

High resolution simulations of the night-time stable boundary layer over snow

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Motivation

- SBL problematic to understand and model
- GCMs: large T_s errors in polar areas
- NWP: warm biases in cold conditions ($T_{2m} < -20^\circ\text{C}$; Atlaskin and Vihma 2012)
- Snow: good insulation, high thermal inertia
 - low night-time T_s , very stable BL, large NWP biases
- How important is Q_{rad} vs Q_{turb} in the SBL, and why?
 - High-resolution idealized 1-D experiments over snow.
 - Vary the wind speed V_g , look for $T(z)$, $V(z)$, $Q_{\text{rad}}(z)$, $Q_{\text{turb}}(z)$, H_0 .
 - Vary the snow scheme & snow properties.

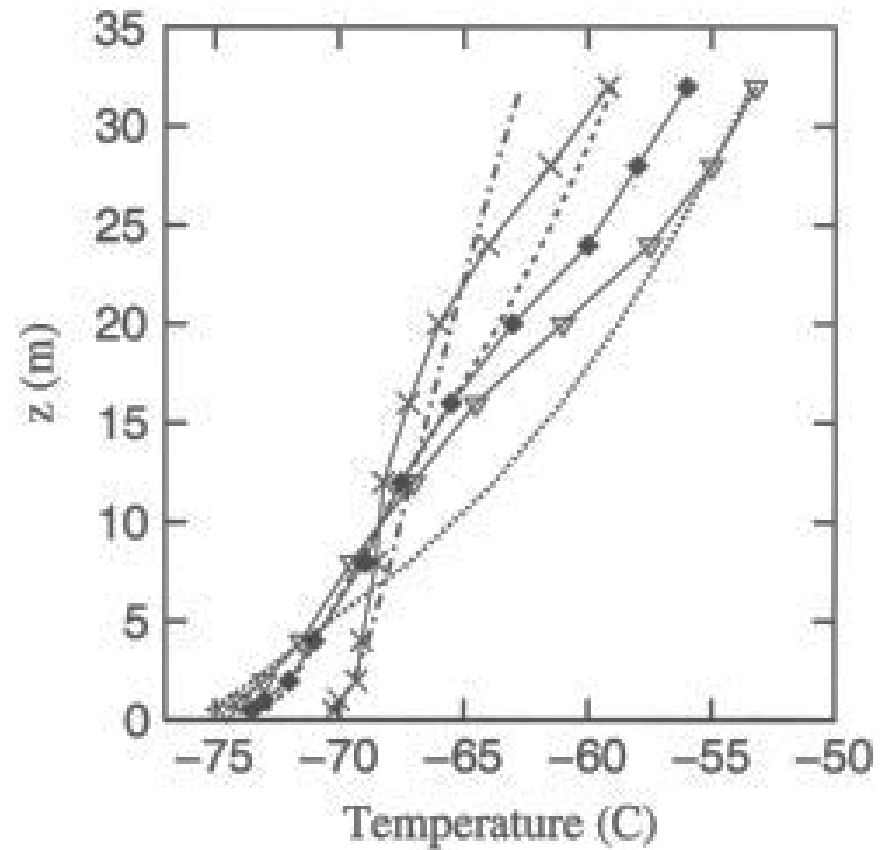
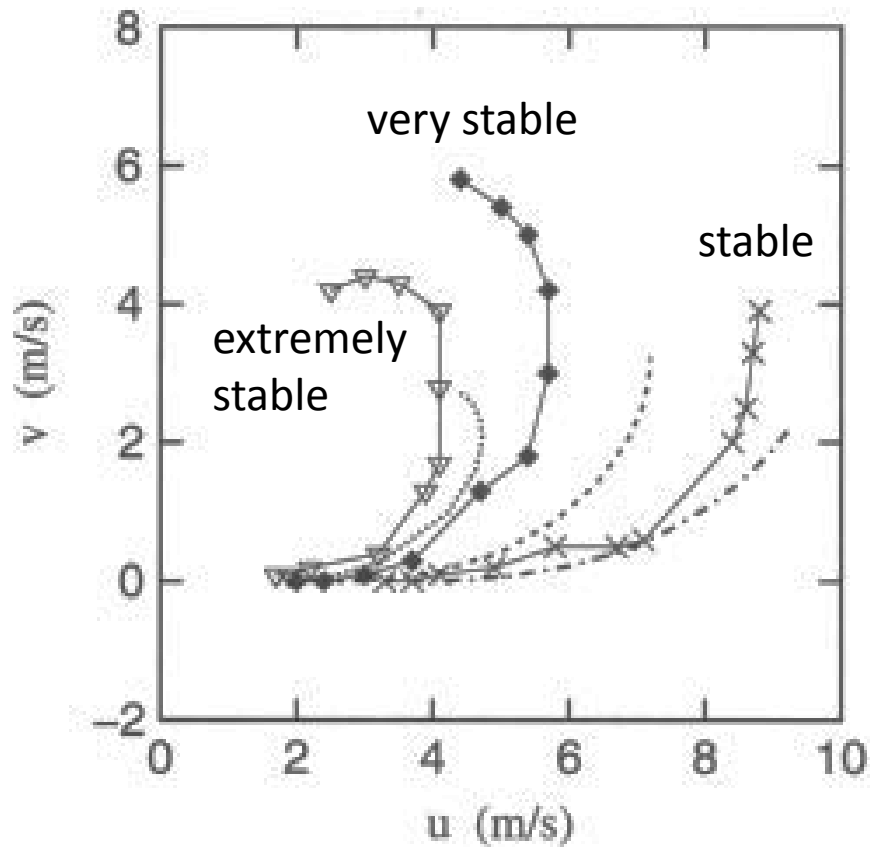
1-D BL model (1)

- 1-D, constant V_g , 10 s time step
- Grid: 0.1 m, 0.3 m, 1, 2, 5, 10, 22 m, ... 30km
- **LW rad**: narrow band model (67 bands, LBL-validated)
SW rad: HIRLAM-type
- **Turbulence**: mixing length Blackadar closure, λ 150 m, $f(Ri) =$ 'short tails' for $Ri > 0$, Dyer-Businger for $Ri < 0$
- **Cloud physics**: saturation at RH 105% (\rightarrow diamond dust in cold air, sedimenting down as observed)

BL model (2)

- **Snow:** thermal diffusion (Crank-Nicholson). **Five optimized depths:** 0, 0.87, 1.85, 6.65, 18.8 cm.
Deep snow: $T = T_{ds} = \text{const}$
- **Scheme max T_s error < 1% of amplitude for diurnal square wave forcing** (cf. 14% in a coarse force-restore, 3% in a 50 layer 0.5 cm gridlength scheme)
- 1-D model tested in midlatitude summer (2006 QJ), in Antarctic winter (2009 QJ)

(and in Mars' very stable nighttime BL: 1-S model vs. Phoenix observations (2010 QJ), vs. Spirit obs (2008, 2012 QJ))

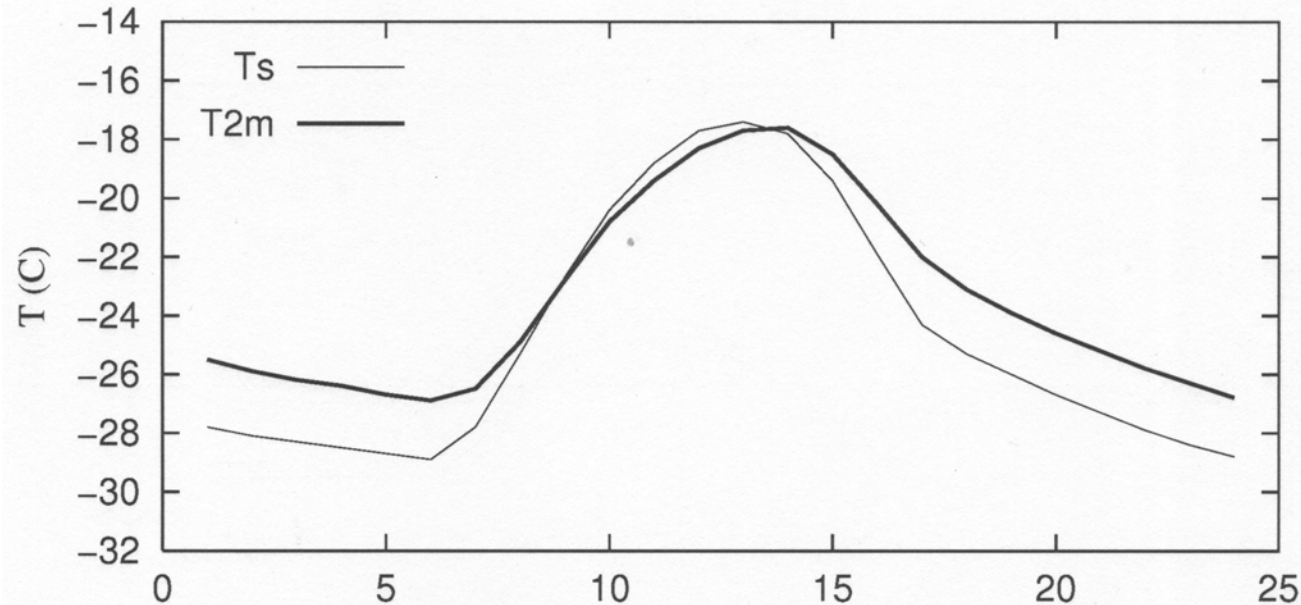


Antarctic High Plateau Station 0-32 m winter mean temperatures and wind hodographs in three stability categories (from King and Turner, 1997) with 1-D simulations (S 2009 QJ)
 (non-flat areas & high Ri \rightarrow local drainage winds \rightarrow $K=K(\sigma(z_s))$)

Present test case

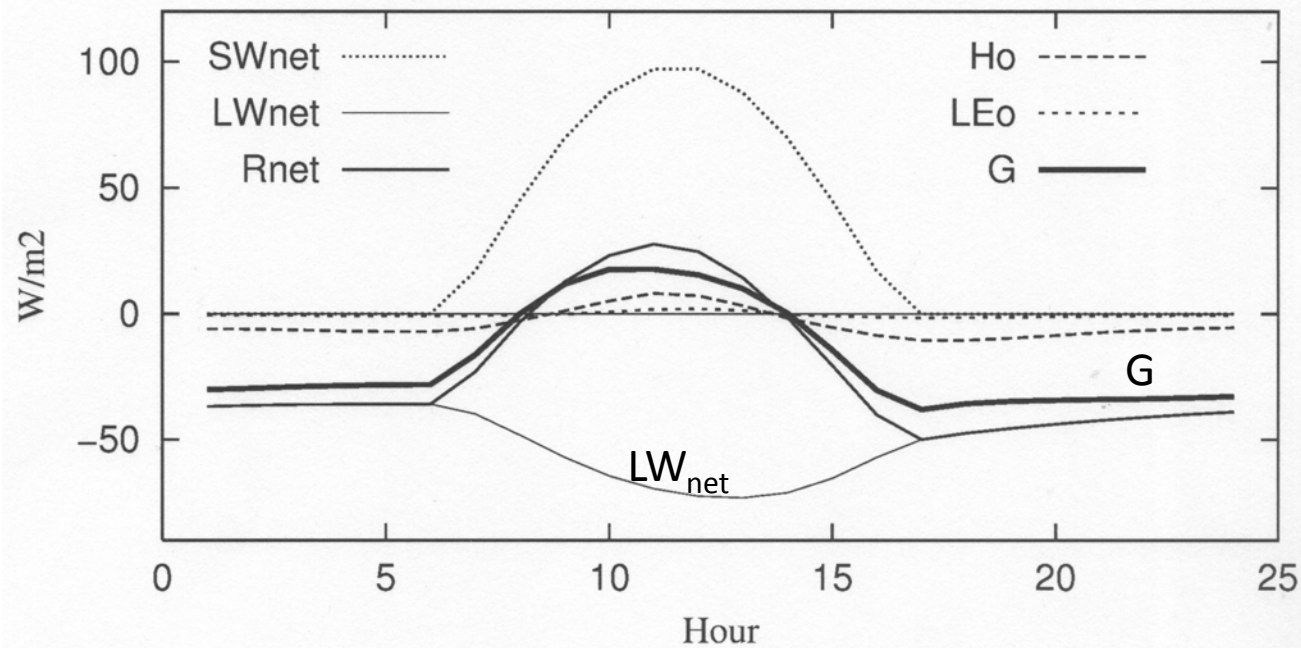
- 67°N, clear-sky early March, snow-covered (50 cm) homogeneous field with typical snow ($z_{0m,h} = 1$ cm, $\alpha = 70\%$), $T_{ds} = -15^\circ\text{C}$.
- Initially: mean Sodankylä winter $T(z)$, 40% RH.
- +54h simulation from 00 local solar time (LT).
- First day: a general Arctic clear-sky inversion to about 300 m is formed, as in the Antarctica case.
- Then: Sunny days \rightarrow shallow well-mixed layer
Clear nights \rightarrow shallow diurnal stable layer, T_s down to $-25\dots-34^\circ\text{C}$ depending on winds ($V_g = 0/1/2/\dots 10$ m/s)

Vg 3 m/s



Day 2, clear sky, weak wind:

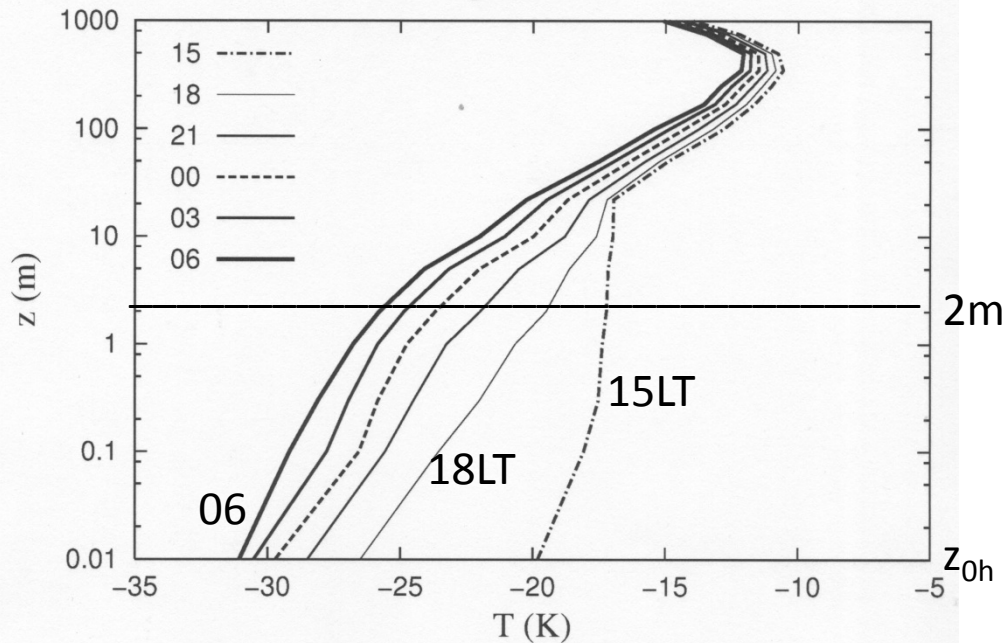
T2m, Ts:
Unstable at 08-14 LT.
 $T_{2m} - T_s$ is about 2 K during the night



Surface heat fluxes:
G: ~ const at night,
 H_o , LE_o small,
 R_{net} dominates in G

→ downwelling LW radiation should be rather accurate for a good G and a good T_s - and T2m prediction.

$V_g = 0.2 \text{ m/s}$



Calm clear-sky night, day 2-3

$T(\log z)$ 15LT \rightarrow 06LT:

General inversion to 300 m

Mixed layer to 20 m at 15LT

Then: Steep inversion

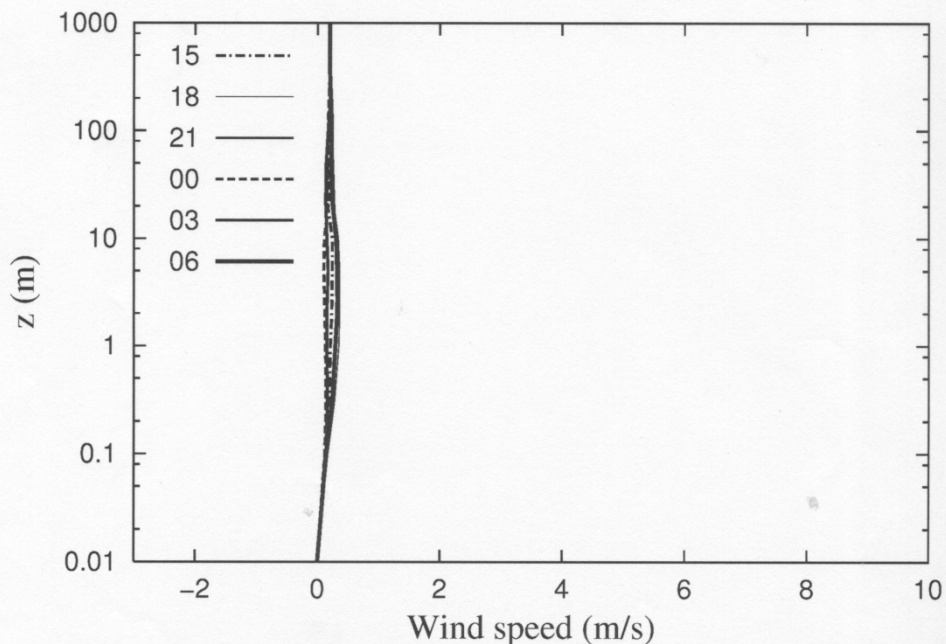
Cold surface, quite logarithmic $T(z)$

$T_{2m} - T_s \sim 6 \text{ K}$

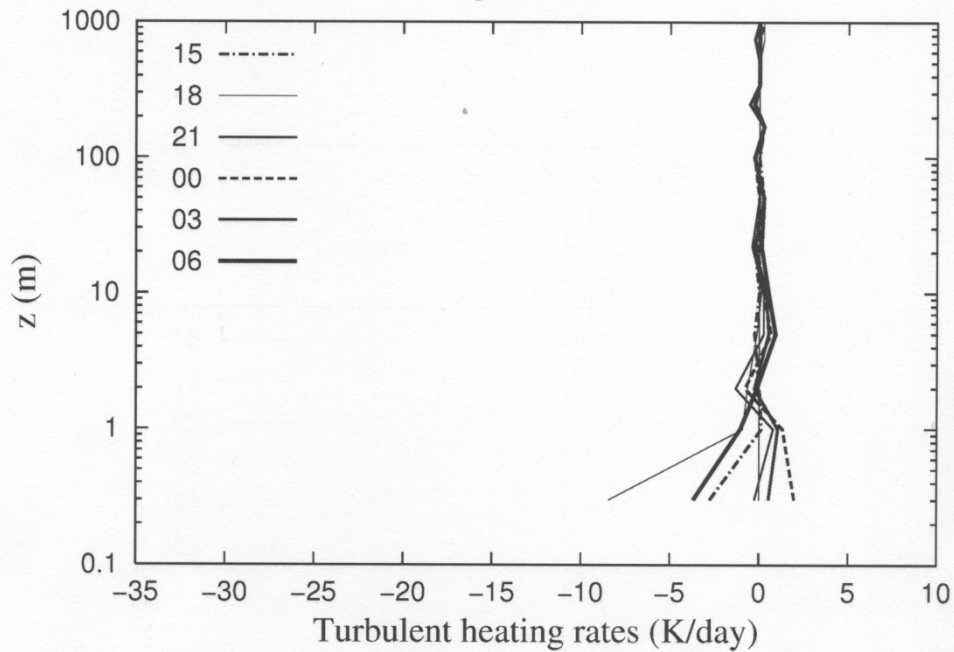
Inversion in $q(z)$, too

$V(\log z)$:

No wind, calm "radiative" night

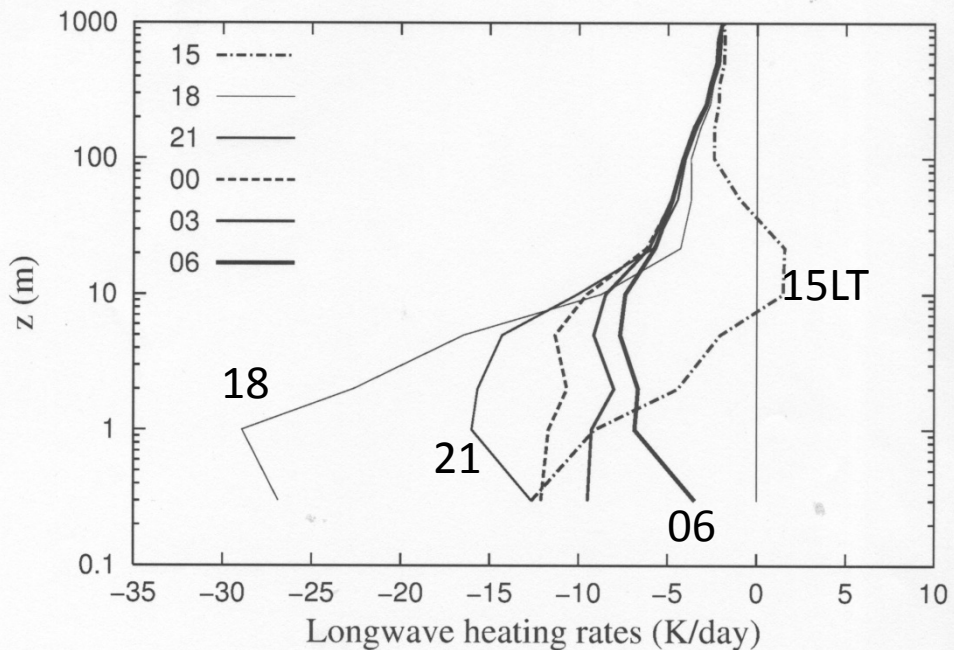


$V_g = 0.2 \text{ m/s}$



Calm clear-sky night

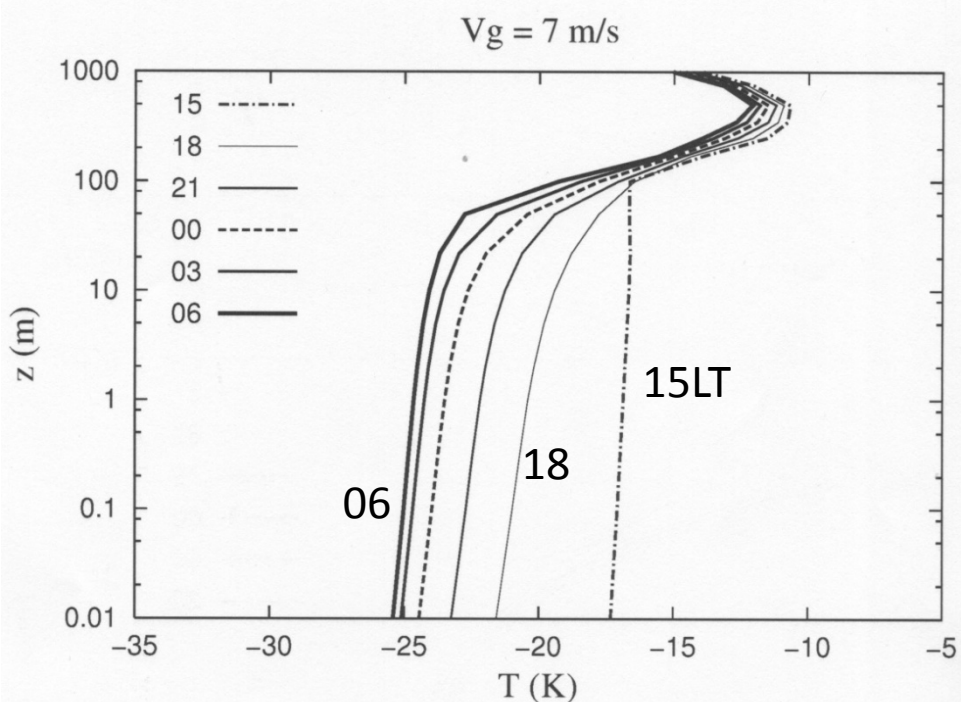
$Q_{\text{turb}}(\log z)$ 15LT \rightarrow 06LT
No wind, no turbulent cooling



Q_{rad} dominates:

In the evening: Strong LW cooling near the surface

During the night: Moderate LW cooling



Windy clear-sky night:

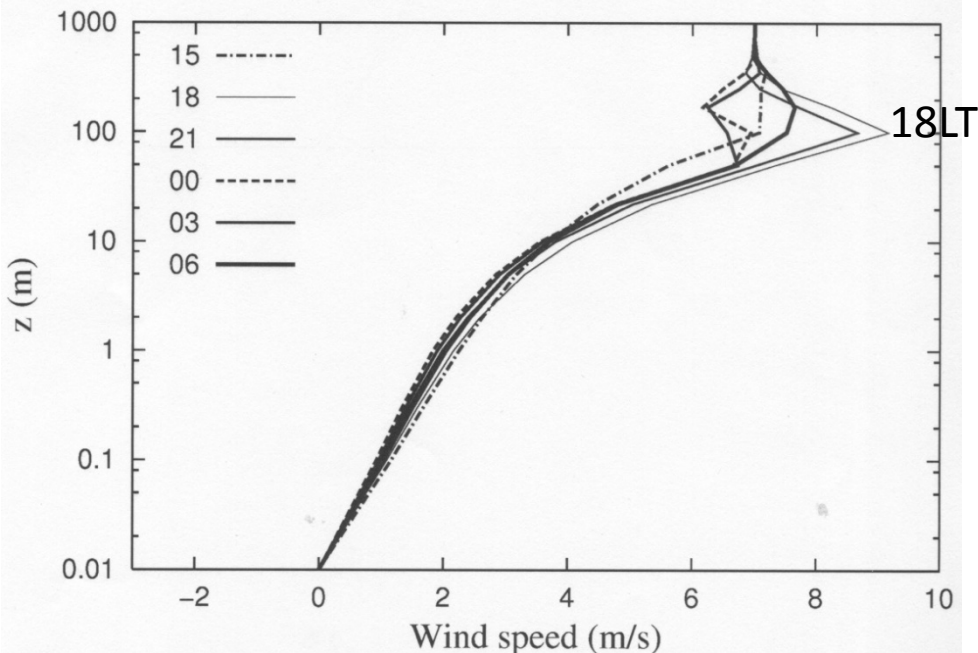
T(log z) 15LT → 06LT:

General inversion to 300 m

Well-mixed BL to 100 m at 15LT.

Nearly isothermal surface layer during the windy night.

T2m is closely coupled to T_s .



V(log z) 15LT → 06LT:

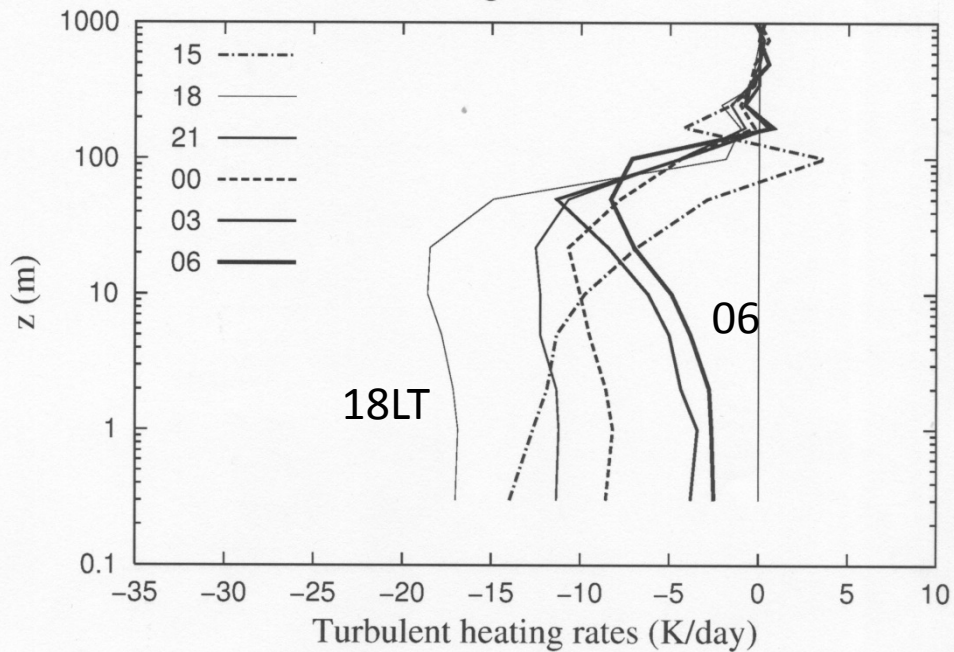
Logarithmic surface layer,
Strong wind shear

Inertial oscillation at 100 m

NLJ at 18-21LT,

Shears time-dependent in
the Ekman layer

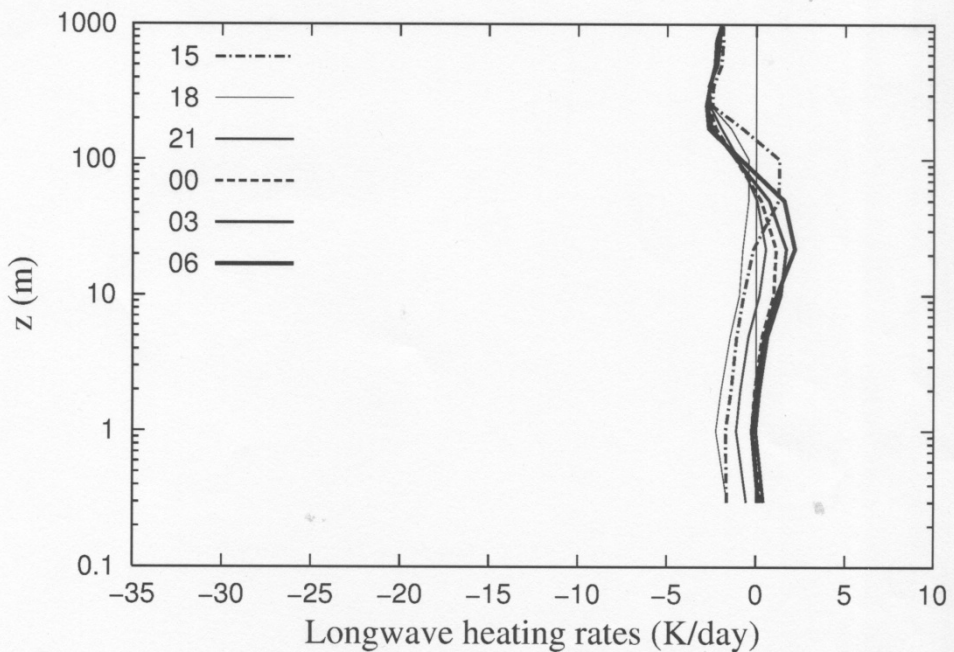
$V_g = 7 \text{ m/s}$



Windy clear-sky night:

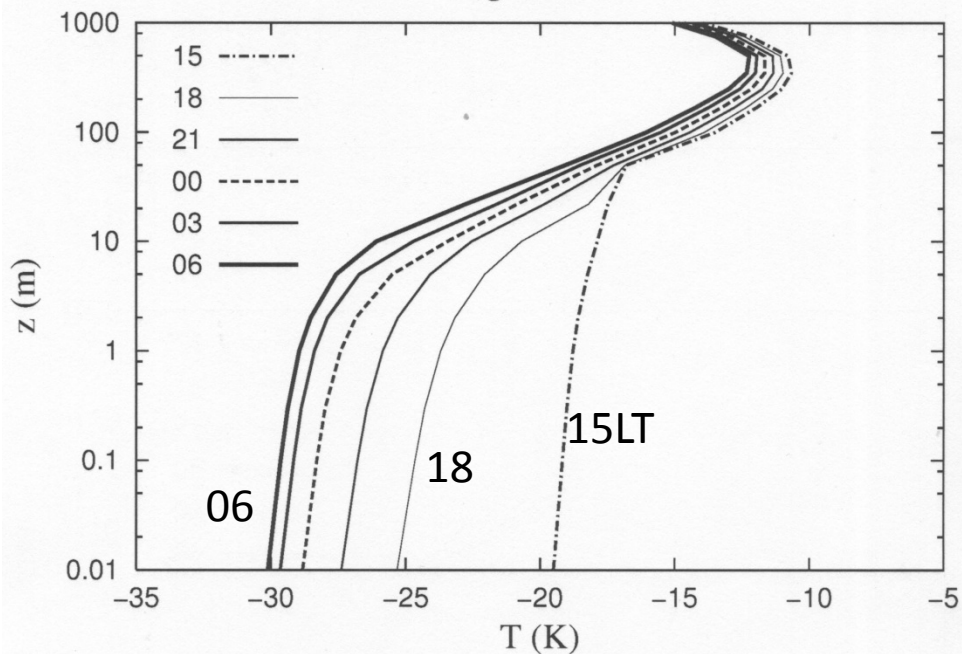
Turbulent cooling dominates

(but is only moderate:
wind shears are strong but
T-gradients are weak)



LW cooling is insignificant in
the windy surface layer
(weak T, q -gradients)

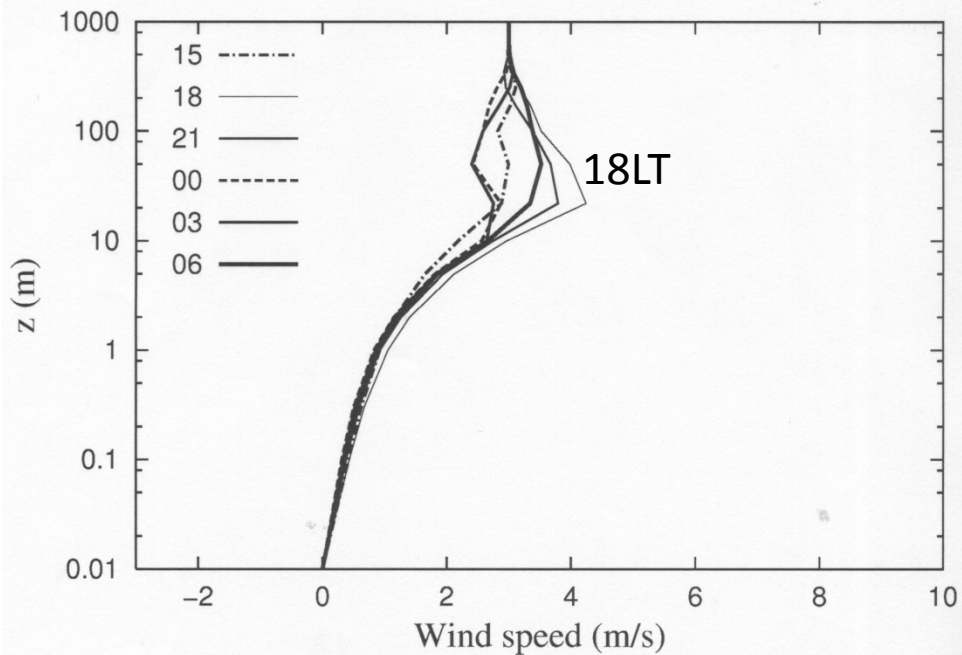
$V_g = 3 \text{ m/s}$



Weak wind, clear sky night:

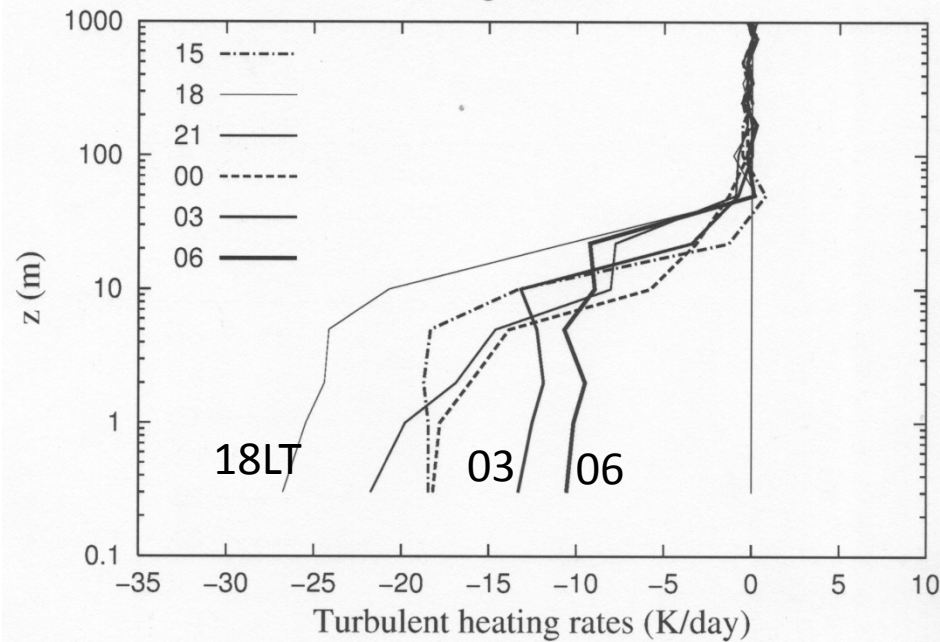
$T(z)$ is well-mixed to 50 m at 15LT, is between isothermal and logarithmic during the night

$$T_{2m} - T_s \sim 2 \text{ K}$$



Wind speed: logarithmic to 5 m, inertial oscillation with weak NLJ at 20-50 m during the night

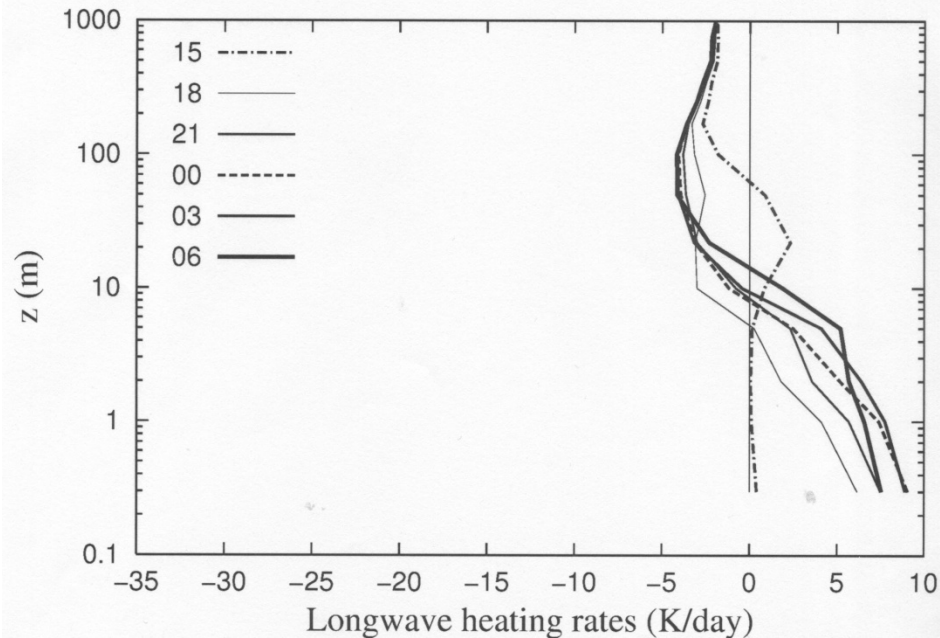
$V_g = 3 \text{ m/s}$



Weak wind clear-sky night:

Turbulent cooling is moderate to strong in the surface layer

(moderate shears, moderate T-gradients)



LW cooling is moderate and dominates the mid-inversion and aloft,
but turns to weak heating in the surface layer

(compensates turb cooling!)

The day 2 0600LT V2m, V10m (ms^{-1}) and T_s , T2m, T10m, T100m ($^{\circ}\text{C}$) as the function of V_g (ms^{-1}).

| V_g | V2m | V10m | T_s | T2m | T10m | T100 m | T2m - T_s |
|------------|-----|------|--------------|--------------|--------------|--------------|-------------|
| 0.0 | 0 | 0 | -29.9 | -23.4 | -19.2 | -11.6 | 6.5 |
| 1.0 | 0.9 | 1.4 | -29.6 | -25.2 | -20.2 | -11.7 | 4.4 |
| 2.0 | 1.1 | 2.3 | -29.3 | -26.5 | -21.8 | -11.9 | 2.8 |
| 3.0 | 1.2 | 2.8 | -28.8 | -26.8 | -23.4 | -12.2 | 2.0 |
| 4.0 | 1.4 | 2.9 | -28.2 | -26.6 | -24.5 | -12.6 | 1.6 |
| 5.0 | 1.7 | 3.2 | -27.5 | -26.0 | -24.6 | -13.0 | 1.5 |
| 6.0 | 1.9 | 3.4 | -26.9 | -25.6 | -24.7 | -13.9 | 1.3 |
| 7.0 | 2.2 | 3.7 | -26.1 | -24.8 | -24.0 | -15.7 | 1.3 |
| 8.0 | 2.6 | 4.1 | -25.4 | -24.1 | -23.5 | -17.3 | 1.3 |

Surprise in the experiments:

The snow surface was at its coldest after a calm night

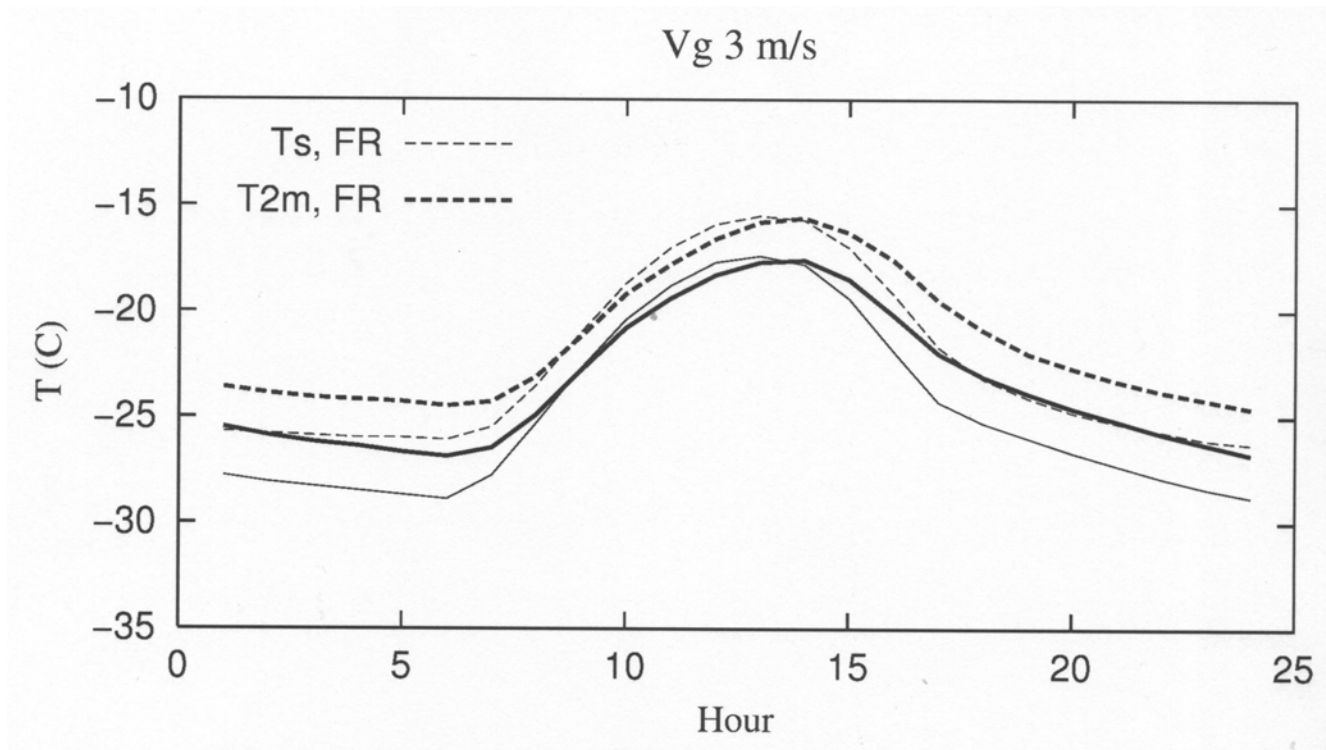
But: T2m was at its coldest during weak wind, V2m of 1.2 m/s
(V_g 3 m/s), not during calms,

and T10m was at its coldest for V10m of 3.5 m/s (V_g 6 m/s)
as was observed in SHEBA.

Why?

A little bit of wind helps to maximize $|H_0|$ (when using the
'short tails' $f(Ri)$, not with 'long tails'), so that the coldness of
the surface is transported effectively to the air and Q_{turb} can
stay strongly negative, against the positive Q_{rad} .

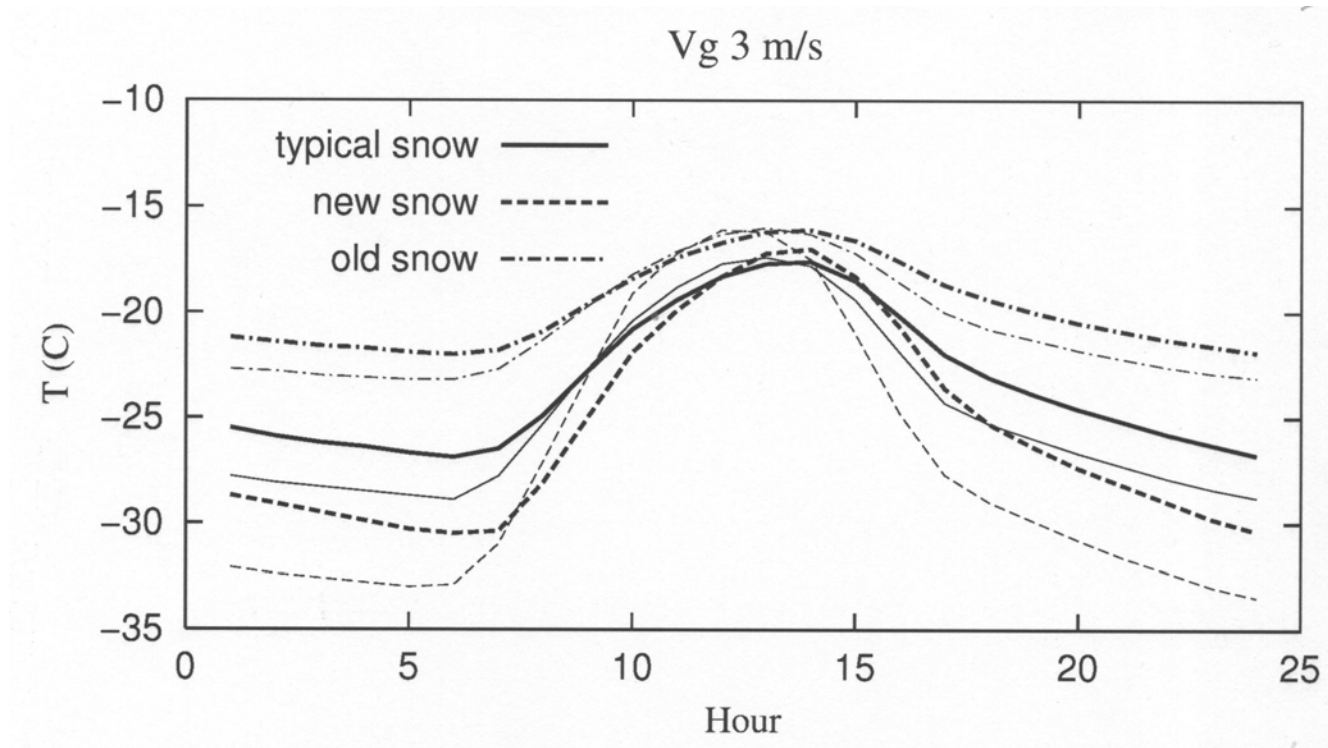
Sensitivity tests: Snow schemes:



A coarse snow scheme (here force-restore = FR) tends to produce too warm night-time temperatures (warm bias).

→ Use multilayer snow schemes

Sensitivity tests: Snow types:



Use of older snow thermal properties (upper curves) when fresher snow actually prevails (lower curves) tends to produce warm bias at night

→ Refresh your snow properties (primarily snow density) by the observed/predicted snowfalls.

Clear-sky nights over snow: conclusions

- Calm night:** LW cooling dominates, leads to nearly logarithmic $T(z)$.
- Windy night:** Turb cooling dominates, $V(z)$ is logarithmic with NLJ, surface layer is \sim isothermal.
- Weak wind:** Turb cooling dominates in the SL against LW warming, LW cooling dominates aloft.
 - T2m is at its coldest as $|H_0|$ is maximized.
 - ($|E_0|$ is also maximized \rightarrow danger of road icing!)
- Coarse snow scheme:** may lead to a warm bias
- Old snow parameters:** may lead to a warm bias
- Inaccurate DLR:** may lead to wrong T_s and T2m.