

# **Energetics of turbulence, Ri-critical, and new closure model for stable stratification**

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# References

- Zilitinkevich, S. S., Elperin, T., Kleeorin, N., Rogachevskii, I., Esau, I., Mauritsen, T., Miles, M., 2006: Critical Richardson numbers: does it exist? Submitted to *Science*.
- Zilitinkevich, S. S., Elperin, T., Kleeorin, N., & Rogachevskii, I., 2005: Energy- and flux-budget (EFB) turbulence closure model for stably stratified flows. *Boundary-Layer Meteorol.* DOI: 10.1007/s10546-007-9189-2.
- Mauritsen T., Svensson G., Zilitinkevich S. S., Enger L., & Grisogono B., 2007: A total turbulent energy closure model for neutrally and stably stratified atmospheric boundary layers. *J. Atmos. Sci.* In press.



# Content

## Traditional approach

Richardson (1920):  $Ri$

Kolmogorov (1941): TKE budget + closure hypotheses  $\rightarrow \tau = K_M dU/dz, F_b = -K_H db/dz$

$Ri_c$  follows from (i) TKE budget and (ii) perturbation analysis ( $Ri_c = 0.25$ )

Not applied in practice: to get rid of  $Ri_c$  turbulence closures use correction functions

## Proposed

**Turbulent potential and total energies: TTE=TKE+TPE**

**No critical  $Ri$  but a threshold between two turbulent regimes**

- strong, chaotic
- weak, intermittent (wave-dominated)



# Key question

Buoyancy

$$b=(g/\rho_0)\rho=\beta\theta$$

$$(\beta=g/T_0)$$

Brunt-Väisälä frequency

$$N^2=db/dz=\beta d\theta/dz$$

Shear

$$S=dU/ds$$

1920: Richardson number

$$Ri=(N/S)^2$$

## Can local shear maintain turbulence at large Ri?

**‘No’ from classical energy- and perturbation-analyses** (Richardson, 1920; Taylor, 1931; Prandtl, 1930, 1942; Chandrasekhar, 1961; Miles, 1961, 1986;...)

**‘Yes’ from overwhelming majority of experiments, LES and DNS**



# Perturbation analysis

Infinitesimal perturbations are indeed stable at  $Ri > Ri_c$  ( $Ri_c \sim 0.25$ )

## But finite perturbations could be unstable...

**Mechanism** → Internal waves – orbital motions – shear instability – turbulent patches (Phillips 1966, Miropolsky, 1981)

**$Ri \gg 1$**  → Turbulence in  
deep ocean (Thorpe, 2005)  
free atmosphere (Lawrence et al., 2004; Tokovinin, 2005)

**Spectral aspect** → short-wave perturbations always stable: stratification shifts stable modes to lower frequencies (Sun, 2006)



# Problems in currently used closure schemes

To afford sufficient turbulent mixing at large Ri (needed for realistic modelling) turbulence closures are equipped with

- Ri-dependencies of  $K_{\{M,H\}}/K_{\{M,H\}\text{neutral}}$  and  $\text{Pr}_T \equiv K_M / K_H$   
preventing appearance of  $\text{Ri}_c$
- and/or **non-zero background turbulent diffusivities**  
(e.g., Shir and Bornstein, 1977 – free atmosphere;  
Kantha and Klayson, 2000 – deep ocean)



# Turbulent potential energy

$$\rho' = \frac{\partial \bar{\rho}}{\partial z} \delta z$$

density fluctuation

$$\delta E_P = \frac{g}{\rho_0} \int_z^{z+\delta z} \rho' dz = \frac{1}{2} \frac{b'^2}{N^2} = \frac{1}{2} \frac{\beta \theta'^2}{N^2}$$

potential energy  
fluctuation



# Budget equations for

$$\mathbf{TKE}, \quad E_K = \frac{1}{2} \overline{u'_i u'_i}$$

$$\frac{DE_K}{Dt} + \frac{\partial \Phi_K}{\partial z} = \tau S + F_b - D_K, \quad \text{where} \quad \tau = \overline{u'_i w'}$$

$$\mathbf{TPE}, \quad E_P = \frac{1}{2} N^{-2} \overline{b'^2}$$

$$\frac{DE_P}{Dt} + \frac{\partial \Phi_P}{\partial z} = -F_b - D_P, \quad \text{where} \quad F_b = \overline{b' w'}$$

$$\mathbf{TTE}, \quad E_T = E_K + E_P$$

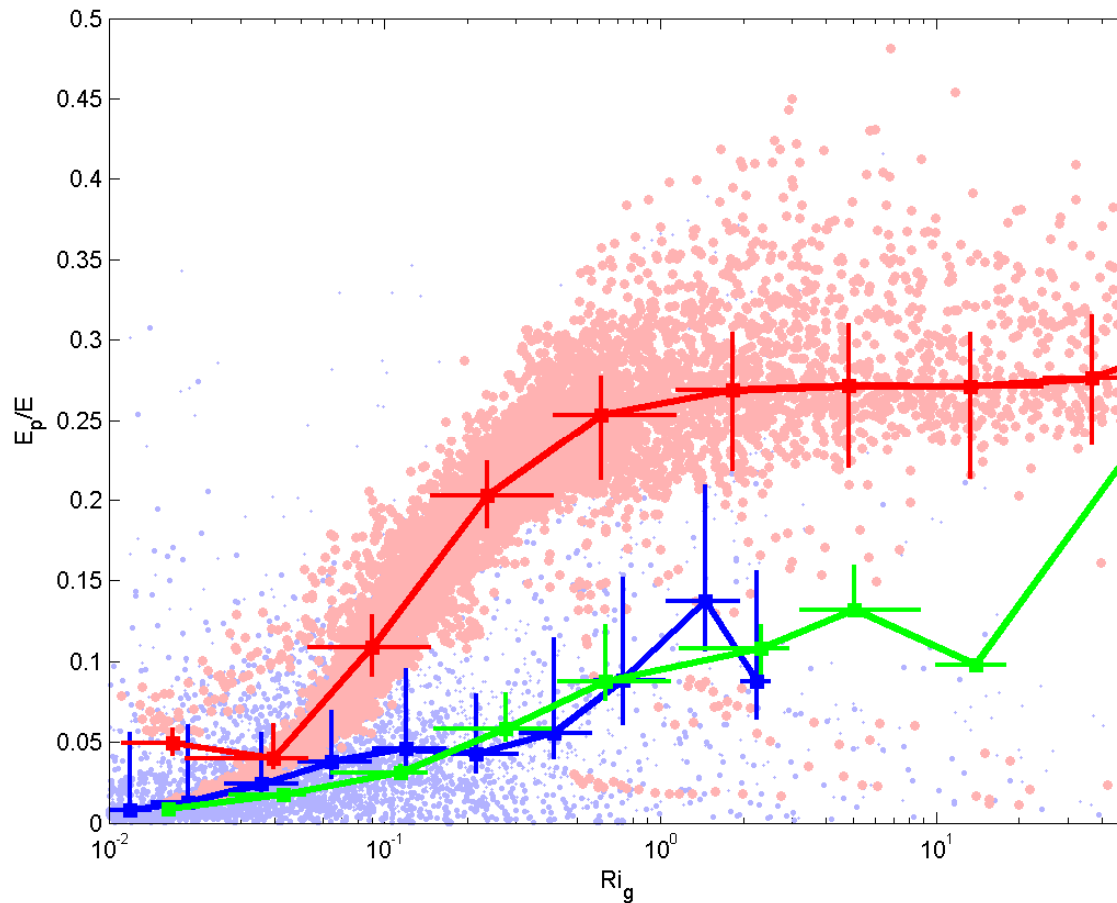
$$\frac{DE_T}{Dt} + \frac{\partial \Phi_T}{\partial z} = \tau S - D_T \quad \text{does not include } F_b!$$





# Role of TPE ( $TPE/TTE = Ri_f = -\beta F_\theta / \tau S < 1$ )

TPE / TKE vs. Ri: lab (Ohya, 2001), LES (Esau)



for LES  
transition at  $Ri=0.25$



# Classical (but erroneous) energy analysis

## From TKE equation:

At large  $Ri$  ( $Ri > Ri_c$ ),  $F_b$  passes a threshold, after which the TKE production,  $P = \tau S$ , is insufficient to compensate its consumption  $\rightarrow$  TKE decays.

## However

The TKE budget equation is not closed. It only shows that **flux Richardson number**,  $Ri_f = -F_b / (\tau S)$  cannot exceed unity.

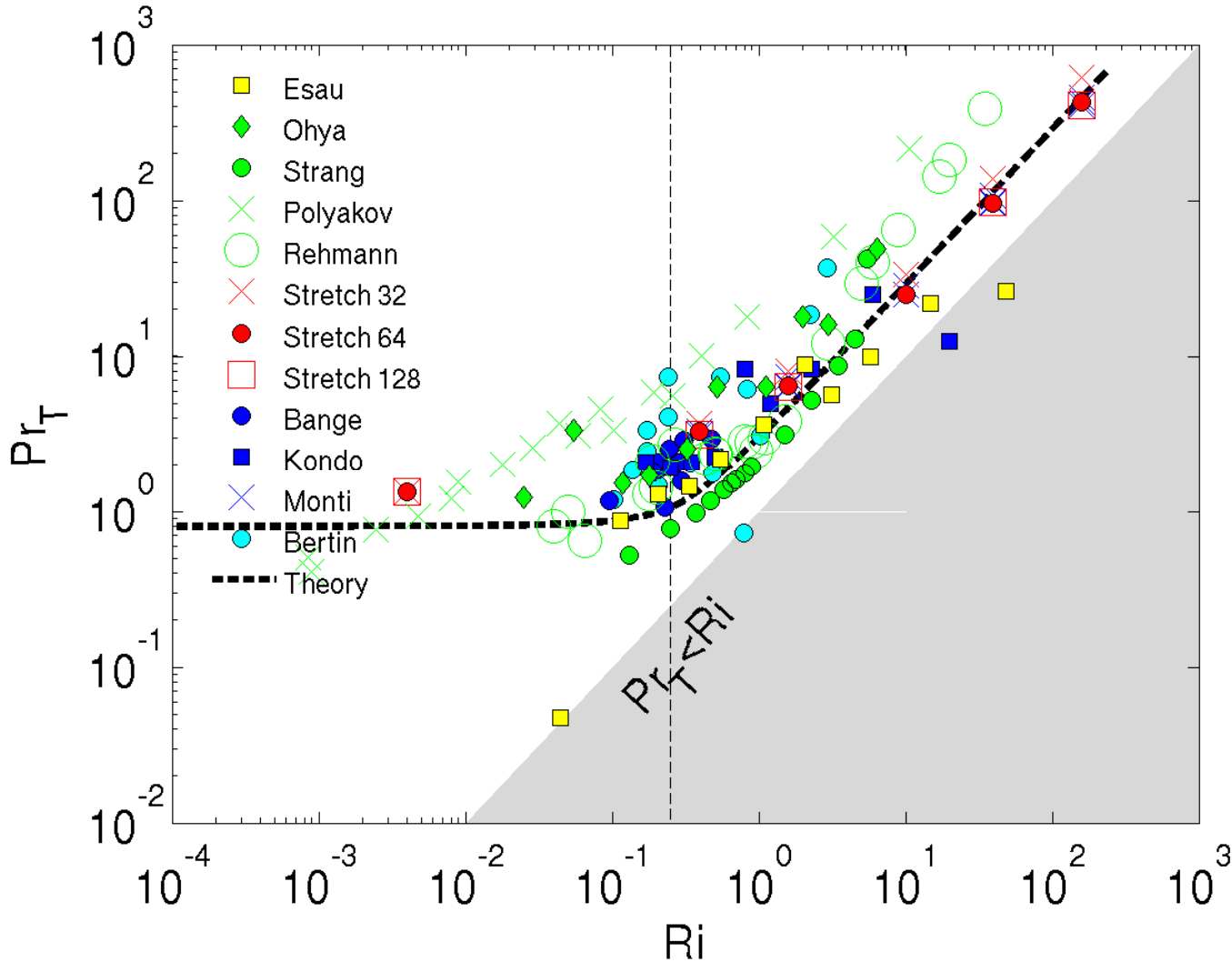
But  $Ri_f$  is an internal turbulent parameter ( $F_b$  and  $\tau$  depend on each other) Inequality  $Ri_f < 1$  does not say anything about maintenance or degeneration of turbulence at large  $Ri$ .

## Traditional derivation of $Ri_c$ assumes

$\tau = K_M dU/dz$ ,  $F_b = K_H db/dz$ , with  $Pr_T = K_M / K_H = \text{constant}$   
**which contradicts** to experimental data (next slides)



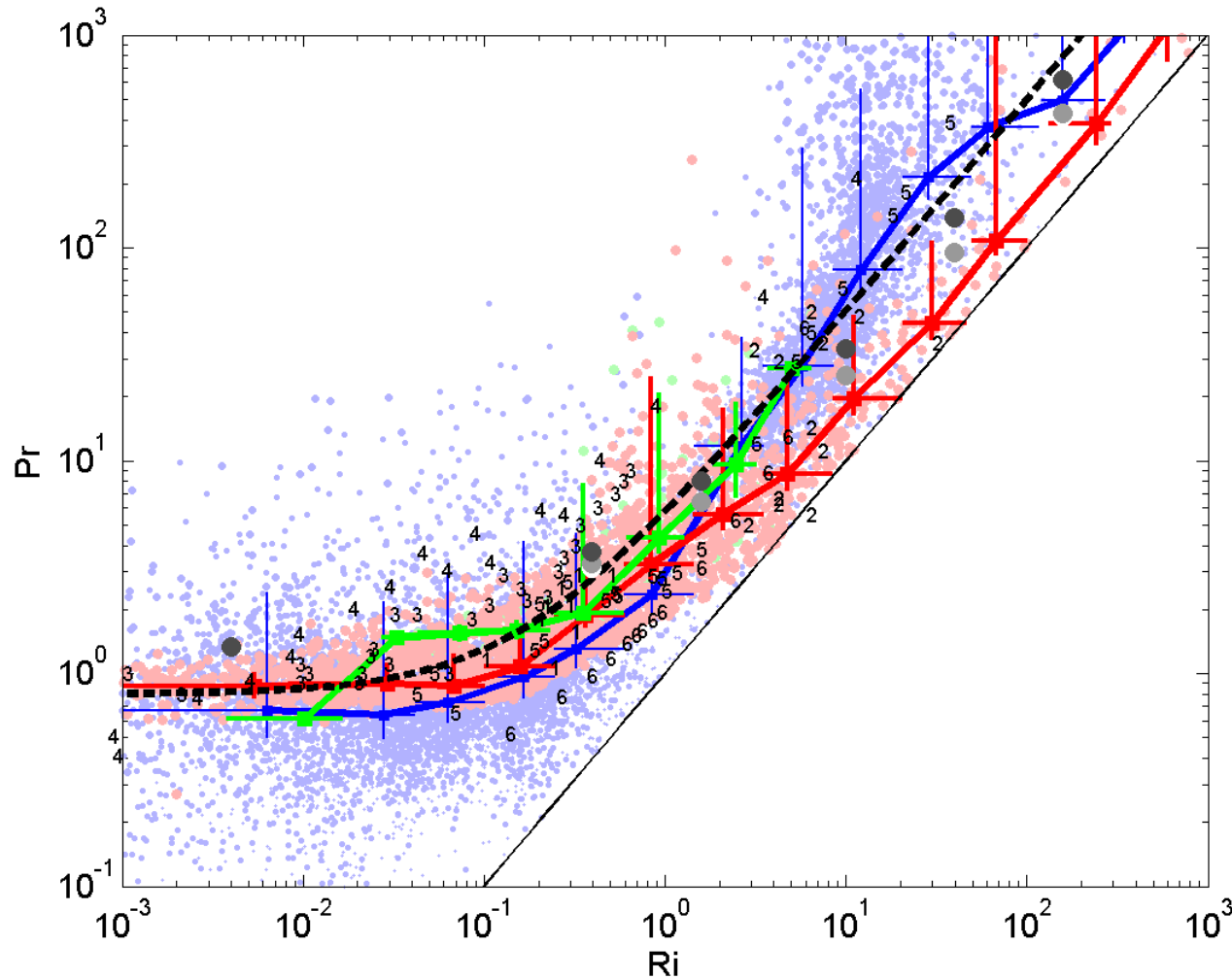
# Ri dependence of turbulent Prandtl number



transition at  $Ri=0.25$



# Ri dependence of the turbulent Prandtl number



Similar dependence  
after recent data

**Blue:** atmosphere  
(SHEBA & CASES-99),

**green:** laboratory,

**red:** LES (NERSC),

**grey:** DNS;

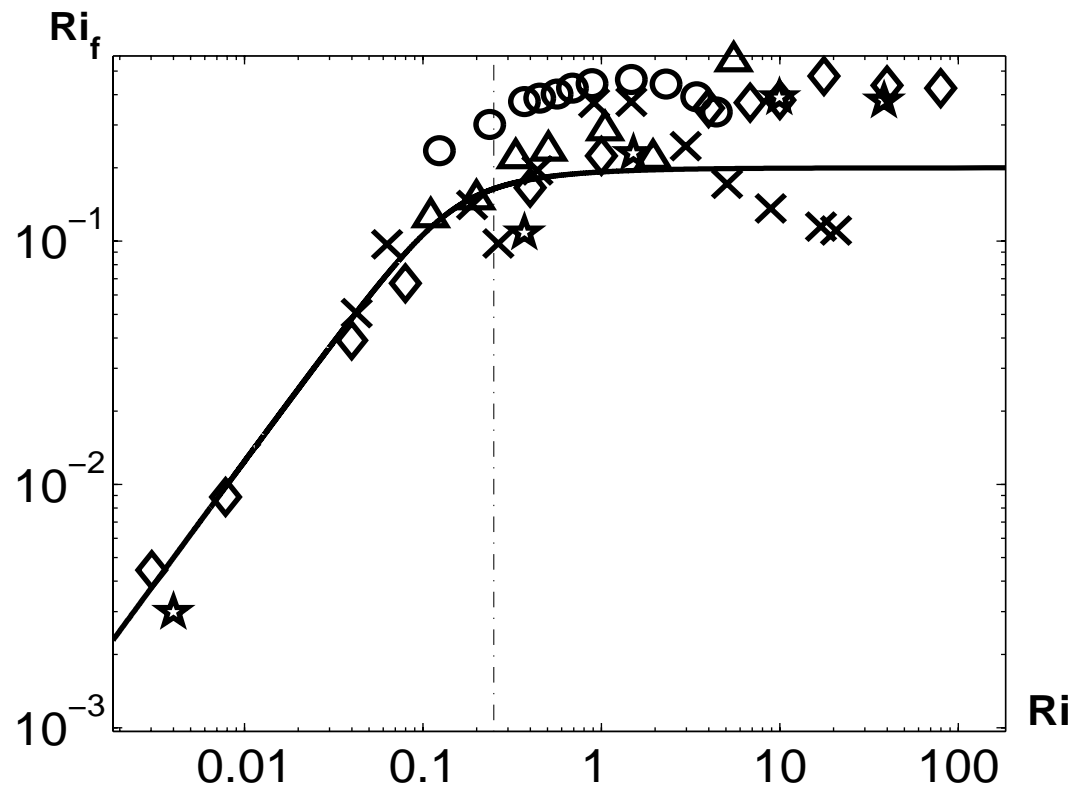
**curve:**  $Pr_T = 0.8 + 5 Ri$

transition at  $Ri=0.25$



# Ri-dependence of the flux Richardson number

**Theory** (for homogeneous turbulence) – solid line; **Laboratory**: circles – Strang & Fernando (2001), diamonds – Ohya (2001), crosses – Rehmann & Koseff (2004); **LES** (Esau, 2006) – triangles; **DNS** (Stretch et al., 2001) – stars



transition at  $Ri=0.25$



# Proposed hierarchy of models

## Budget equations for the basic statistical moments:

Turbulent energies: TKE and TPE (or TTE and either TKE or TPE)

Vertical turbulent fluxes of potential temperature,  $F_\theta$ , and momentum  $\tau_{\{x, y\}}$

## Closure hypotheses:

Kolmogorov's dissipation rates (generalised – for “dissipation” of  $\tau_{\{x, y\}}$ )

The “pressure  $\leftrightarrow$   $\Theta$ -gradient” correlation term in the  $F_\theta$ -equation is  $\sim -\beta\theta^2$   
(spectral analysis and data)  $\rightarrow$  **crucial role of  $\beta\theta^2$**  (counter gradient flux)

Generalised “return-to-isotropy”  $\rightarrow$  redistribution of TKE between  $E_x, E_y, E_z$   
accounting for the stability effect on the  $x \rightarrow z$  energy transfer

Mixing length,  $l=z\Phi(\text{Ri}_f)$       asymptotes  $\rightarrow \Phi \sim (1 - \text{Ri}_f / \text{Ri}_{f,\text{crit}})^n$   
experimental data  $\rightarrow n=4/3$       (next slides)

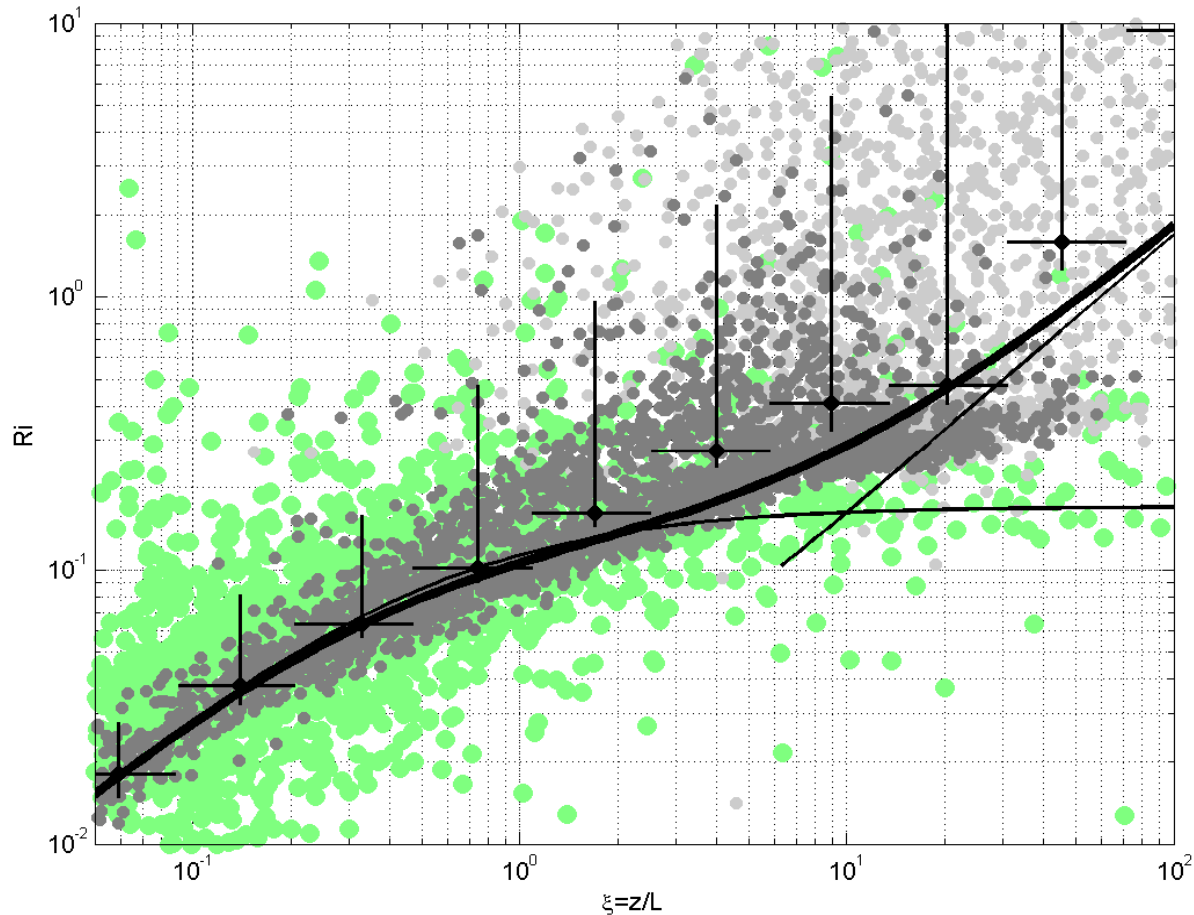


# Part I: Local (algebraic) closure for the steady-state homogeneous regime of turbulence



# Ri versus $z/L$ in stably stratified ABLs

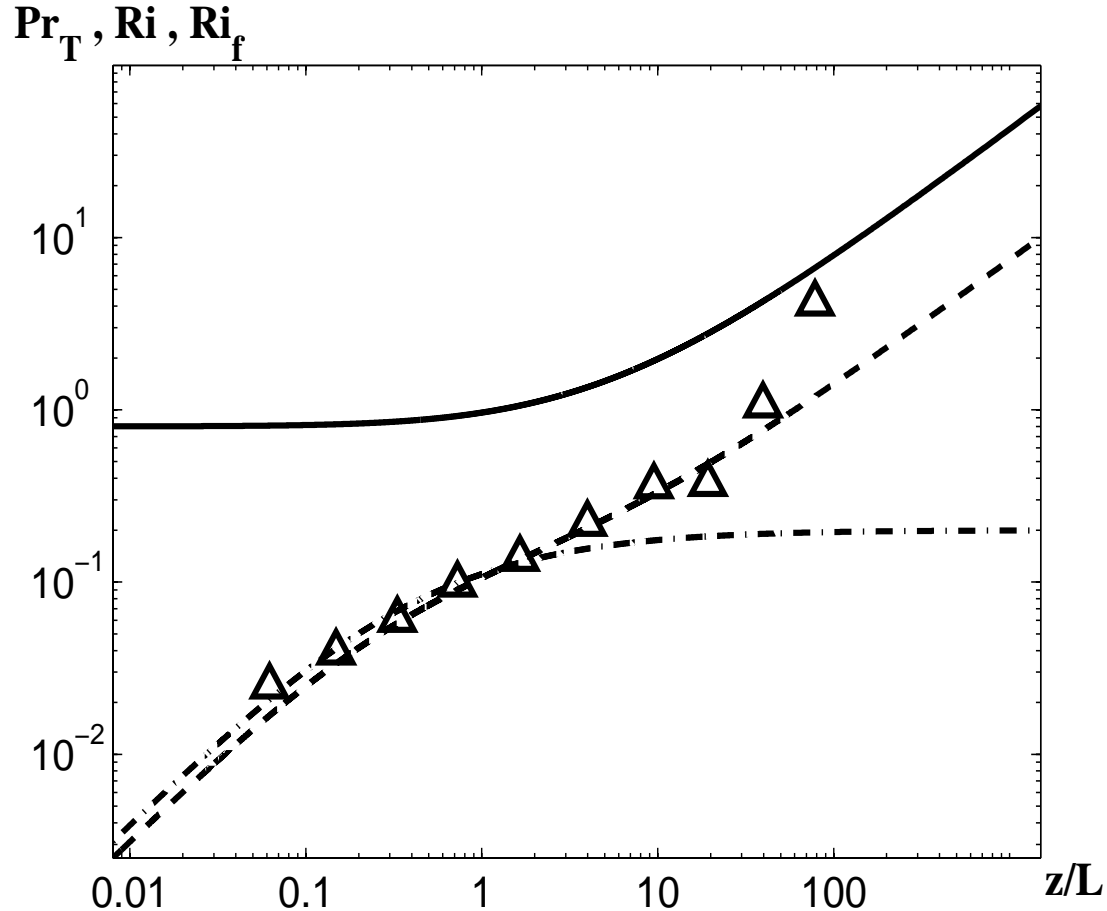
LES: dark grey  $z < h$ , light grey  $z > h$ ; SEBA field data: green; Bin-averaged: black with error bars; Bold line: analytical approximation:  $L = \tau^{3/2} (-\beta F_\theta)^{-1}$





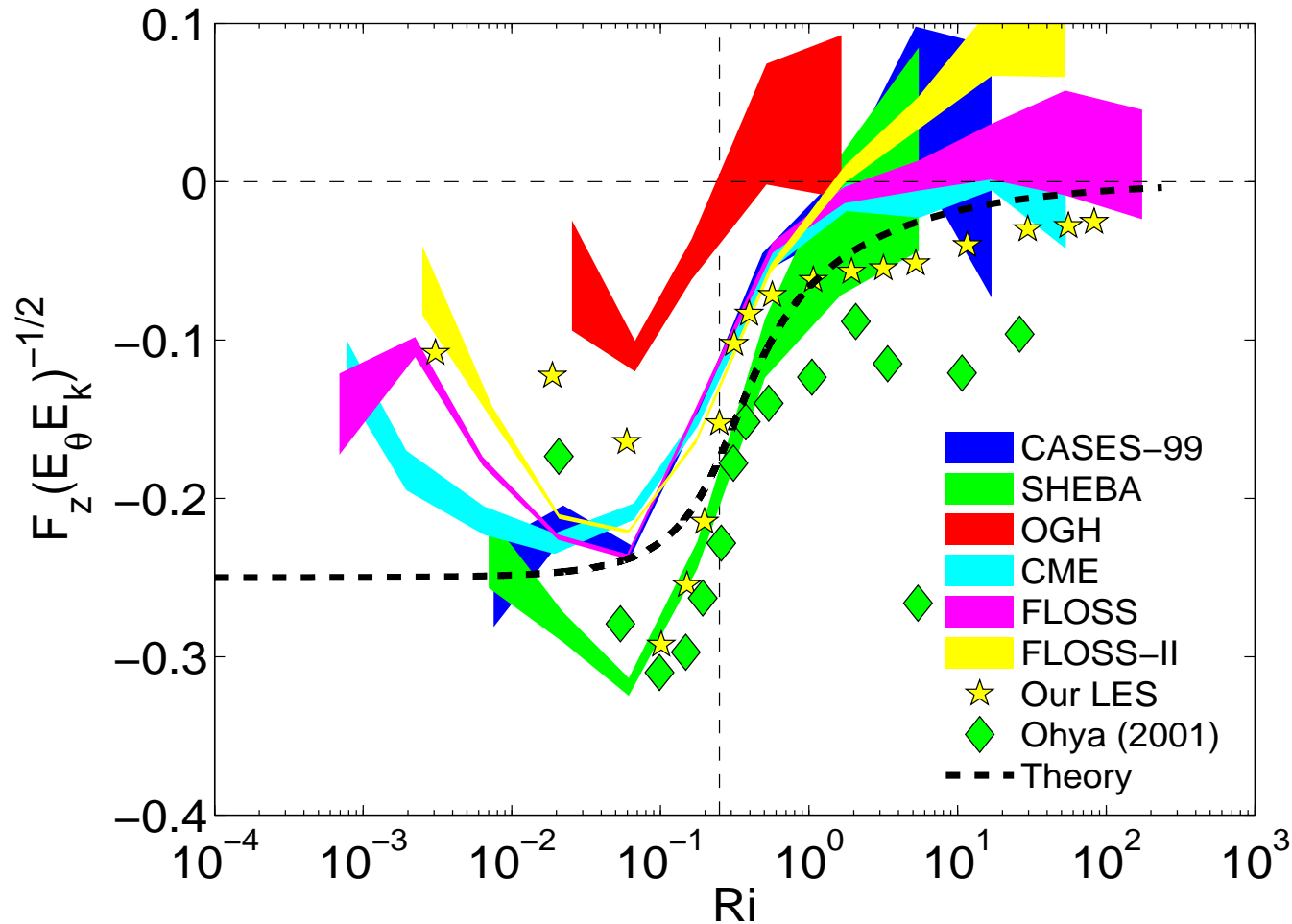
# Ri versus $z/L$ in stably stratified ABLs

triangles: LES; dashed line:  $l/z = (1 - Ri_f/Ri_{f,crit})^{4/3}$ ;  $L$  is MO length



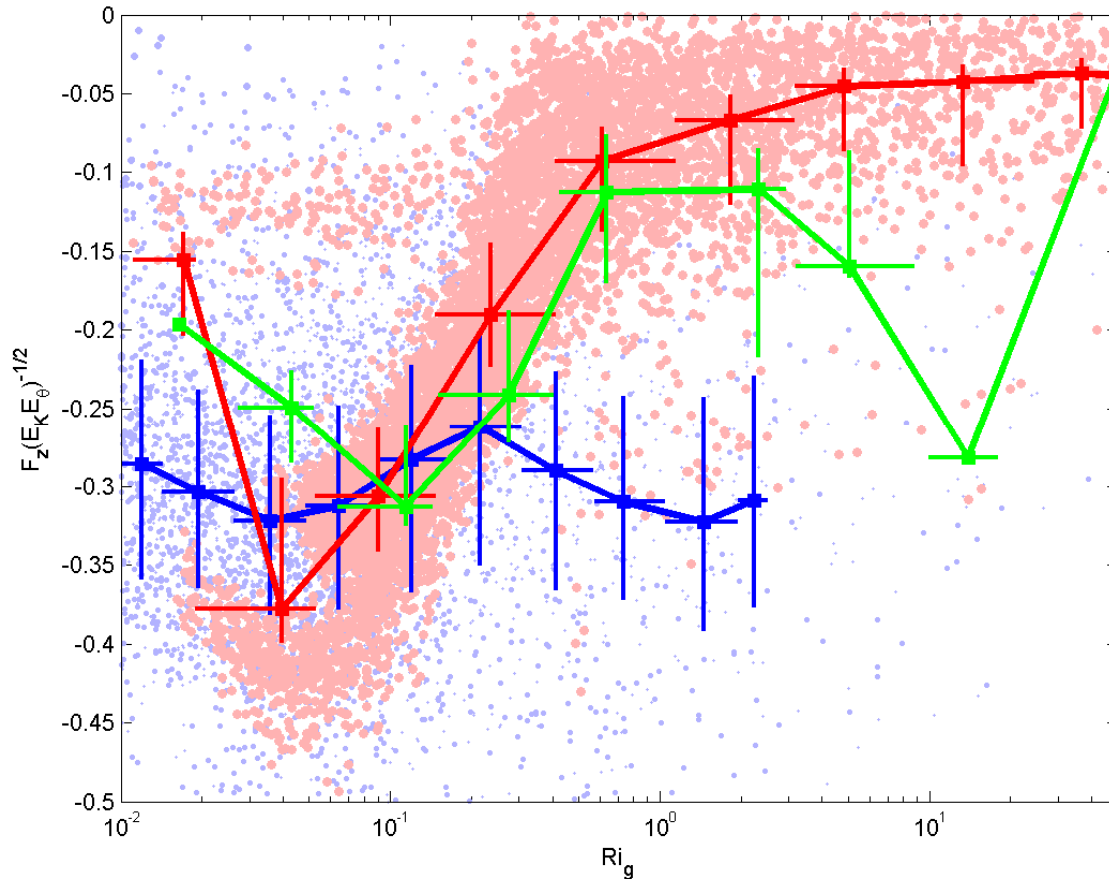
# Normalised turbulent heat flux versus Ri

analysed by Mauritsen: transition at Ri=0.25 (dashed line: theory)



# Normalised turbulent heat flux vs. $Ri$

analysed by Esau: LES (red), lab (green, Ohya, 2001), blue (Sheba)

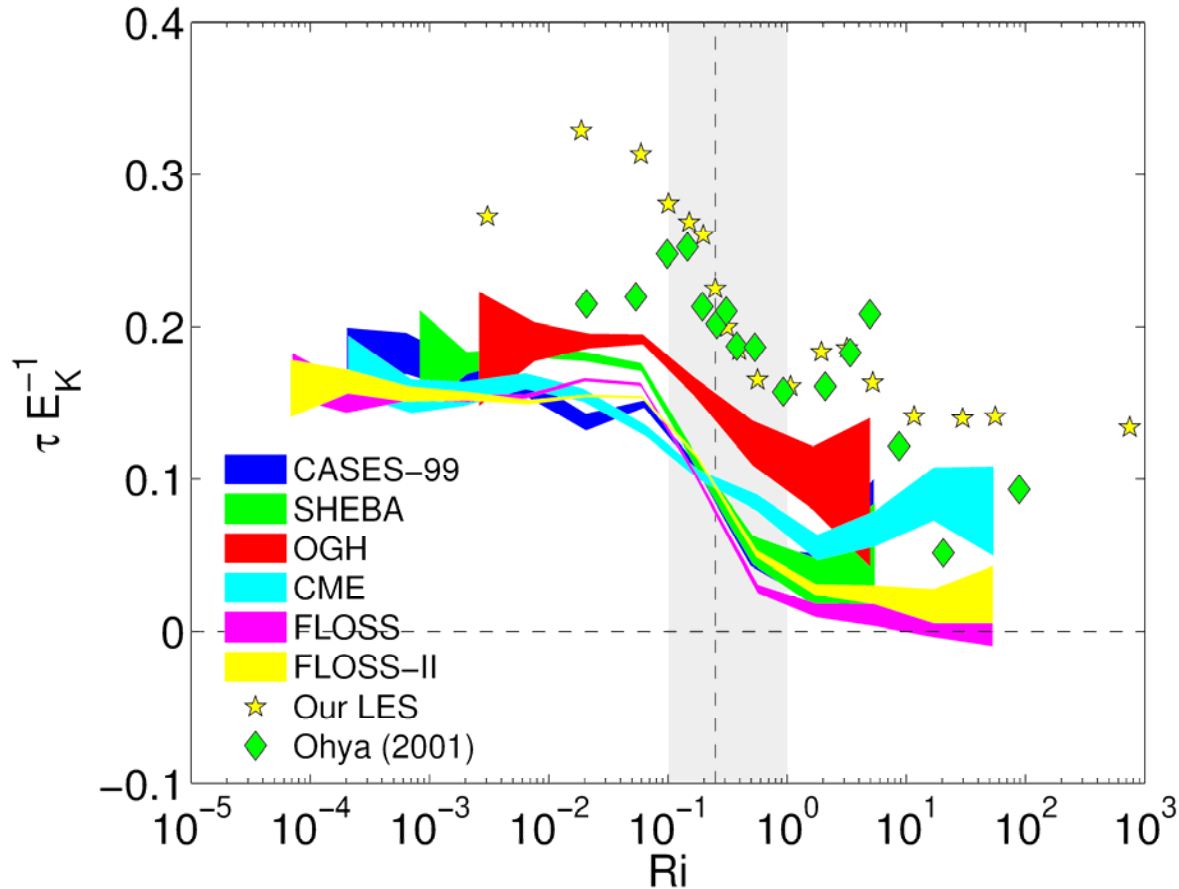


for LES  
transition at  $Ri \sim 0.25$



# Normalised flux of momentum vs. Ri

analysed by Mauritsen: transition at  $Ri=0.25$

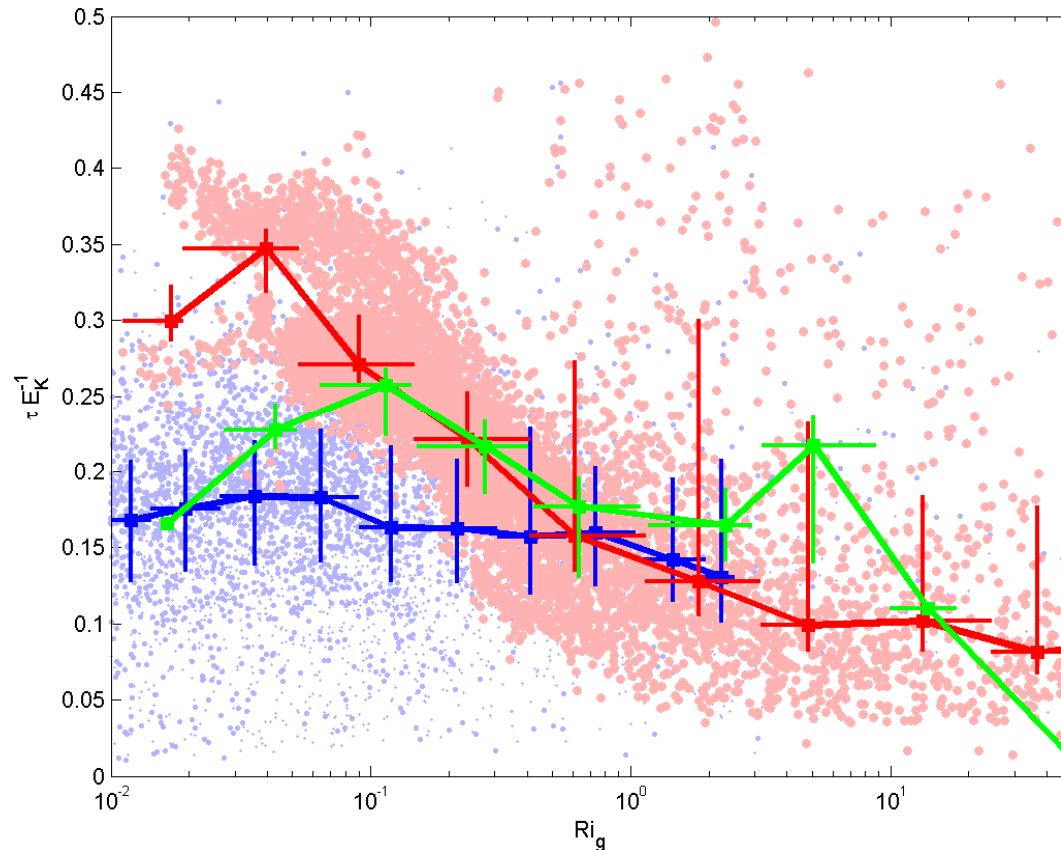


transition at  $Ri=0.25$



# Normalised flux of momentum vs. Ri

analysed by Esau: LES (red), lab (green, Ohya, 2001), blue (Sheba)



for LES  
transition at  $Ri \sim 0.25$



# Sharp transition – between two regimes

Anisotropy and normalised heat flux (also other dimensionless parameters) turn into constants in the two alternative regimes: (i) near-neutral, (ii) very stable, with **sharp transition** in a narrow interval of Ri around  $Ri=0.25$

Coincidence of with **classical hydrodynamic instability threshold**

However, this threshold separates not **turbulent** and **laminar** regimes, as classical concept says, but two different turbulent regimes: **strong, chaotic** ( $Pr_T$  of order unity) and **weak, intermittent** (wave-induced momentum transfer,  $Pr_T$  linearly increases with Ri)



## Part II: Minimal non-local closure (future work)

Five budget equations (for TKE, TPE and turbulent fluxes) and one relaxation equation for the mixing length accounting for

- temporal evolution,
- transport by mean wind,
- vertical turbulent transport (at 1<sup>st</sup> stage through turbulent diffusion)



# Conclusions: physics

- The TKE equation do not fully characterise energy transformations. Equally important are the **TPE or TTE equations** (TTE = TKE+TPE)
- TTE **(invariant) conserves along motion** in the absence of shear and dissipation. TPE-equation allows developing advanced turbulence closures
- It proves that velocity shear generates turbulence in any stratification, and **disapproves the critical Ri concept**. Field, lab, LES and DNS data show maintaining of turbulence at  $Ri > 100$  – two orders larger than  $Ri_c = 0.25$
- Ri-dependences disclose **sharp transition** at  $Ri \sim 0.25$  (linear instability limit), which separates “strong, chaotic” and “weak, intermittent” regimes
- The model is extended to **turbulent diffusion of passive scalars**





# Conclusions: modelling

L'vov, Procaccia and Rudenko, 2007: Energy conservation and second-order statistics in SBLs. Submitted to *BLM* (Weizmann Institute of Science, Israel) → detailed analysis of the momentum-flux budget

Mauritsen et al., 2006: TTE closure model for SBLs. *JAS* in press (Stockholm University) → a model employing TTE and TPE budgets and empirical dependencies of  $\tau/E_K$  and  $F_z(E_K E_\theta)^{-1/2}$ , consistent with Z et al. non-local scaling and experimental & LES data for mean profiles and ABL height, including conventionally neutral and long lived stable ABLs



**Application to wind energy**

Revised high resolution modelling of the wind field in the lower 200-300 m to refine the wind energy potential, especially for coastal zones



# Future: unsolved problems for convective flows

**Decomposition: mean, turbulent, semi-organised (cells / rolls)**

**Nature and theory of rolls: aspect ratio? energetics?**

**Parameterization**

**Application to turbulent transport: advanced closure**

**Application to wind energy (contribution from rolls)**



**END**

