Atmospheric transport and deposition of coarse solid particles: relations with PBL (and underlying terrain)

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Basic: turbulent dispersion depending on PBL stability

Flux of a passive admixture to the surface:

J = w'C'

Or in gradient-transfer theory formulation:

$$J = -K_{(h)} \partial C / \partial z$$

Often taken the same *K* as for heat:

$$H = -K_h \rho c_p \,\partial \overline{\theta} / \partial z \qquad \text{or} \qquad H = \rho c_p \,\overline{w' \theta'}$$

Briggs' dispersion parameters: "Old stuff", but still useful in many cases.

For idealised Gaussian dispersion $\sigma_i = (2K_i t)^{1/2}$ for plume

$$C = \frac{Q}{2\pi\sigma_z\sigma_y u} \exp\left[-\frac{y^2}{2\sigma_y^2} - \frac{(z-H)^2}{2\sigma_z^2}\right],$$

but real stratification makes certain correctives and therefore Briggs (1970's) proposed on experimental basis, e.g. for open country for small releases (0.1 < x < 10 km):

$$\sigma_z = 0.06x(1+1.5x)^{-1/2}$$
Pasquill stability "D"
$$\sigma_z = 0.03x(1+0.3x)^{-1}$$
Pasquill stability "E"
$$\sigma_z = 0.016x(1+0.3x)^{-1}$$
Pasquill stability "F"





Location of snow sampling sites, Narva power plants (PP-s) and minor industrial sources

Deposition fluxes (average and standard deviation) of fly ash estimated from measurements and computed applying SILAM, December 2 - 14, 2002. Number of samples is indicated in brackets.

Site type (number of samples)	Measured deposition flux, mg/m ² per day, <u>based on</u> :						SILAM:			
	Ca ²⁺		spheroidal particles		total mass		deposition flux, mg/m² per day		concen- tration, µg/m ³	
Woodland (6)	29.0	±4.5	30.2	±5.4	25.6	±2.2	3.7	±0.3	6.0	±0.5
Open land (5)	28.3	±3.6	38.5	±4.1	26.5	±1.7	3.7	±0.3	6.0	±0.5



0.1 0.5 2

7 10

4

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7 10

4

SILAM dispersion and deposition, HIRLAM meteo

December 2 – 14, 2002

Vertical dispersion



SILAM (FMI, Finland)



AEROPOL (Tartu Observatory, Estonia)



AEROPOL: HIRLAM meteo, Gaussian reflection with gravitational sedimentation only, 8.15 μ m ash particles (ρ = 2800 kg/m³).

Reflection, partial or complete adsorption?





AEROPOL:

HIRLAM meteo, complete Gaussian adsorption, gravitational sedimentation and vertical flow, 8.15 μ m ash particles ($\rho = 2800 \text{ kg/m}^3$).

What a kind of particles?

The main admixture in flue gases is water vapour – about 60 g/m³ or 30 times more than fly ash.

Cooling rapidly from +300 °C down to -10...-20 °C that water most likely gets frozen onto the particles making them much larger. v_d must increase.





AEROPOL:

HIRLAM meteo, complete Gaussian adsorption, gravitational sedimentation and vertical flow, 25 μ m ice/ash particles (ρ = 1000 kg/m³).

Are HIRLAM met. data correct?



Wind roses

Surface heat fluxes

(MRF - http://www.arl.noaa.gov/ready)

HIRLAM \rightarrow mainly Pasquill stability "D" MRF \rightarrow often Pasquill stability "E" or "F"



AEROPOL: <u>MRF meteo</u>, complete Gaussian adsorption, gravitational sedimentation and vertical flow, 25 μ m ice/ash particles ($\rho = 1000 \text{ kg/m}^3$).



Deposition flux, computed for measurement point No. 1 (AEROPOL, 25 μm particles) Deposition fluxes: HIRLAM vs. MRF met. data, different options (AEROPOL)

met. data	HIRLAM	MRF	
options			
Reflection, 8.15 µm particles	1.4		3.9
Adsorption, 8.15 µm particles	3.5		5.3
Adsorption, 25 µm particles	4.9		17.0
Measured	Ca-based:		28.7
	Sphparticles-based		34.0
	Total mass-based		26.0

Conclusions

So large deposition fluxes are not possible otherwise than particles must be concentrated into a thin (compared to the stack height) near-surface layer.

Thus, forced mixing of plume within PBL is not always justified in a lower meso-scale model, but vertical dispersion must be treated carefully!

More complex measurements and modelling exercises are needed.

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