

INCLUDING CORIOLIS EFFECTS IN THE PRANDTL MODEL FOR KATABATIC FLOW

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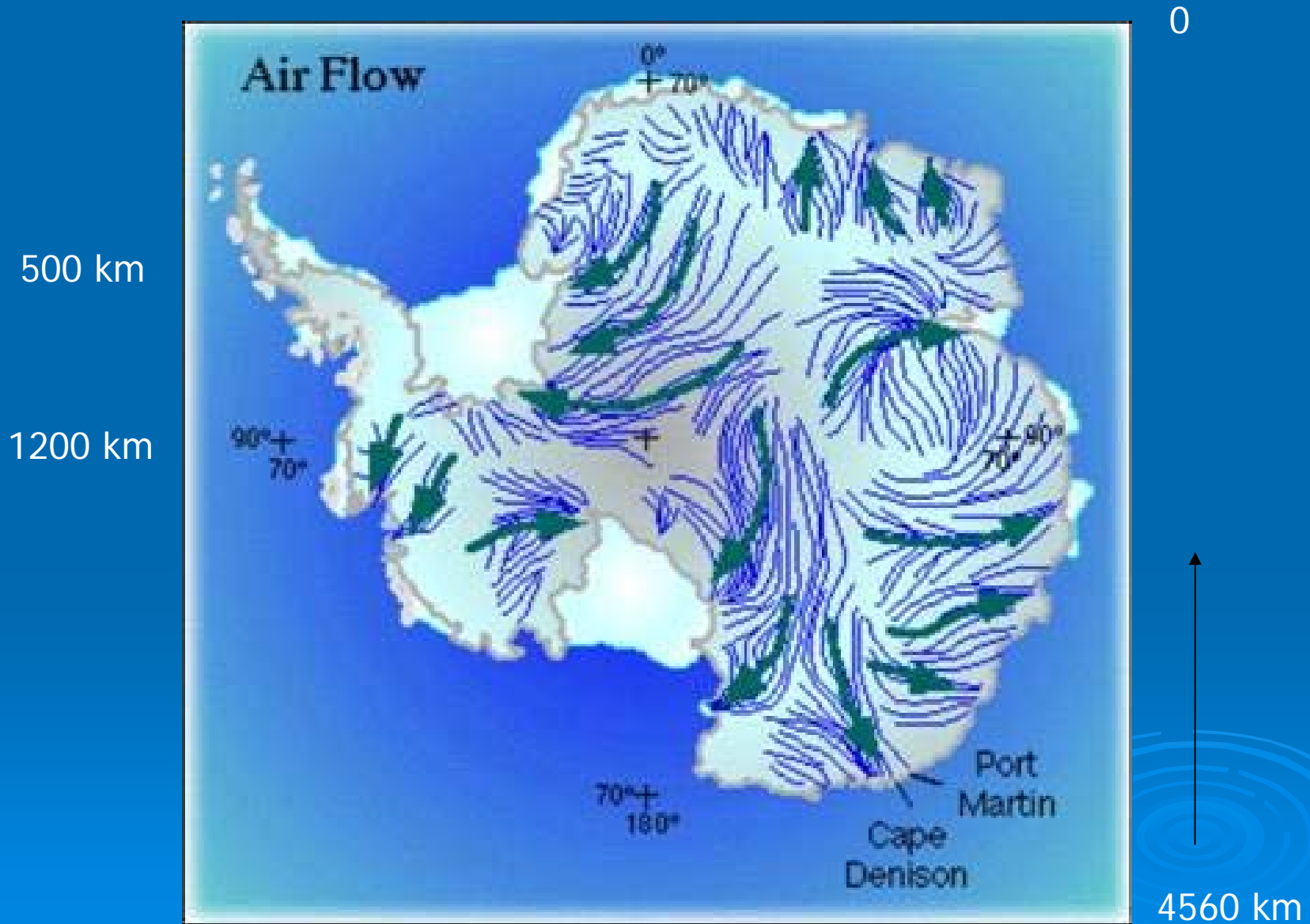
OBJECTIVES

- INTRODUCTION
- ROTATING PRANDTL MODEL
- ASYMPTOTIC TIME-DEPENDENT MODEL AND
NUMERICAL RESULTS
- CONCLUSION

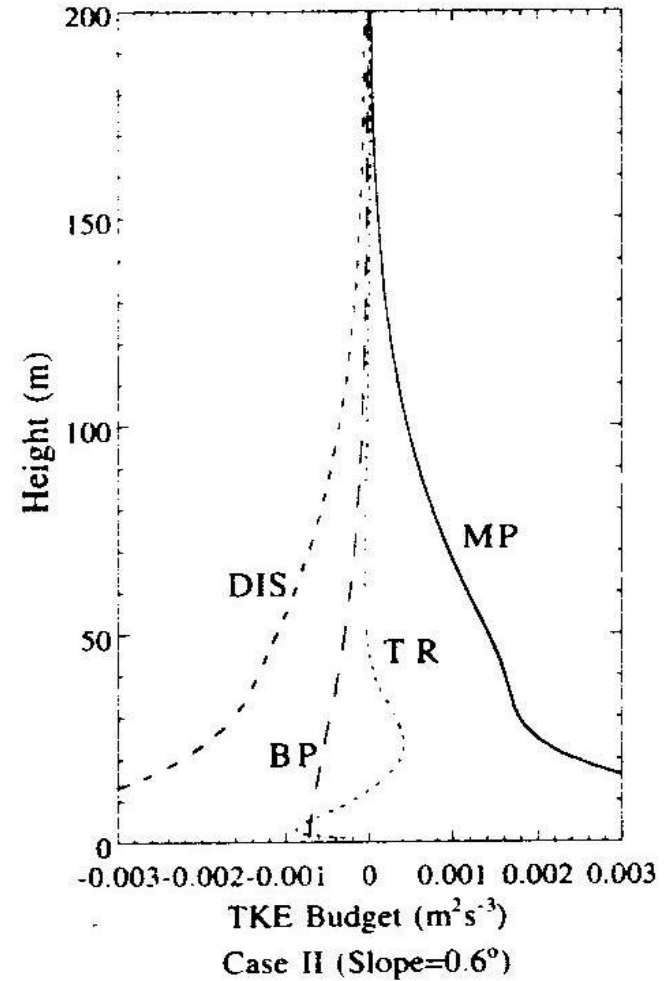
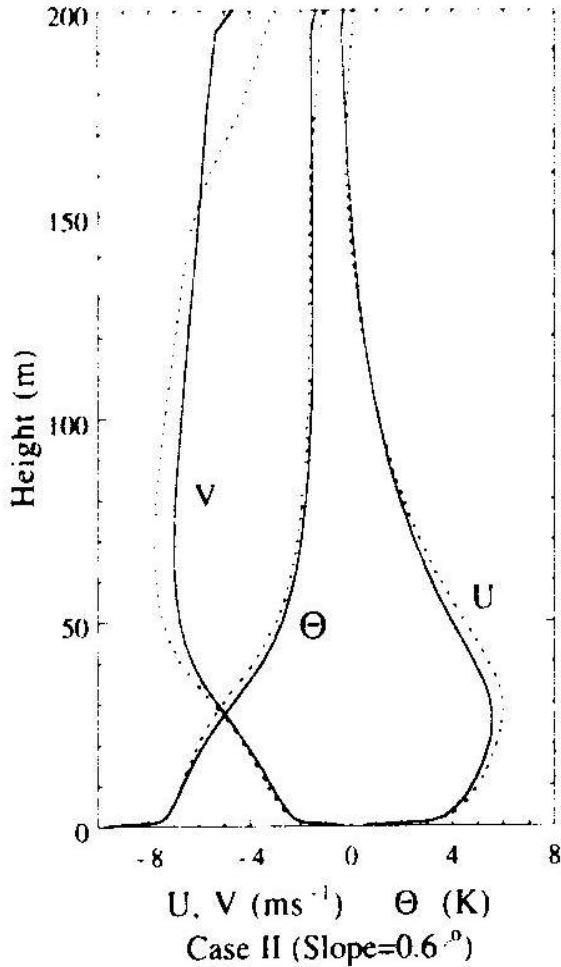
KATABATIC FLOWS - IMPORTANCE

- GLACIERS – climate barometers
- LOCAL CIRCULATIONS – urban effects, traffic, convection...
- CUMMULATIVE EFFECTS IN CLIMATE PROCESSES
 - to be properly parameterized in climate models

KATABATICS + f long-lasting ABL

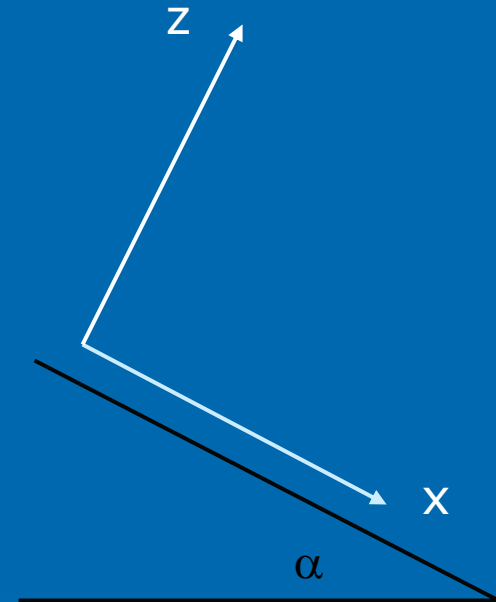


Katabatic wind with rotation



PRANDTL MODEL+ CORIOLIS EFFECT

- hydrostatic balance
- linear (no advection)
- clockwise rotation ($\alpha < 0$)
- constant K^*



* Parmhed et al., QJ04 - Prandtl model with $K(z)$ may be treated via WKB method and applied to a real glacier (e.g. no problem with any large Ri and Pr but missing pressure gradient and variable advection)

X:
$$0 = \frac{g}{\theta_0} \sin\alpha \theta + f \cos\alpha V + KPr \frac{d^2U}{dz^2}$$

Y:
$$0 = -fU \cos\alpha + KPr \frac{d^2V}{dz^2}$$

θ :
$$0 = -U\gamma \sin\alpha + K \frac{d^2\theta}{dz^2}$$

B. C.
$$\begin{aligned} \theta(z = 0) &= \tilde{C}, \quad U(z = 0) = V(z = 0) = 0 \\ \theta(z \rightarrow \infty) &= U(z \rightarrow \infty) = V(z \rightarrow \infty) = 0 \end{aligned}$$

$$F = (\theta, U, V)$$

$$\frac{d^2}{dz^2} \left(\frac{d^4 F}{dz^4} + \sigma^4 F \right) = 0$$

 \Rightarrow

$$\frac{d^4 F}{dz^4} + \sigma^4 F = az + b$$

- B. C. give $a = b = 0$

$$\tilde{C} = \frac{C}{1 + \Delta}$$

$$\Delta = \frac{f^2 \cot^2 \alpha}{N^2 \text{Pr}}$$

$$\sigma^4 = \frac{N^2 \text{Pr} \sin^2 \alpha}{K^2 \text{Pr}^2} (1 + \Delta)$$

STEADY-STATE SOLUTIONS FOR U AND θ

- Analogy to classic Prandtl model

$$\theta_s = \tilde{C} e^{\frac{-\sigma Z}{\sqrt{2}}} \cos\left(\frac{\sigma Z}{\sqrt{2}} + \Delta\right),$$

$$U_s = \frac{\tilde{C} K \sigma^2}{\gamma \sin \alpha} e^{\frac{-\sigma Z}{\sqrt{2}}} \sin\left(\frac{\sigma Z}{\sqrt{2}}\right)$$

$$V_s = \frac{\tilde{C} f \cot \alpha}{\text{Pr} \gamma} \left(1 - e^{\frac{-\sigma Z}{\sqrt{2}}} \cos\left(\frac{\sigma Z}{\sqrt{2}}\right) \right)$$

NUMERICAL UNSTEADY SOLUTIONS

- Main time scale: $T=2\pi/[N\sin\alpha]$

$$X: \quad \frac{\partial U}{\partial t} = \frac{g}{\theta_0} \sin\alpha \theta + f \cos\alpha V + KPr \frac{\partial^2 U}{\partial z^2}$$

$$Y: \quad \frac{\partial V}{\partial t} = -fU \cos\alpha + KPr \frac{\partial^2 V}{\partial z^2}$$

$$\theta: \quad \frac{\partial \theta}{\partial t} = -U \gamma \sin\alpha + K \frac{\partial^2 \theta}{\partial z^2}$$

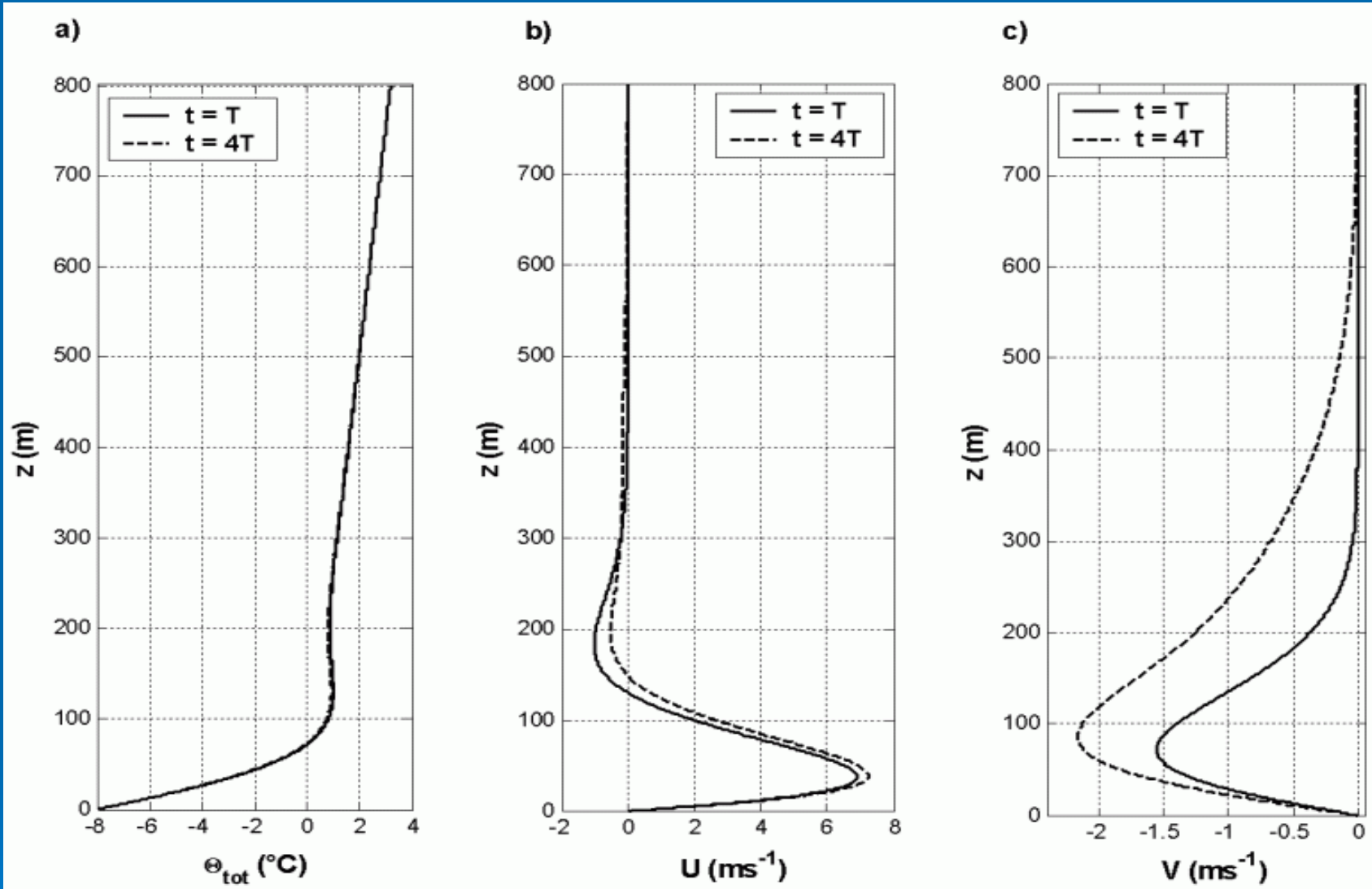


Figure 1. Numerical solution for time-varying Prandtl model with f , $T=2.1\text{h}$, $\alpha, \gamma, K, Pr, C, f = -4^\circ, 4\text{K/km}, 1\text{m}^2\text{s}^{-1}, 1.1, -8\text{K}, 1.1 \cdot 10^{-4}\text{s}^{-1}$

QUASI-UNSTEADY SOLUTION

- Numerical results and scale analysis: U and θ become nearly steady after $T \approx 1/N \sin \alpha$, **but not V !**
- Thus, we solve a simplified time-varying '**Prandtl** + **f'** ' problem analytically and compare it to the fuller, numerical solution shown

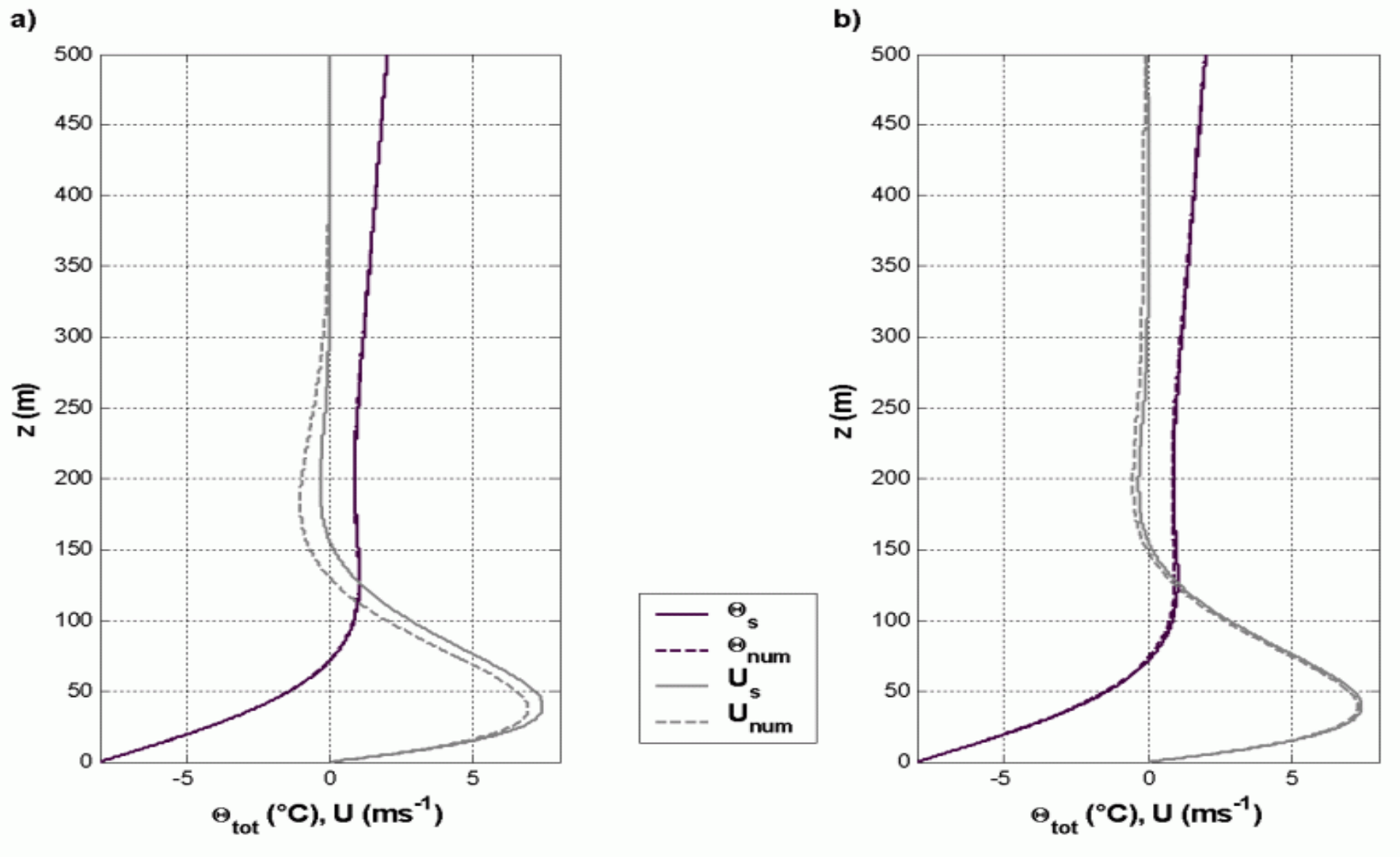


Figure 2. Numerical (dashed) and approximate (solid) steady solutions for the Prandtl model, at (a) $t = T$ and (b) $t = 4T$. The rest as in Fig. 1.

V: diffusion equation forced by Prandtl model

$$\frac{\partial V}{\partial t} - K \text{Pr} \frac{\partial^2 V}{\partial z^2} = -f \cos \alpha U_s, \quad t > T$$

$$V = V_0 \left[1 - \text{erf} \left(\frac{z}{2\sqrt{\tau K \text{Pr}}} \right) - e^{\frac{-\sigma z}{\sqrt{2}}} \cot \left(\frac{\sigma z}{\sqrt{2}} \right) \right], \quad \tau = t - T$$

$$V_0 = \frac{Cf \cot \alpha}{\text{Pr} \gamma}$$

- $U(z,t)$ and $\theta(z,t)$ kept as before from the classic Prandtl model
- V-effect in the x-momentum, since weakest, is dropped:

$$f \theta_0 / \alpha g \theta \sim O(10^{-2})$$

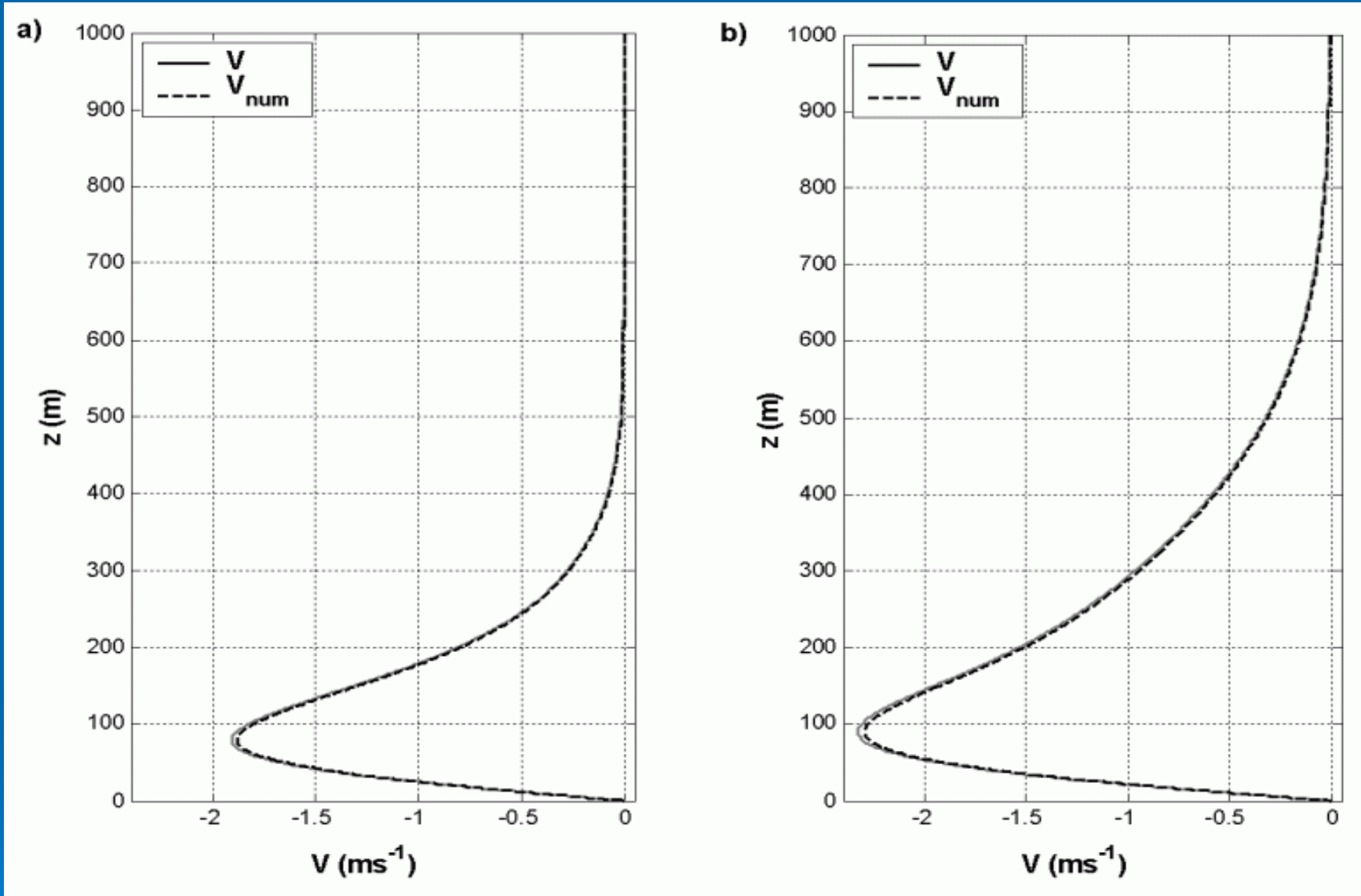


Figure 3. Numerical (dashed) and time-dependent V (solid) solutions obtained for (a) $t = 2T$ and (b) $t = 6T$. The rest as in Fig. 1.

CONCLUSION

- Steady Prandtl model + f is not equivalent to its time dependent counterpart
- U and θ reach steady state profiles after T
- V diffuses upwards in time without well defined time scale
- Approximate quasi-unsteady system (U, V, θ) is in agreement with numerical solution after $t > T$
- We got a closed form for new simple “kata-parameterization”