range 90 km

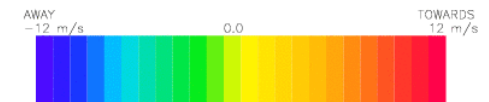


Superobbed data
Remapping method



Radial Velocity Superobs (F Rihan)

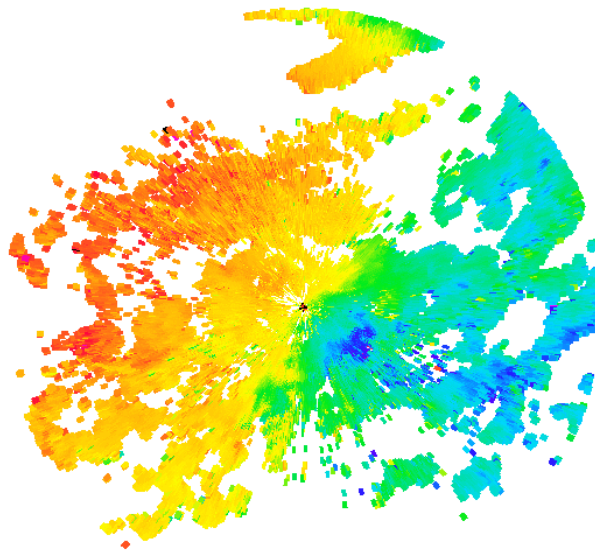
Chilbolton PPI scan 1215z 010703, 1.0 deg elev, max range 90 km



Errors of Superrobbed radar doppler winds

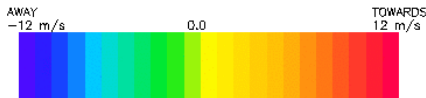


Raw data

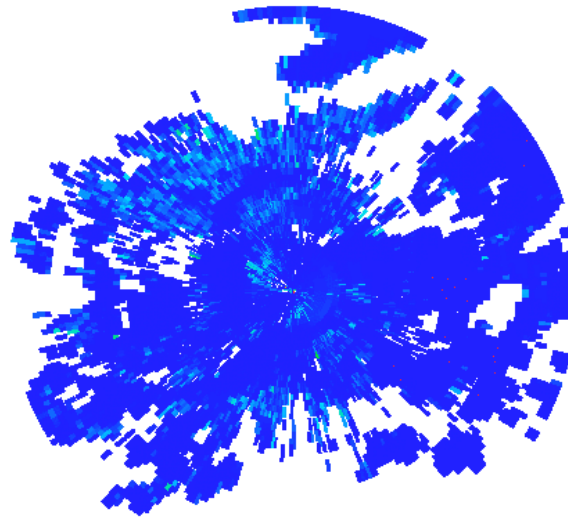


Raw Radial Velocity Obs

Chilbolton PPI scan 12z 010703, 1 deg elev, max range 90 km



Superrobbed data
Difference from background
Method – observation errors



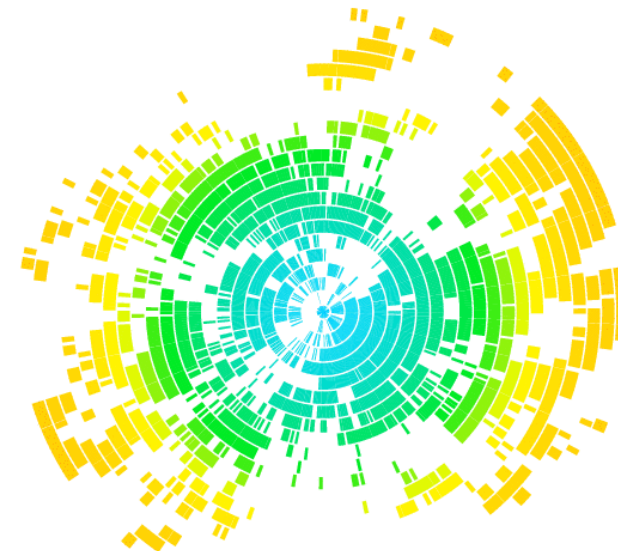
Radial Velocity Superob Error

Cell Size 1 degree x 2 km

Chilbolton PPI scan 12z 010703, 1 deg elev, max range 90 km



Superrobbed data
Remapping method
Total observation error



Radial Velocity Superob Error (F Rihan)

Chilbolton PPI scan 1215z 010703, 1.0 deg elev, max range 90 km



- Need sophisticated quality control
- Need specification of error as variance/standard deviation in units of observation
- Ideally need to allow for correlation of errors
- Need to thin data or perform super-obbing

High resolution data assimilation

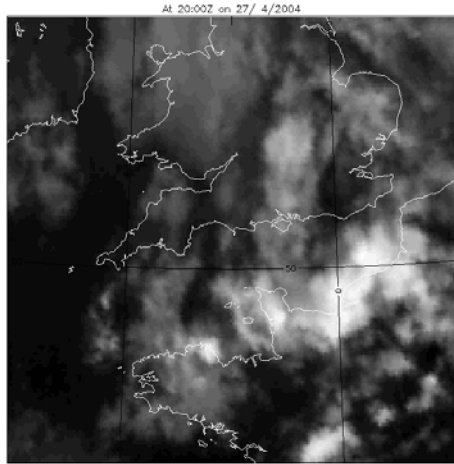


4km 3D-VAR with continuous cycles

With or without moisture and Latent Heat Nudging (LHN) using AC scheme
(referred to as MOPS data – moisture observation processing system)

- IAU – increments output from 3D-Var and fixed over time window
- AC scheme – increments depend on latest model fields so vary with timestep through weighting factor and model evolution/impact of data

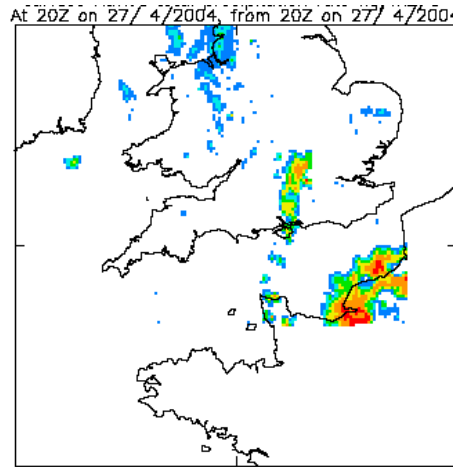
Moisture Observation Preprocessing



Satellite data



Surface reports



Radar data

Precipitation
5km
smoothed to
15km
Hourly
Testing 15min



**3D Cloud
fraction**

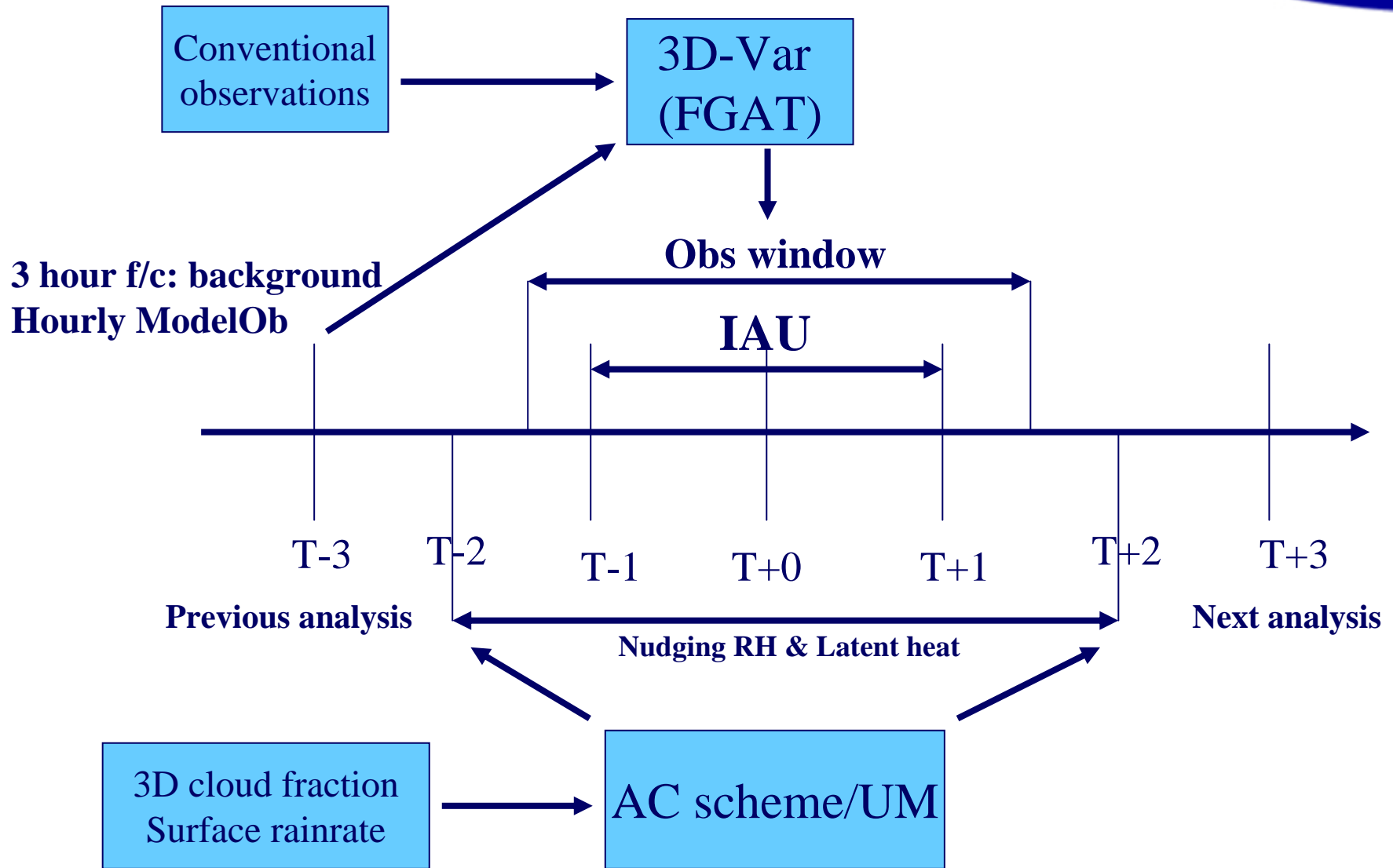


**3D Relative
humidity**



**Nudge
model state**

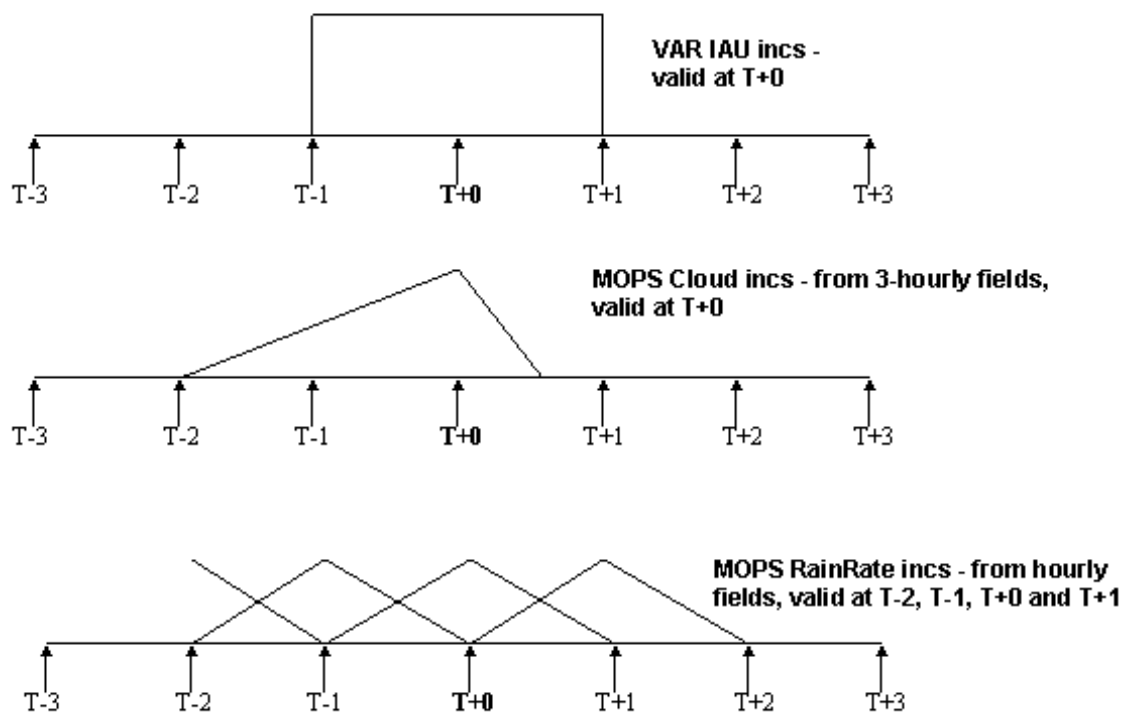
- Resolution: 15km, 3 hours
(Testing 1 hour)



Period over which observations and analysis increments are nudged into Unified Model



Initialisation of the Mesoscale Model: Weights given to Var & MOPs data



Impact of cloud and precipitation data

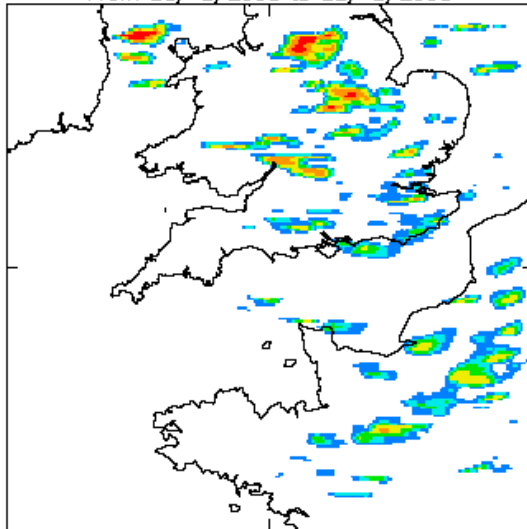
14UTC 25 August 2005 – CSIP IOP 18

T+2 forecast
No MOPS data

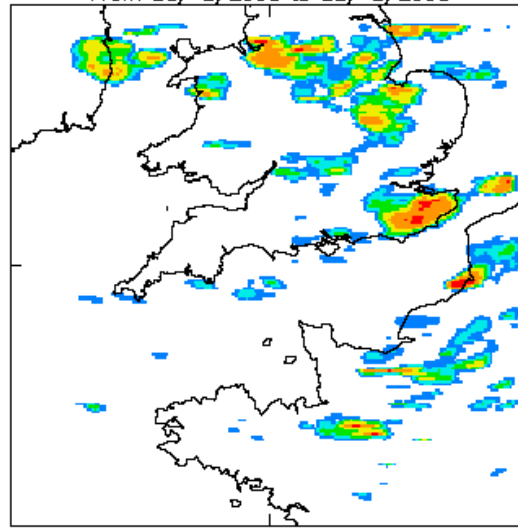
T+2 forecast
15min precip and hrly cloud

Radar
1 hour accumulation

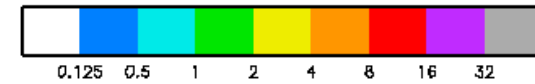
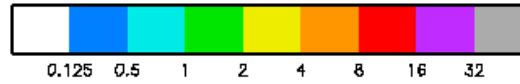
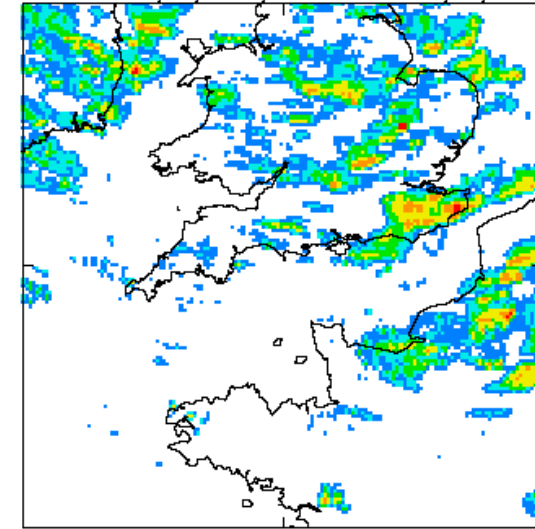
14Z MDC05_20050825QM12_000 4km
XAQQD Time mean
surface Atmos total precipitation amount kg/m2/ts
From 25/ 8/2005 to 25/ 8/2005



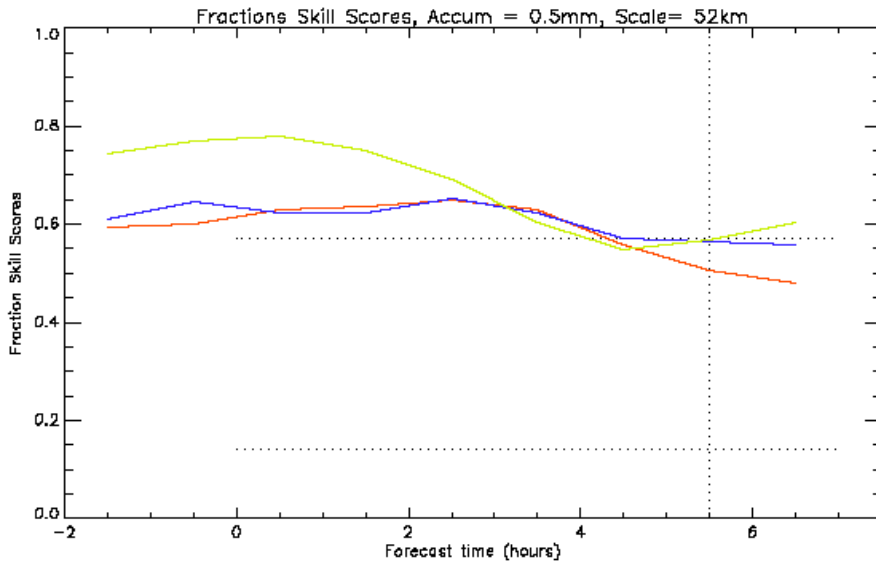
14Z MDC19_20050825QM12_000 4km
XAQQG Time mean
surface Atmos total precipitation amount kg/m2/ts
From 25/ 8/2005 to 25/ 8/2005



14Z RADAR HOURLY ACCUMULATION
AAAAJ Time mean
surface Atmos total precipitation amount kg/m2/ts
At 14Z on 25/ 8/2005, from 14Z on 25/ 8/2005



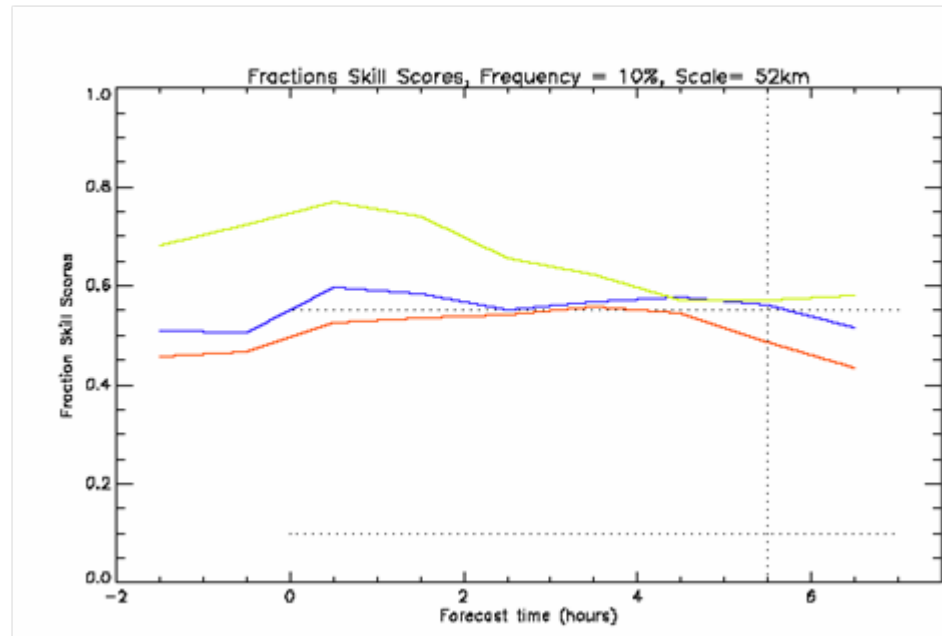
Impact of cloud and precipitation data



Skillscore for top 10% accumulations

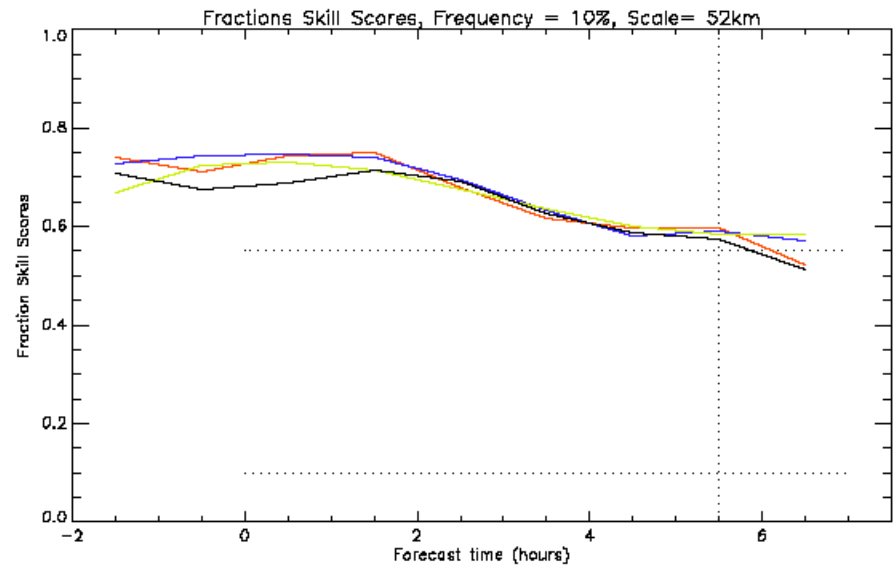
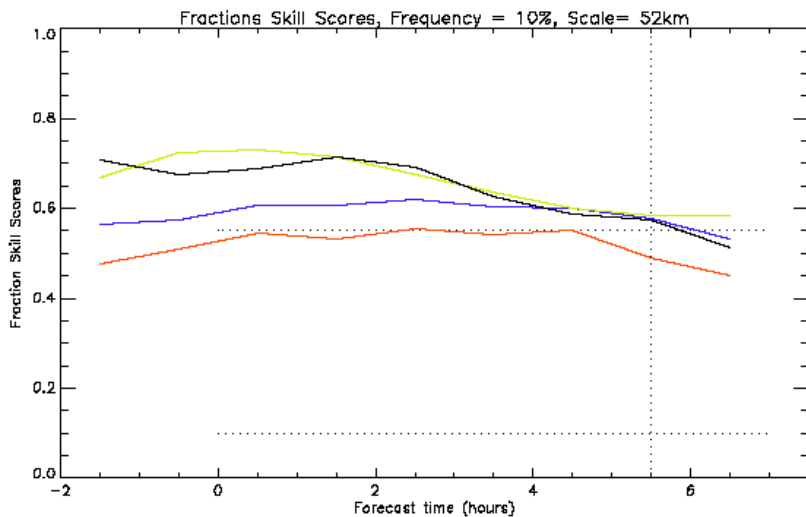
Skillscore for accumulation > 0.5mm

- 15min precip, hrly cloud reduced filtering
- hourly precip, 3 hr cloud full filtering
- no precip and cloud



Impact of cloud and precipitation data

5 cases



Skillscore for top 10% accumulations

- _____ cloud = 3hrs; rain = 1hr; no filter
- _____ cloud = 1hr; rain = 15min; no filter
- _____ cloud=3hrs; rain=1hr
- _____ full filtering fwhm=42km
- _____ no precip and cloud

Skillscore for top 10% accumulations

- _____ cloud = 3hrs; rain = 1hr; no filter
- _____ cloud = 1hr; rain = 15min; no filter
- _____ cloud=3hrs; rain=1hr; no filter
- _____ diagnostic rain
- _____ cloud=1hrs; rain=15min; no filter
- _____ no subcloud LHN

Impact of cloud and precipitation data



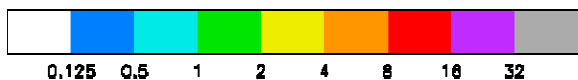
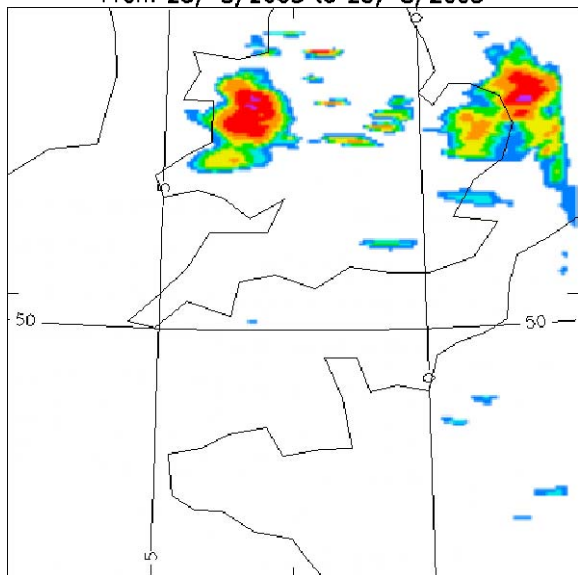
17UTC 25 August 2005 – CSIP IOP 18

T+2 forecast
42km filter
cloud=3hr rain=1hr

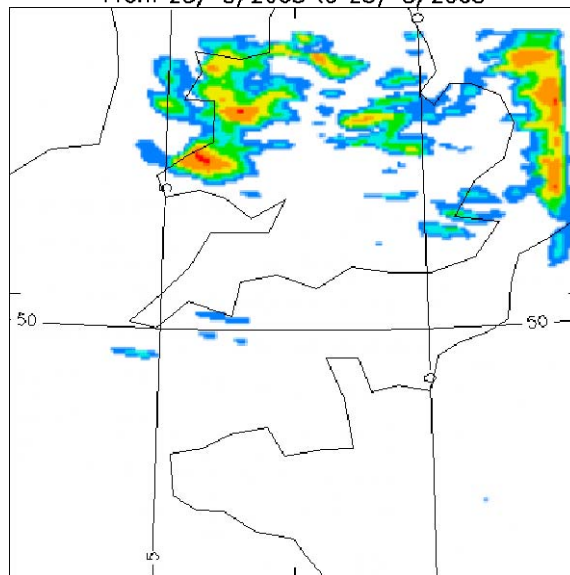
T+2 forecast
1km filter(no filter)
Cloud=3hr rain=1hr

Radar
1 hour accumulation

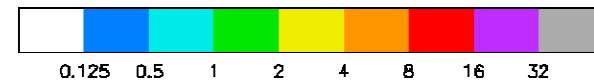
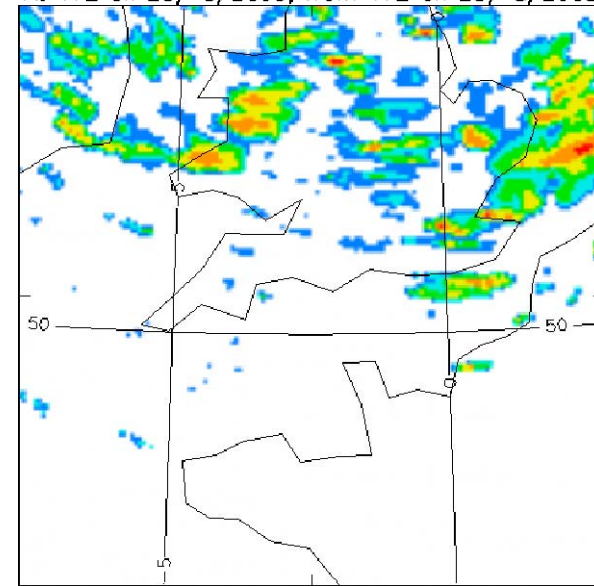
XAPYF Time mean
surface Atmos total precipitation amount kg/m2/1s
From 25/ 8/2005 to 25/ 8/2005



XAQOK Time mean
surface Atmos total precipitation amount kg/m2/1s
From 25/ 8/2005 to 25/ 8/2005



AAAAJ Time mean
surface Atmos total precipitation amount kg/m2/1s
At 17Z on 25/ 8/2005, from 17Z on 25/ 8/2005



- Analysis is very sensitive to specification of background errors and observation errors
- Need to bias correct and quality control data
- Specification of errors tends to be a matter of tuning to find best overall forecast skill score
- Need careful balance of errors and data quantity for different data types to get best forecast

1. Dealing with model error. Statistical characteristics of model error are difficult to determine. However, relatively simple models may bring benefits.
2. Incorporating 'errors of the day'. I.e. making the background error covariances more synoptically dependent. Also tying them to boundary layer depth.
3. Ensemble techniques – enables calculation of background error distribution/covariances – time evolving and allows for observation and forecast uncertainty
4. Variational quality control and bias correction
5. Specification of bias correction, quality control, thinning/superobbing, errors for radar doppler radar winds, reflectivity and refractivity data
6. Allow for correlated observation errors

Questions & Answers