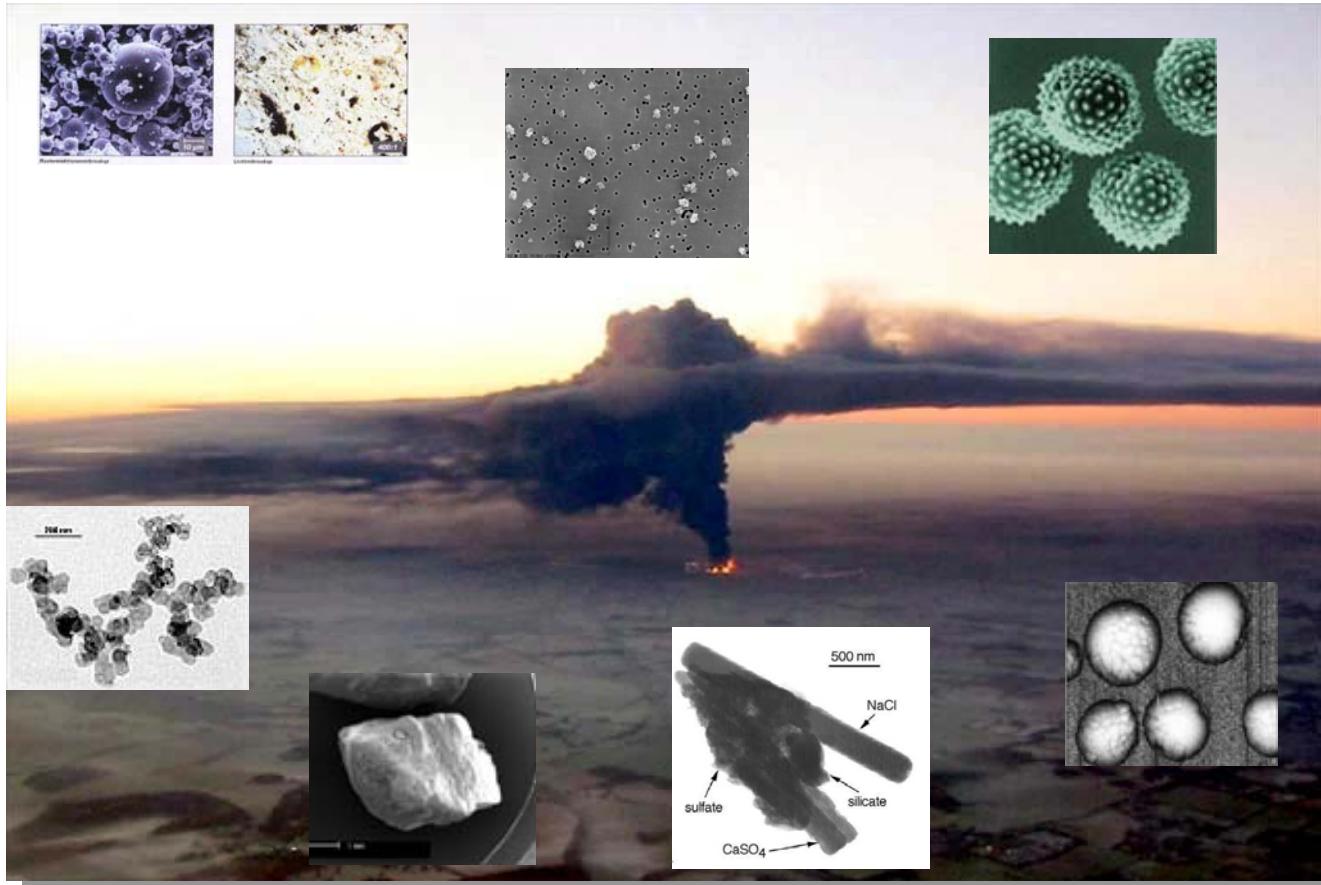


COSMO LM-ART

Aerosols and Reactive Trace Gases within LM

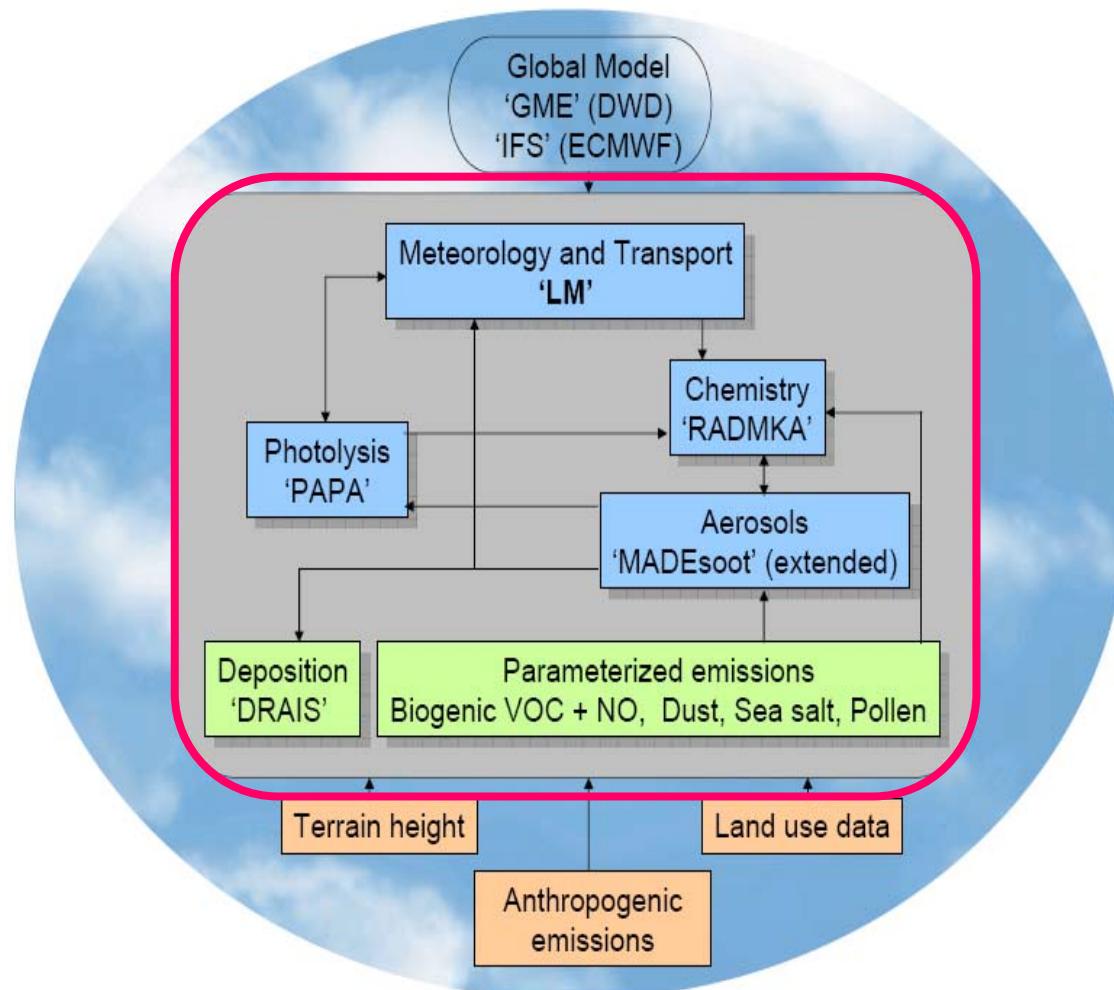


Dominique Bäumer
Max Bangert
Kristina Lundgren
Rayk Rinke
Tanja Stanelle
Bernhard Vogel
Heike Vogel

Questions

- Interaction aerosol and regional climate
(e.g. secondary circulations, mineral dust and monsoon)
- Interaction aerosol and gas phase
- Parameterization of both processes for global scale models
- Modelling of the visibility and AOD
- Dispersion of pollen

COSMO LM – ART (ART = Aerosols and Reactive Trace Gases)



Concept:

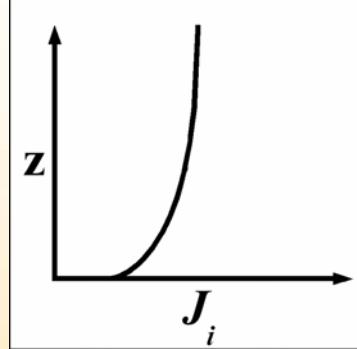
LM-ART is **online** coupled.

Identical methods are applied for all scalars as temperature, humidity, and concentrations of gases and aerosols to calculate the transport processes.

It has a **modular** structure.

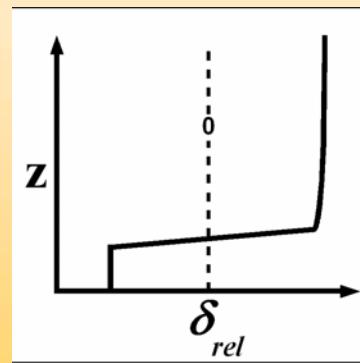
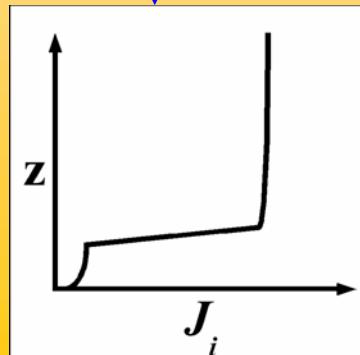
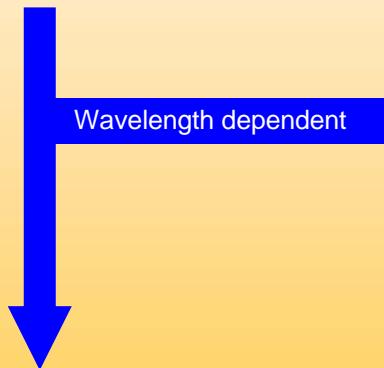
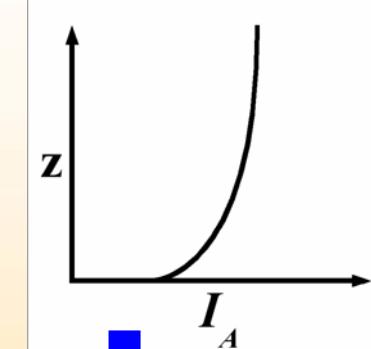
Therefore **LM-ART** can easily be used in the **forecast mode**.

Parameterizations of the Photolysis Frequencies

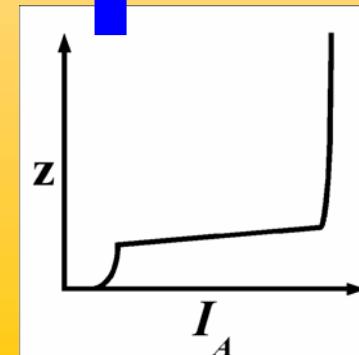
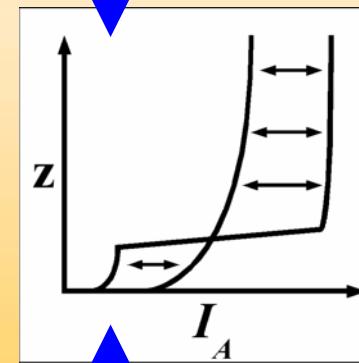


Pre - calculated profiles
For identical conditions with
STAR and GRAALS (LM)

Cloud free, no aerosols, sea level



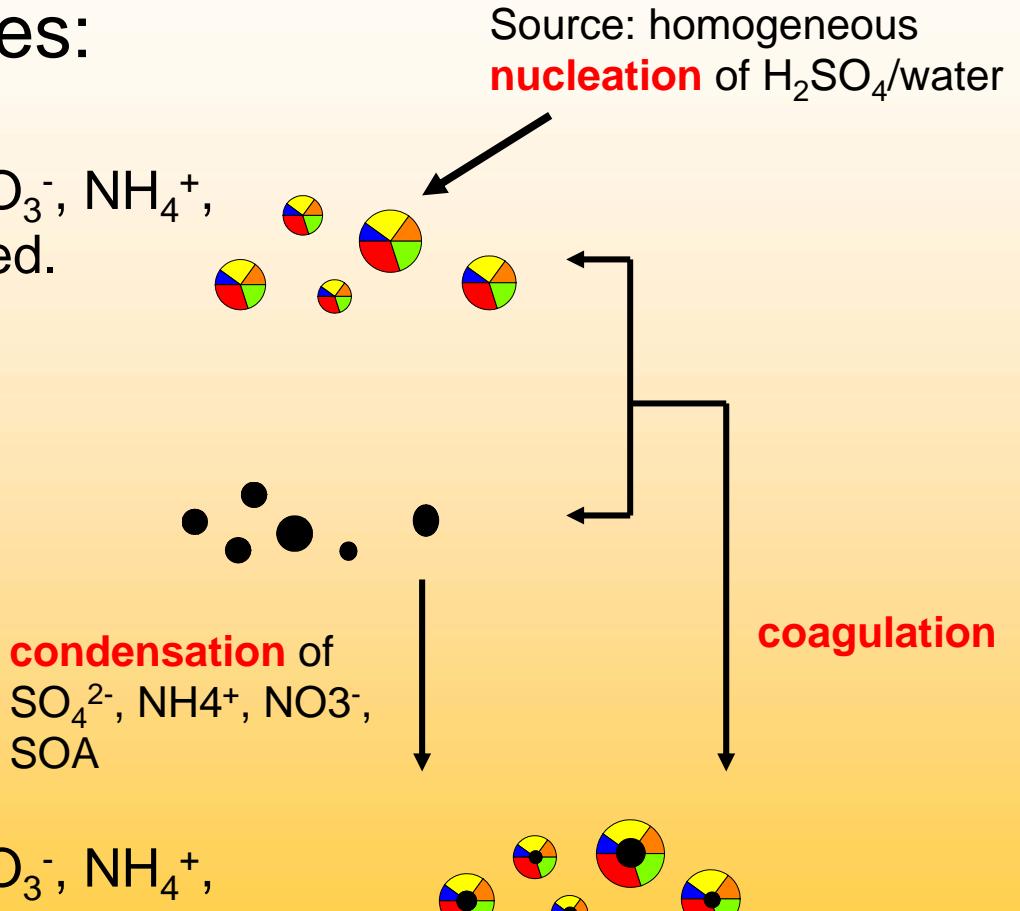
In LM-ART
Online coupled
Accounts for
clouds, aerosol,
terrain height, albedo



Treatment of the Aerosol Particles

Interaction of five modes:

- **Two modes** for SO_4^{2-} , NO_3^- , NH_4^+ , H_2O , SOA, internally mixed.
- **One mode** for pure soot.
- **Two modes** for SO_4^{2-} , NO_3^- , NH_4^+ , H_2O , SOA, and soot internally mixed.



Three modes for **mineral dust** particles + three modes for **sea salt** particles + **pollen**

Optical Properties of the Aerosols

Refractive index of aerosols

Mie Calculations

Single scattering albedo (ω),
specific extinction coefficient (b),
asymmetry parameter (g)

New Routine in LM-ART:
Computation of ω , b, g for
prevailing aerosol concentration

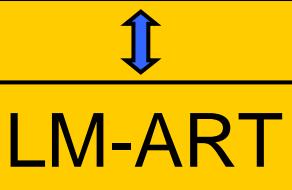
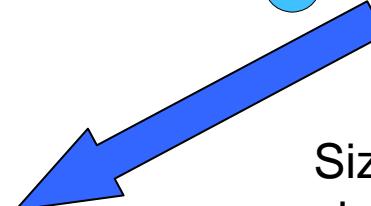
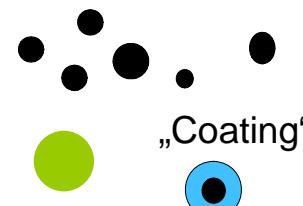
ω , b, g

Modified radiation in LM:
Substitution of climatological optical
properties based on current aerosol
concentrations

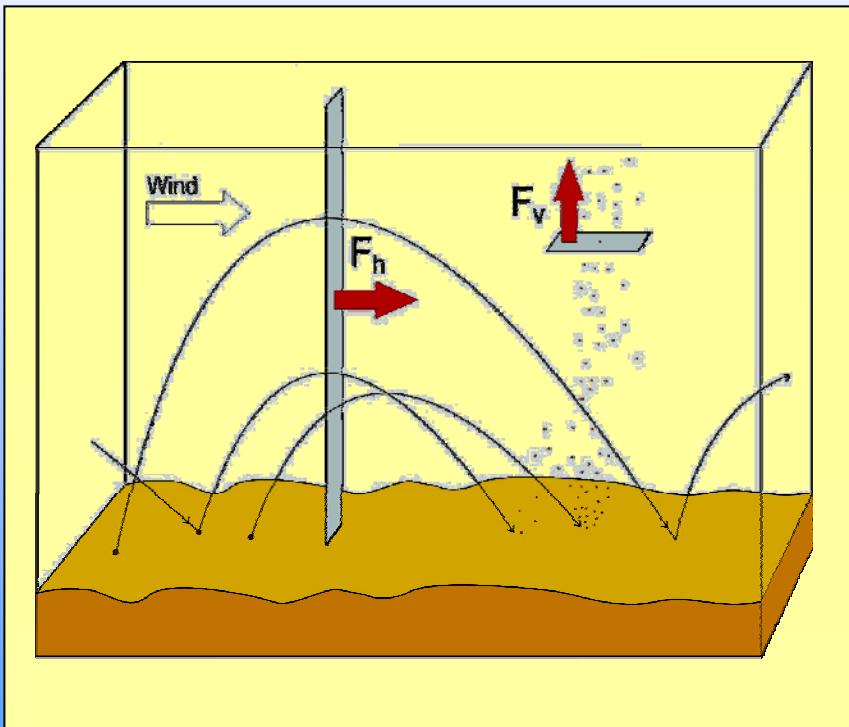
Transport, Sedimentation,
Deposition

LM-ART

Size distribution,
chemical composition
of each mode



Parameterization of the Dust Emissions

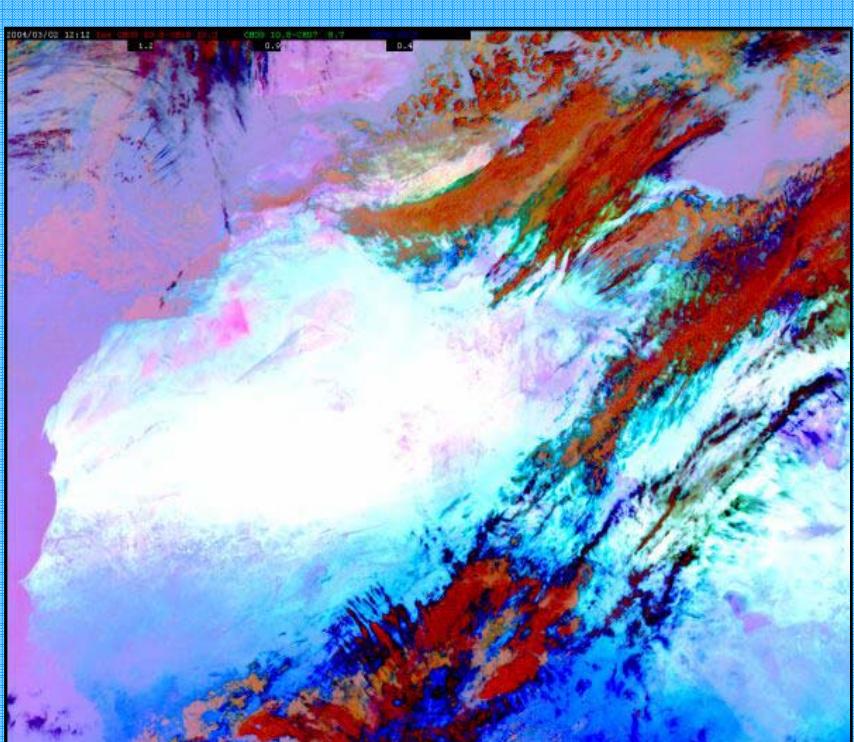


Parameterization of the horizontal and vertical saltation and emission flux (Vogel et al, 2006)

3 different Modes ($d = 1.5, 6.7, 14.2 \mu\text{m}$)

Log normal distributions

Case Study: Mineral dust over West Africa in March 2004



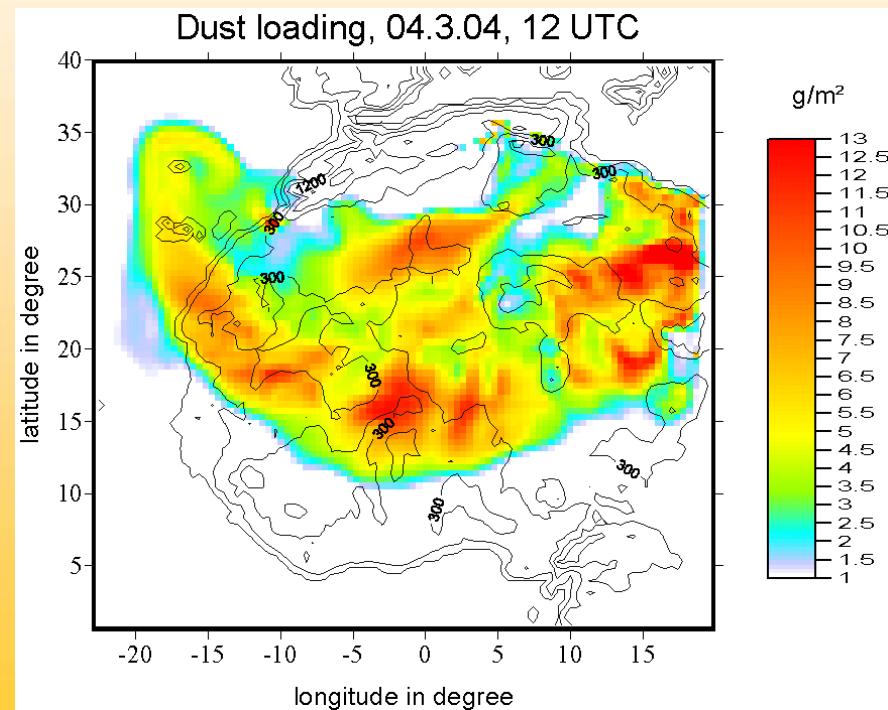
Meteosat-8 Image, 2. – 3. März 2004

Prevailing Situation:

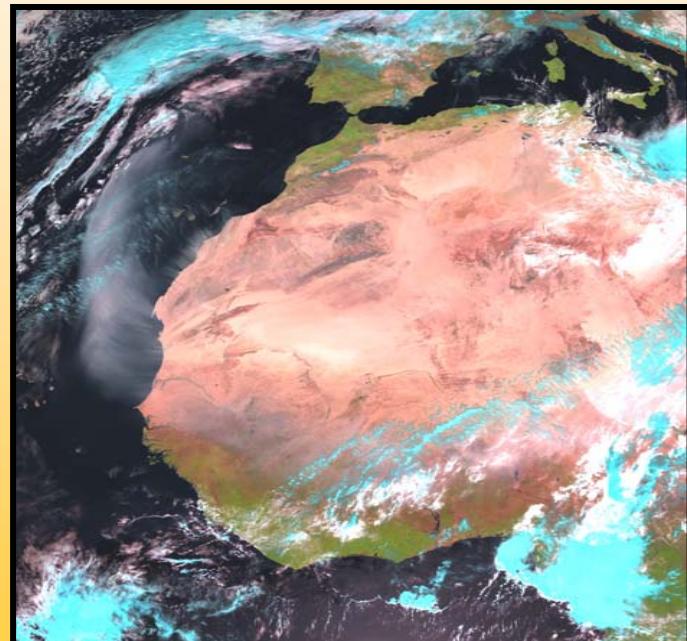
- Unusual
- Low temperatures and high wind speeds in the Sahara
- Heavy Precipitation in Libya
- After the mineral dust event ITC was shifted southwards

(Knippertz and Fink, 2006)

Case Study: Mineral dust over West Africa in March 2004



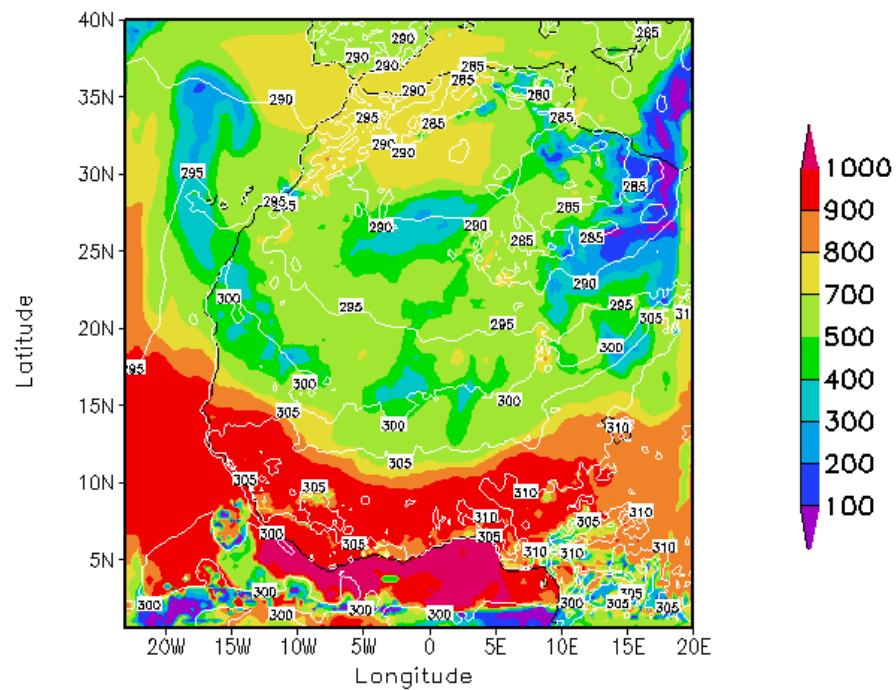
Modis, 04.03.2004, 12 UTC



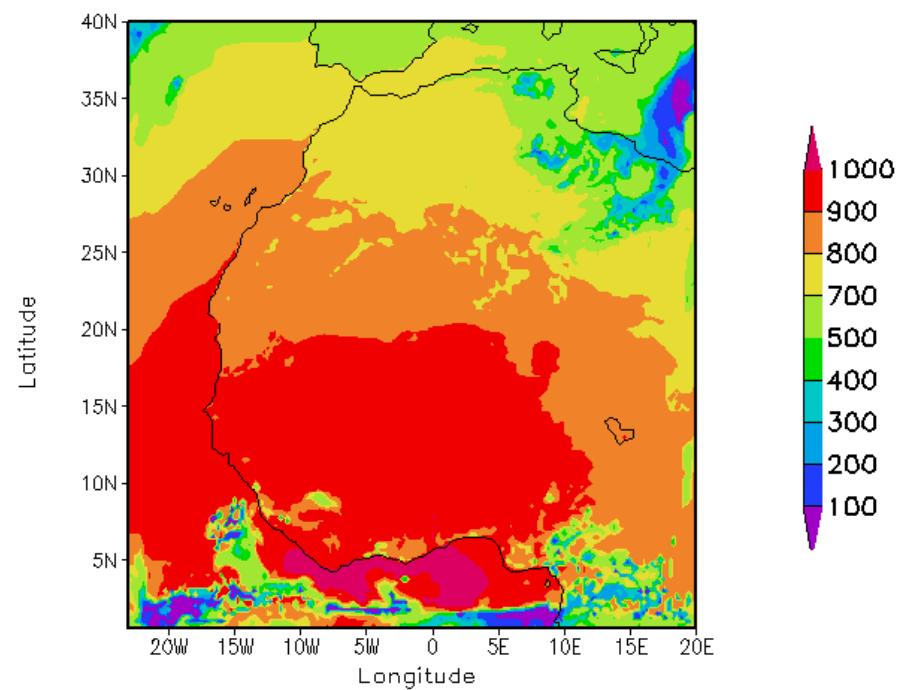
Interaction of Radiation and Mineral Dust

Shortwave radiation balance

with interaction



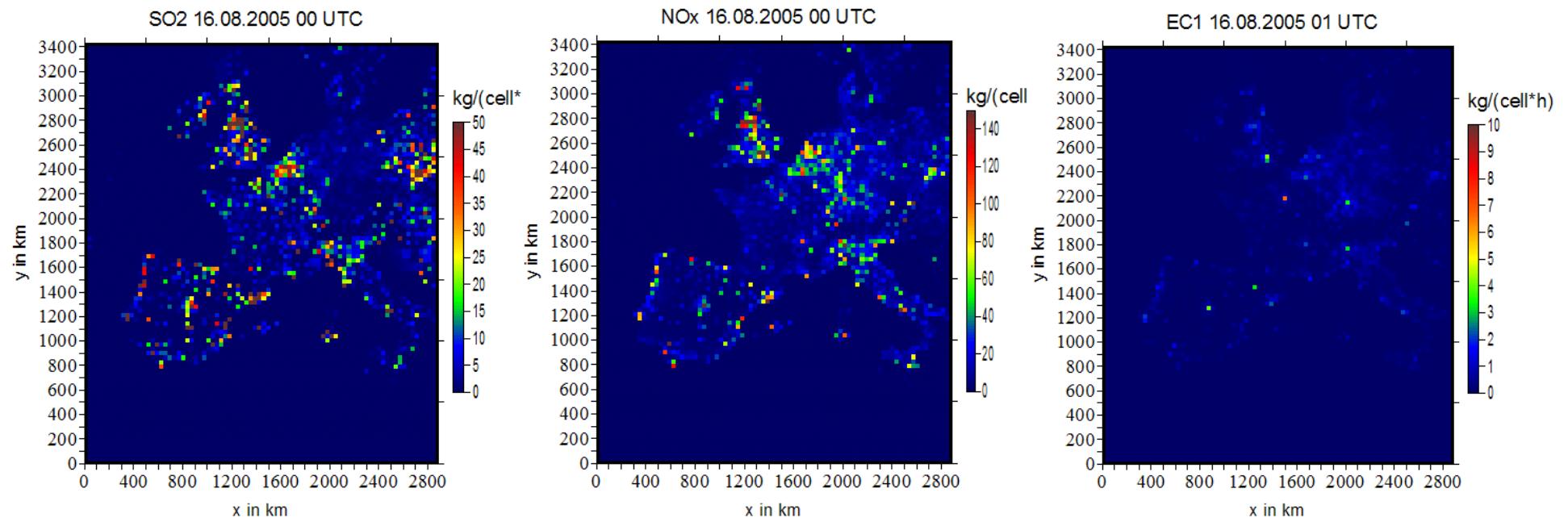
without interaction



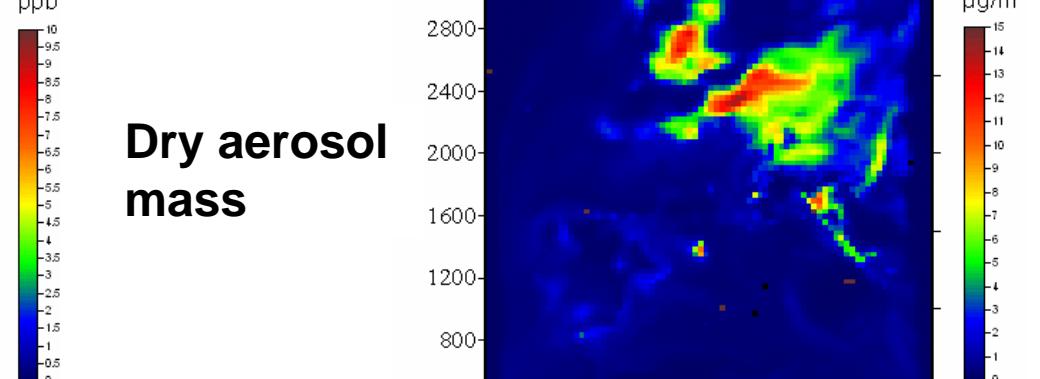
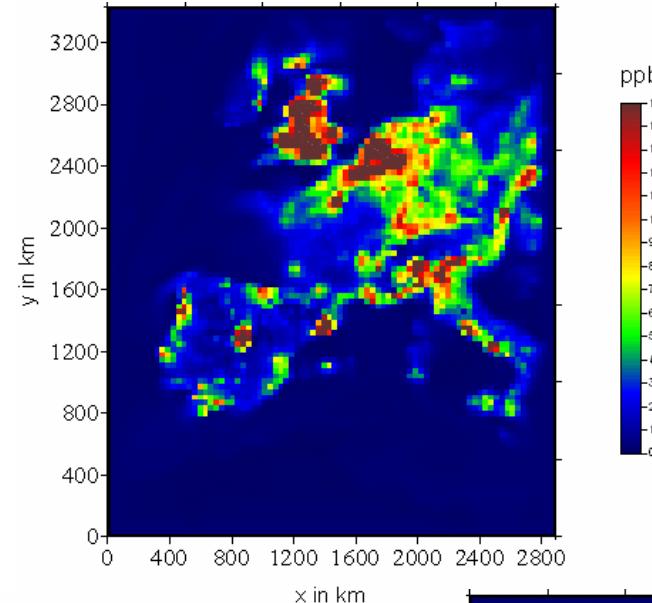
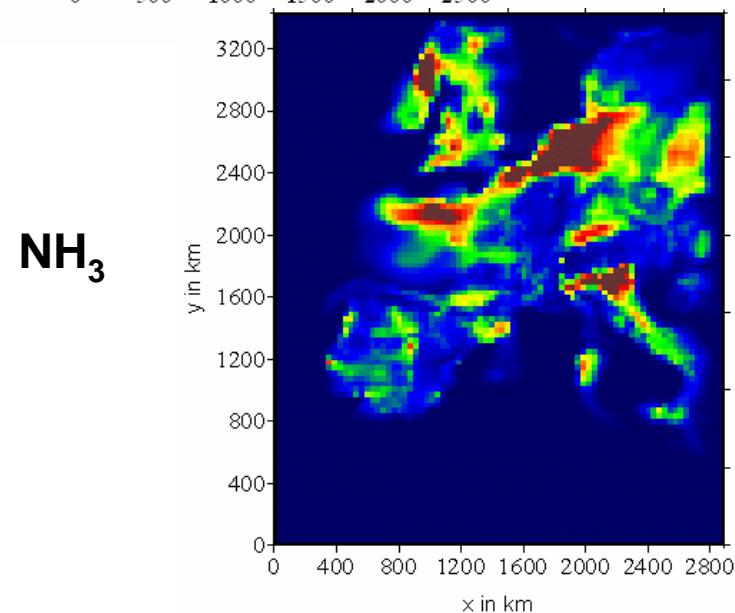
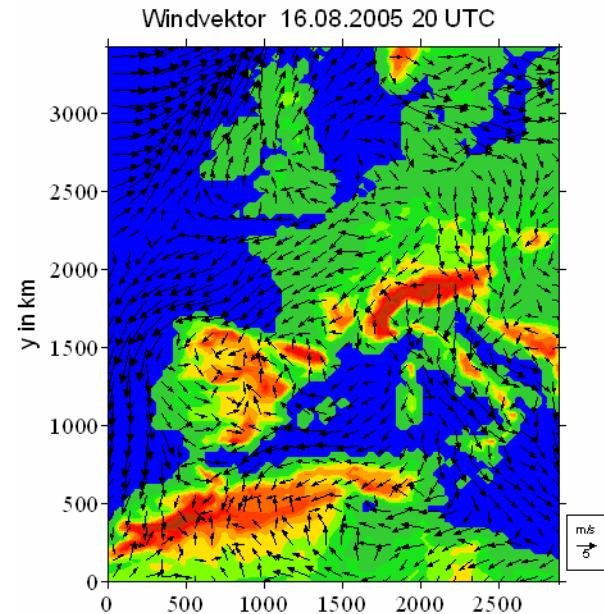
Emissions

(Pegger et al., IER, Uni Stuttgart)

- Discrimination of point and area sources.
- Temporal resolution 1h, spatial resolution 7km x 7km
- 20 gas phase species (5 inorganic species, 15 VOC's)
- Primary particle emissions as EC and PM (1, 2.5, 10 μm)
- Biogenic emissions calculated online



Trace Gases and Aerosols over Europe



Anthropogenic Aerosol Radiation Interaction

Simulation period:

16.08.05 - 22.08.05

Simulation domain:

Southwest Germany + adjacent areas

Horizontal Resolution 7km x 7km

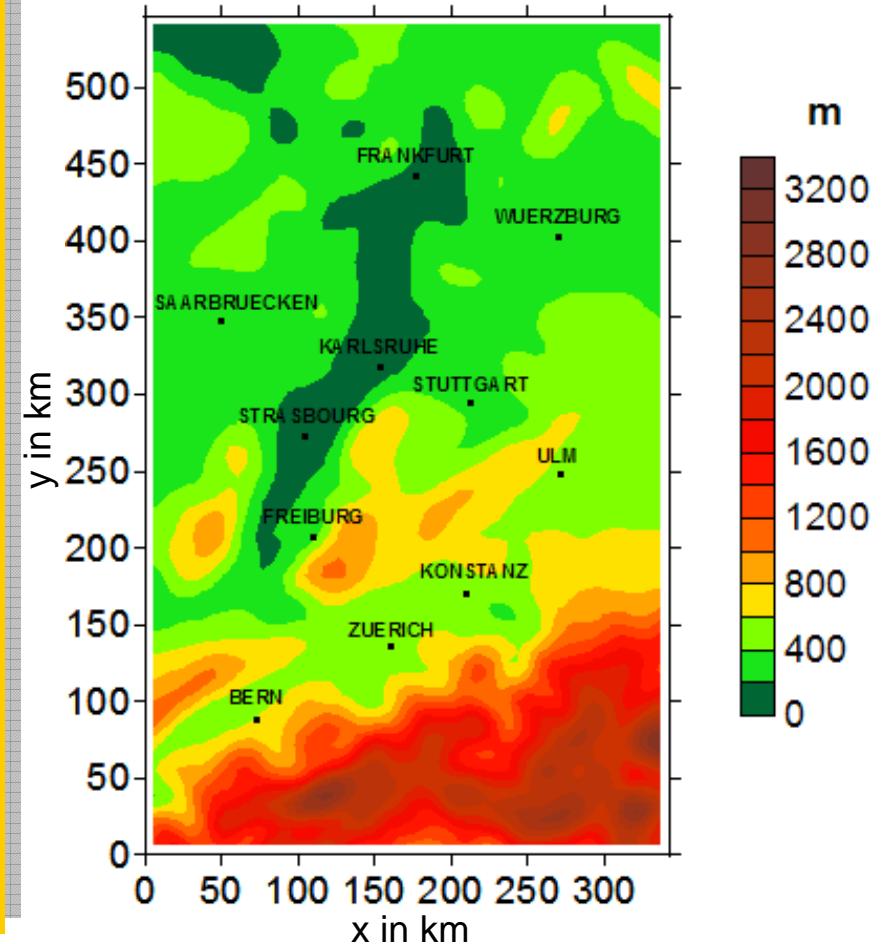
Input data:

- Meteorology: GME Reanalysis (DWD)
- Emission data (IER, Stuttgart)
- Land use (JRC-IES, Ispra)

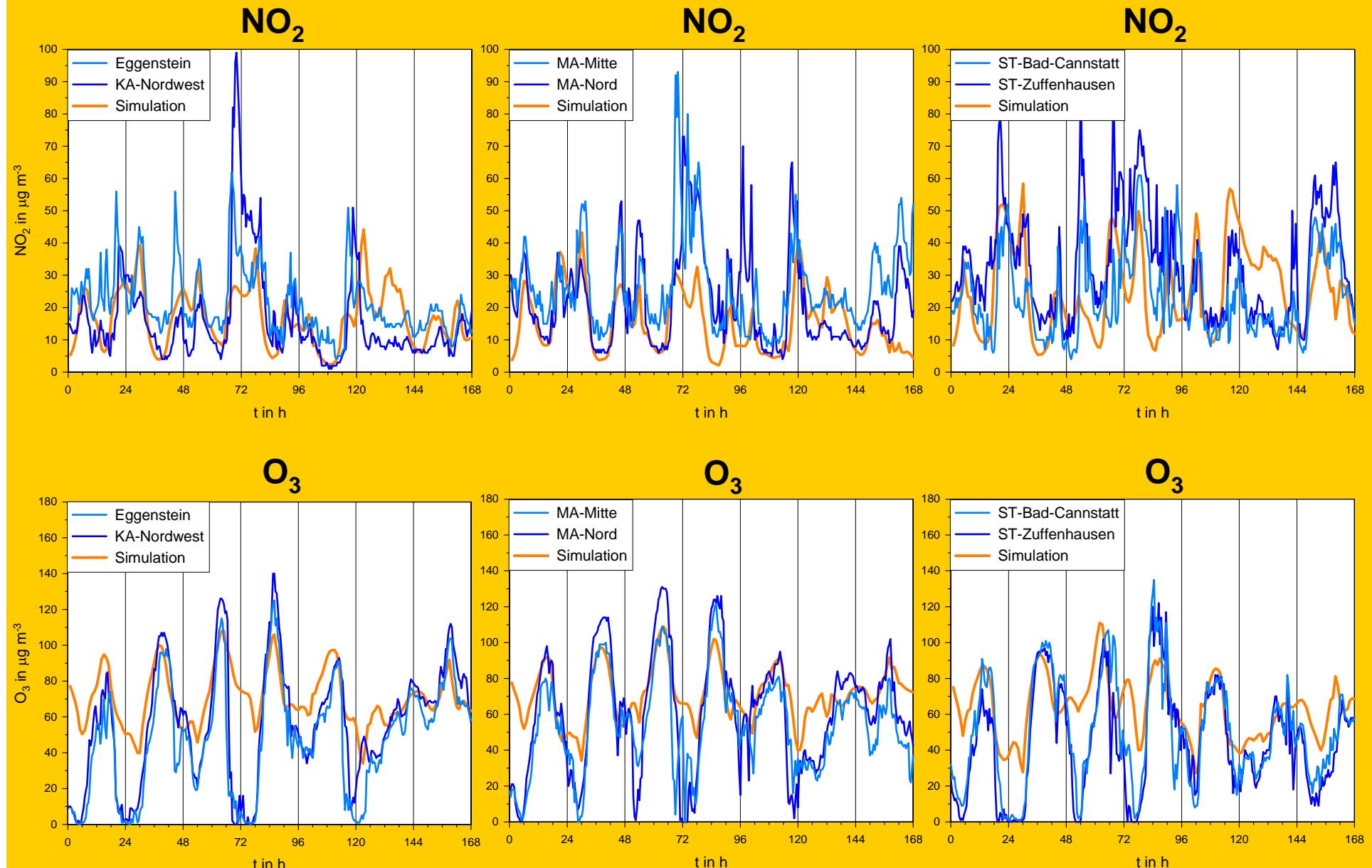
Run A: with aerosol radiation interaction.

Run B: without aerosol radiation
interaction

Model domain



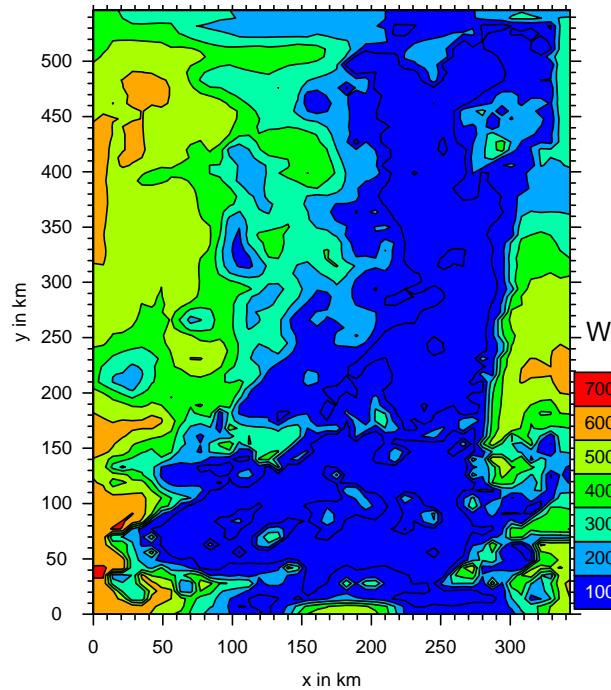
Comparison with Measurements (16.08. – 22.08.2005)



Direct Aerosol Effect on Radiation and Temperature

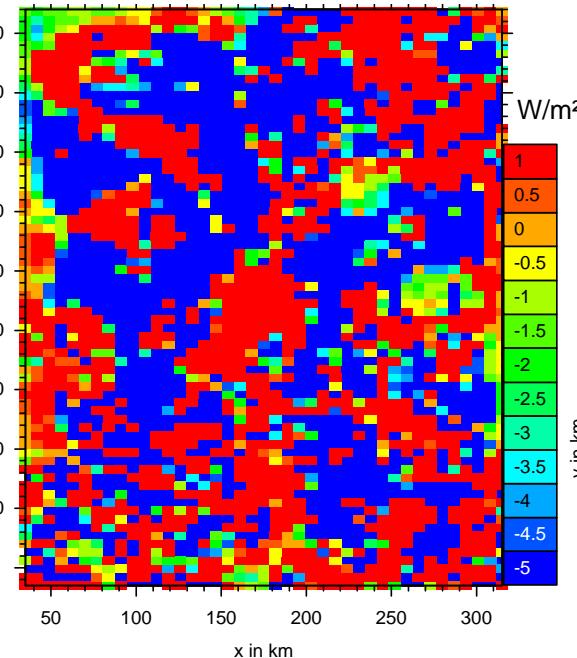
(20.08.2005 12 UTC)

Shortwave radiation balance



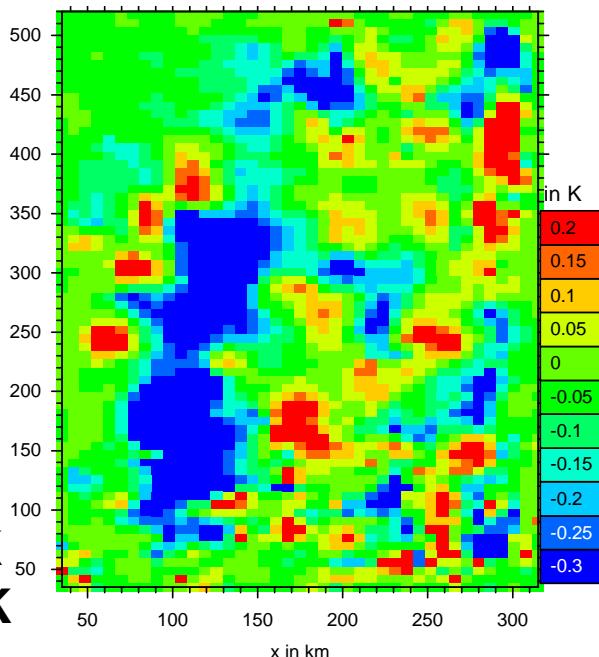
Min: -342.6 Wm⁻²
Max: +349.5 Wm⁻²

Difference of the shortwave
radiation balance
(Run A – Run B)



Min: -1.23 K
Max: +0.85 K

ΔT (Run A – Run B)



Parameterization of the Source Function of Pollen Emission

$$F_p = c_p \cdot \frac{q_p}{LAI \cdot h} \cdot u_* \cdot K_e$$

F_p : flux of pollen grains

c_p : plant specific factor

q_p : total pollen number of a season per m⁻²

LAI: leaf area index

h : canopy height

u_* : friction velocity

K_e : meteorological correction factor

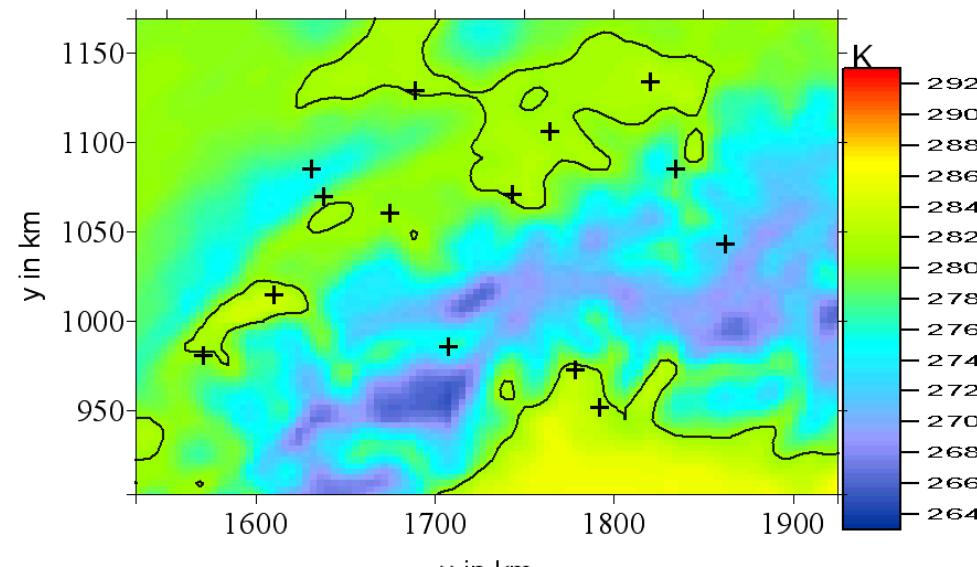
Birch Pollen Episode in Switzerland

In collaboration with A. Pauling, Meteoswiss

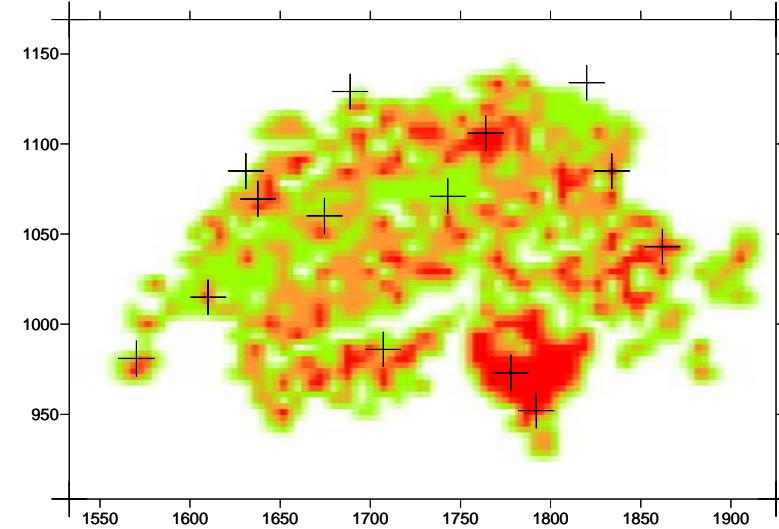


Schweizerische Eidgenossenschaft
Confédération suisse
Confederazione Svizzera

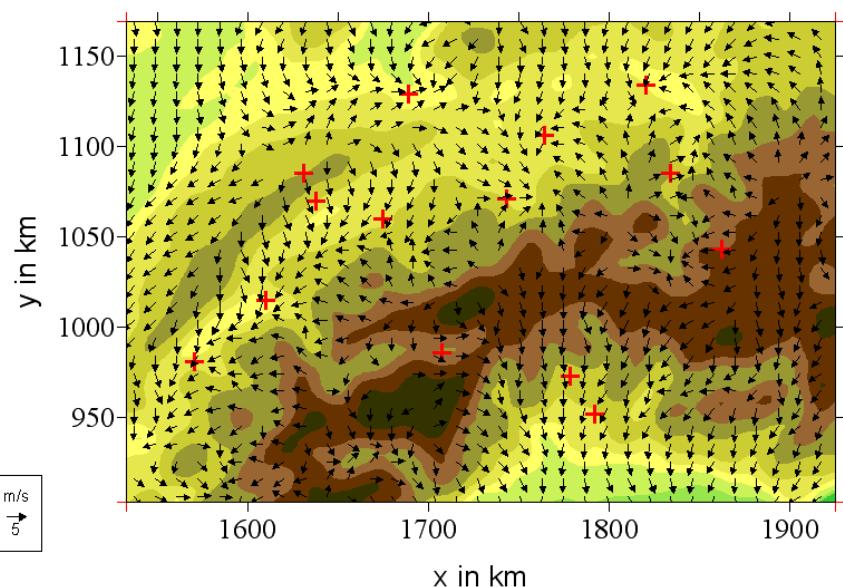
Temperature, 19.04.2006, 01 UTC



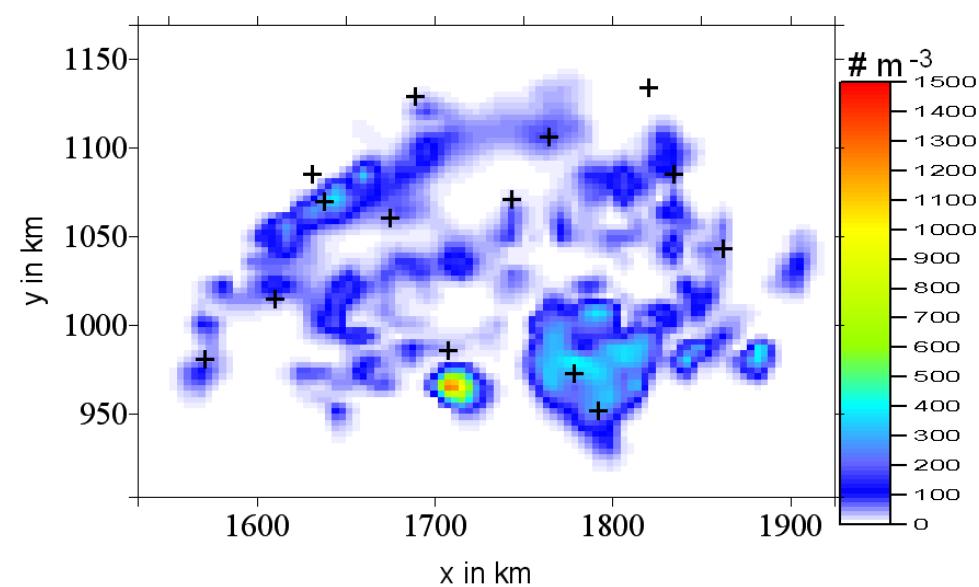
Birches



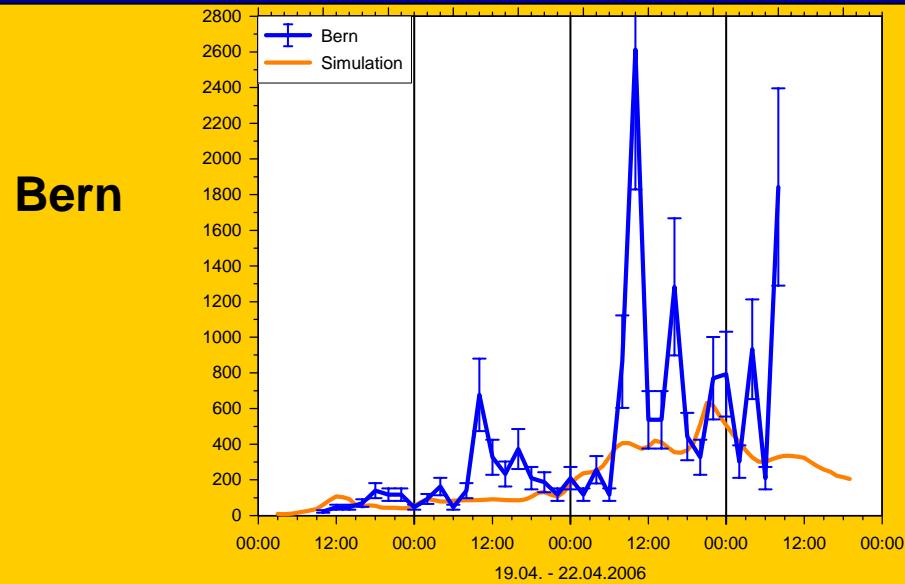
Wind , 19.04.2006, 01 UTC



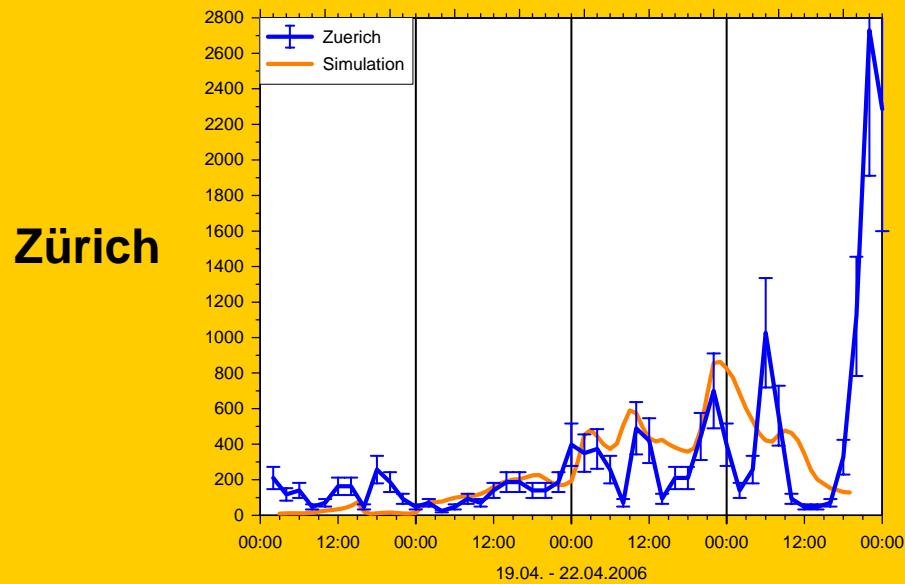
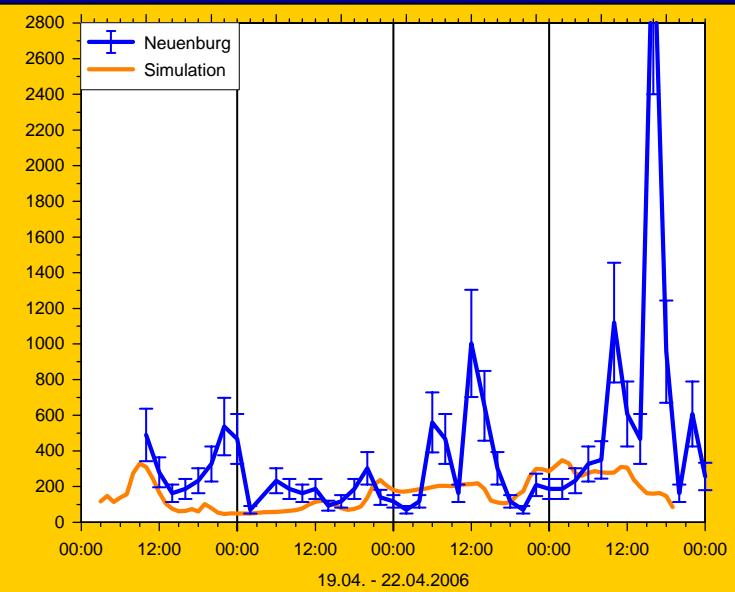
Pollen, 19.04.2006, 01 UTC



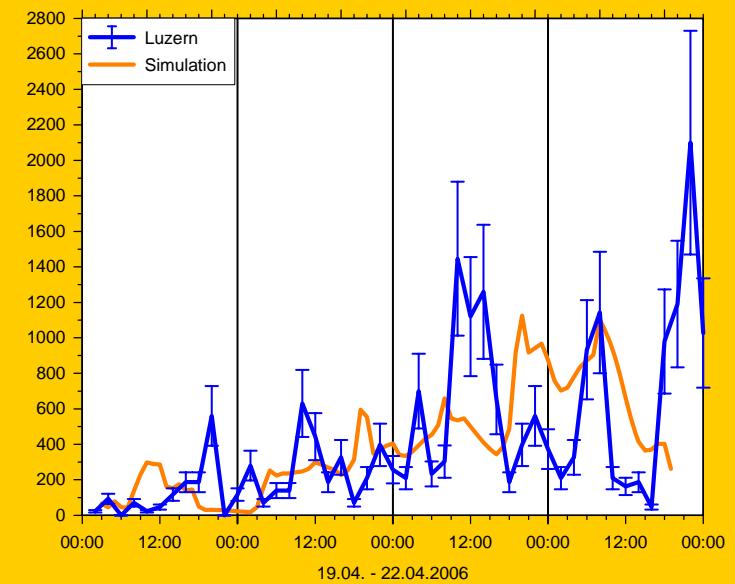
Comparison with Measurements



Neuenburg



Luzern



Summary

- **LM-ART contains a variety of natural and anthropogenic aerosol particles.**
This includes an explicit treatment of the soot ageing. The treatment of their impact on the atmospheric radiation allows the quantification of feedback mechanisms.
- **Dust-radiation-interaction in LM-ART:**
Strong influence on radiation balance and consequently on the dynamics
- **Anthropogenic aerosol-radiation-interaction in LM-ART:**
Very low aerosol loads cause changes in shortwave radiation balance at the ground of reasonable sign and order.
Surprisingly strong effect on cloud pattern (semi-indirect effect) was found.
- Pollen dispersion is also included. This allows an operational pollen forecast with **LM-ART** in the near future.

COSMO LM – ART (ART=Aerosols and Reactive Trace Gases)

Gas phase chemistry: **58 transported variables**,
Aerosol: **77 transported variables**

Nesting option is currently realized in collaboration with U. Schättler

Future Developents:

Interaction of aerosols and clouds

Wet phase chemistry

Aerosols and Climate Processes

Aerosols have an impact on human health,

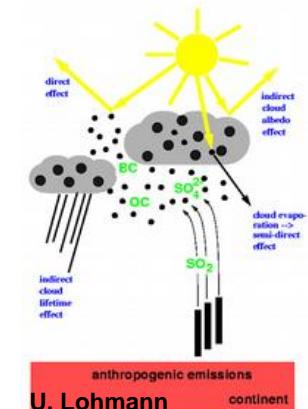
BUT they have also an impact on climate (and weather).

by:

modifying the atmospheric radiation (direct effect),

modifying cloud formation (indirect effect),

and mixtures of both.



AND:

They are changing the chemical composition of the atmosphere.

Emissions of Sea Salt Particles

◆ Mårtensson *et al.* [2003]: $\frac{dF_0}{d \log D_p} = \Phi \cdot 3.84 \cdot 10^{-6} \cdot U_{10}^{3.41}$, $\Phi = \Phi(T_w, D_p)$

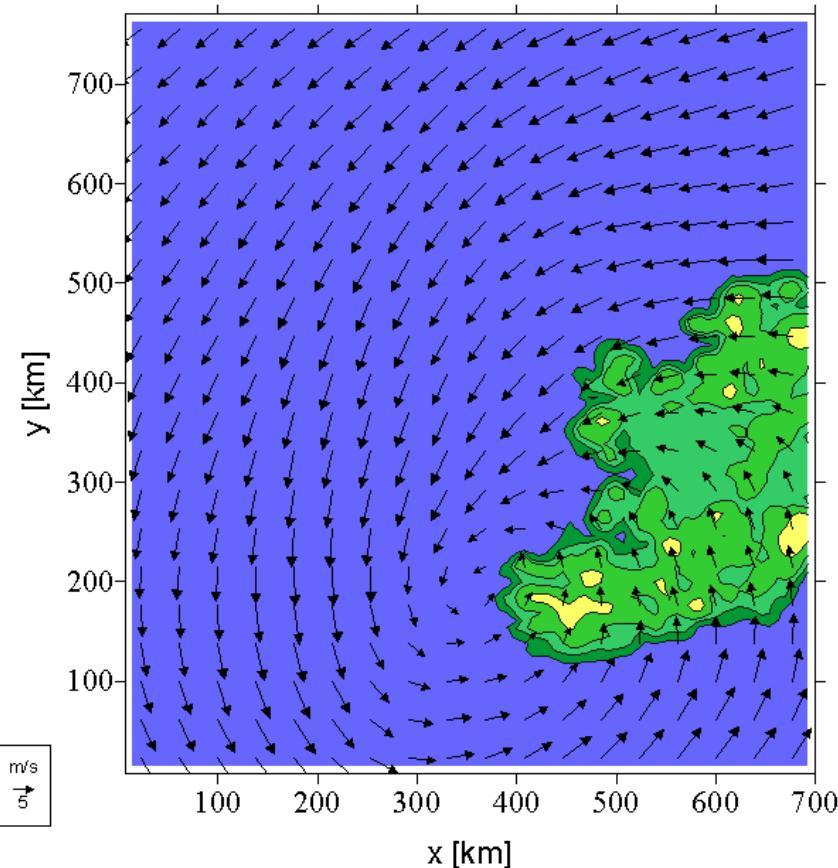
◆ Monahan *et al.* [1986]: $\frac{dF_0}{dr_{80}} = 1.373 \cdot U_{10}^{3.41} \cdot r_{80}^{-3} \left(1 + 0.057 r_{80}^{1.05}\right) \cdot 10^{1.19e^{-B^2}}$

◆ Smith *et al.* [1993]: $\frac{dF_0}{dr_{80}} = \sum_{i=1,2} A_i \exp\left(-f_i \left(\ln \frac{r_{80}}{r_{0i}}\right)^2\right)$, $A_i = A_i(U_{10})$

Lewis and Schwartz [2005]: $r_{RH} = (r_d) \left(\frac{4.0}{3.7}\right) \left(\frac{2.0 - RH}{1 - RH}\right)^{\frac{1}{3}}$

Wind and Number Density

Wind , 28.05.05, 00 UTC



number concentration, 28.05.05, 01 UTC

